

**Università di Trieste**  
**LAUREA MAGISTRALE IN**  
**GEOSCIENZE**  
**Curriculum Geofisico**  
**Curriculum Geologico Ambientale**

**Anno accademico 2018 – 2019**

**Geologia Marina**

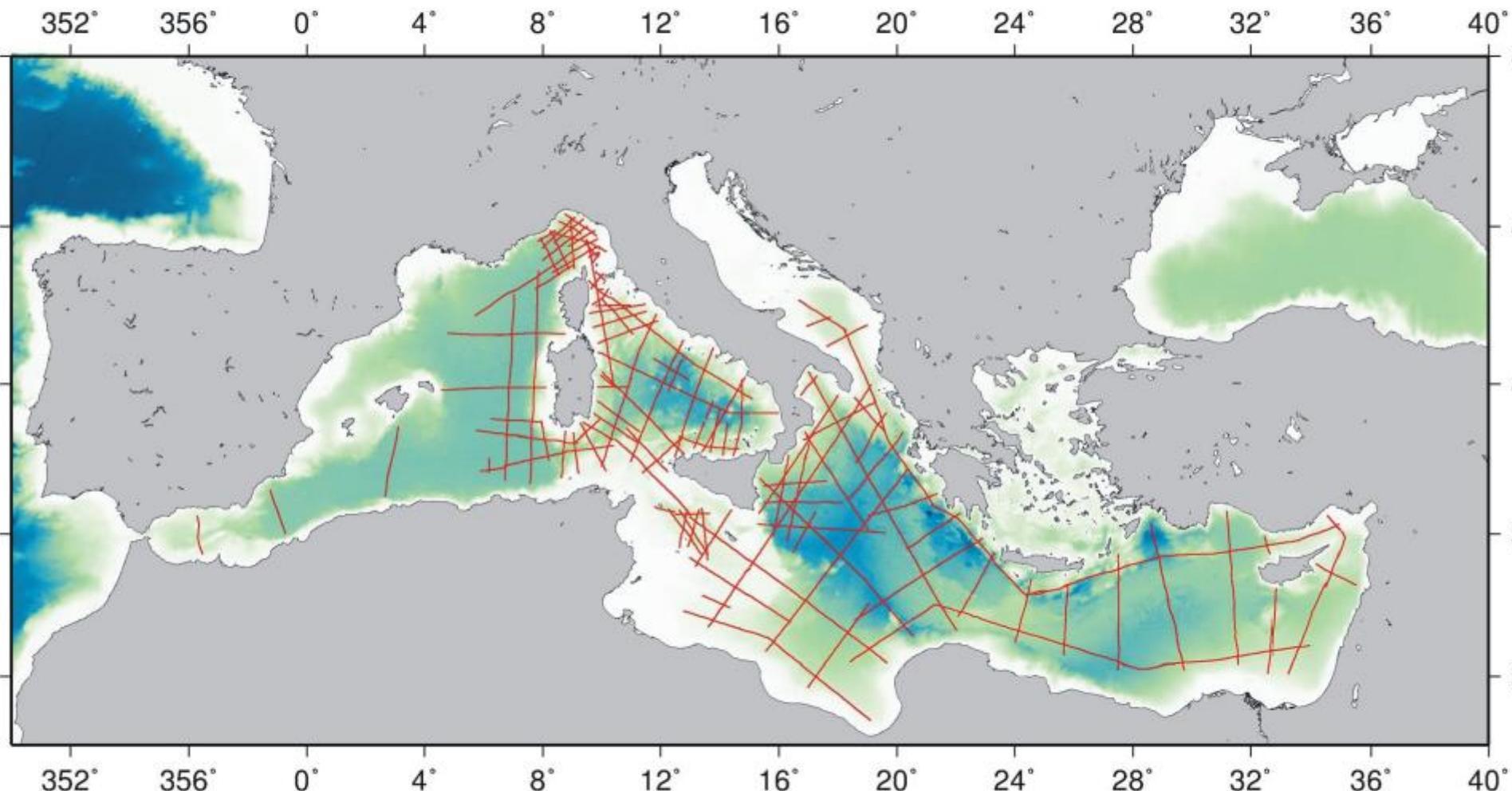
**Modulo 5.2**

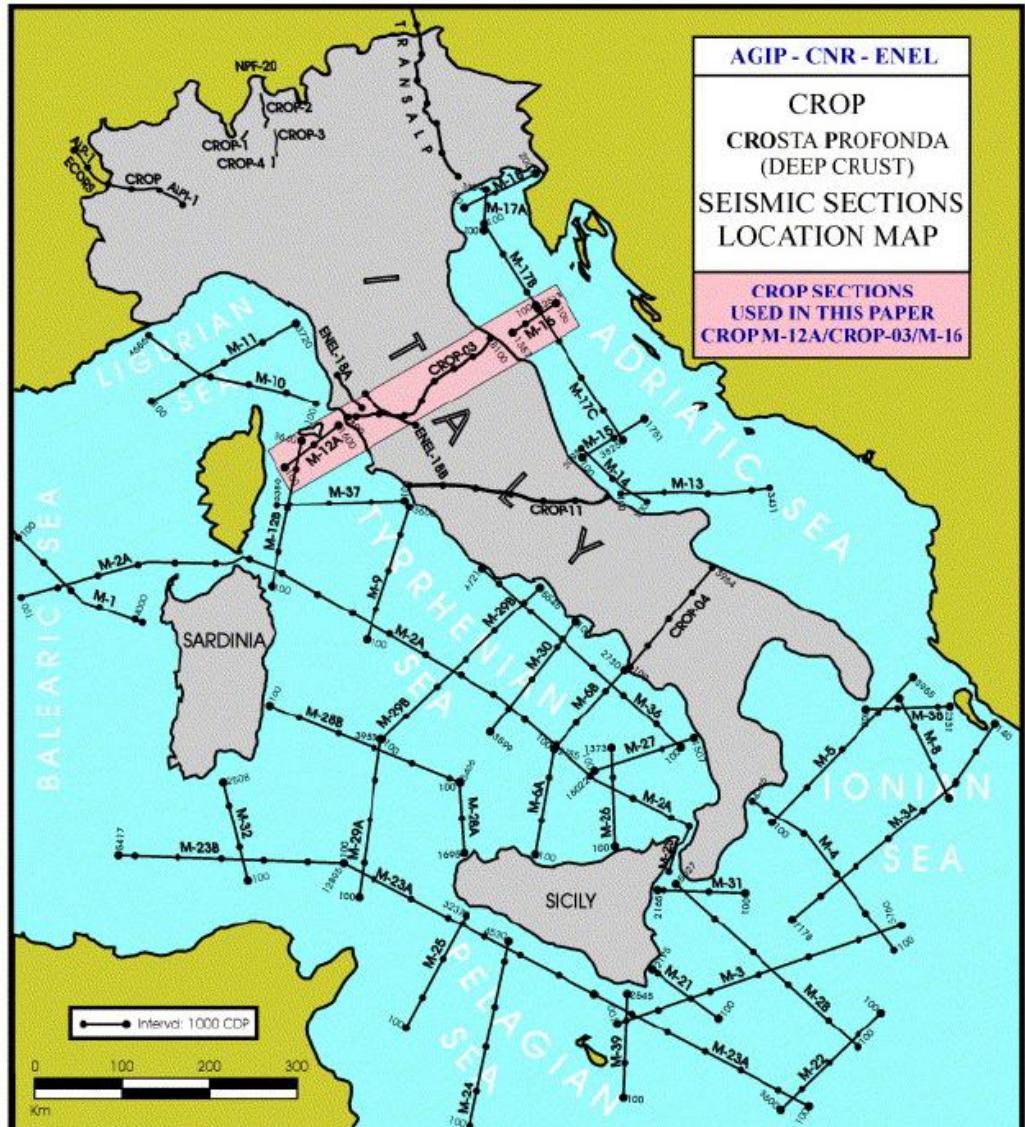
**Mediterraneo 2 (Alboran, Balearico e Ionio)**  
**Part 2**

Docente  
**Silvia Cereamicola**  
(sceramicola@inogs.it)

# MS map

(seismic profiles collected from 1968 to 1982)





# CROP map

(Seismic profiles collected both onshore and offshore)

**FRENCH (ECORS) - ITALIAN (CROP) COOPERATION**  
-WESTERN ALPS (ALP-I/ALPI-I)  
-BALEARIC SEA (GULF OF LYON/SARDINIA, M-1)

**SWISS (NPF-20) - ITALIAN (CROP) COOPERATION**  
-CENTRAL ALPS (CROP-1, 2, 3, 4 AND NPF-20  
CONTINUATION TO NORTH)

**GERMAN (DEKORP) - AUSTRIAN (OEKORP) -  
ITALIAN (CROP) COOPERATION**  
-EASTERN ALPS (TRANSALP)

**GREEK - ITALIAN (CROP) COOPERATION**  
-NE IONIAN SEA (EAST - EXTREMITIES OF M-34 & M-38)





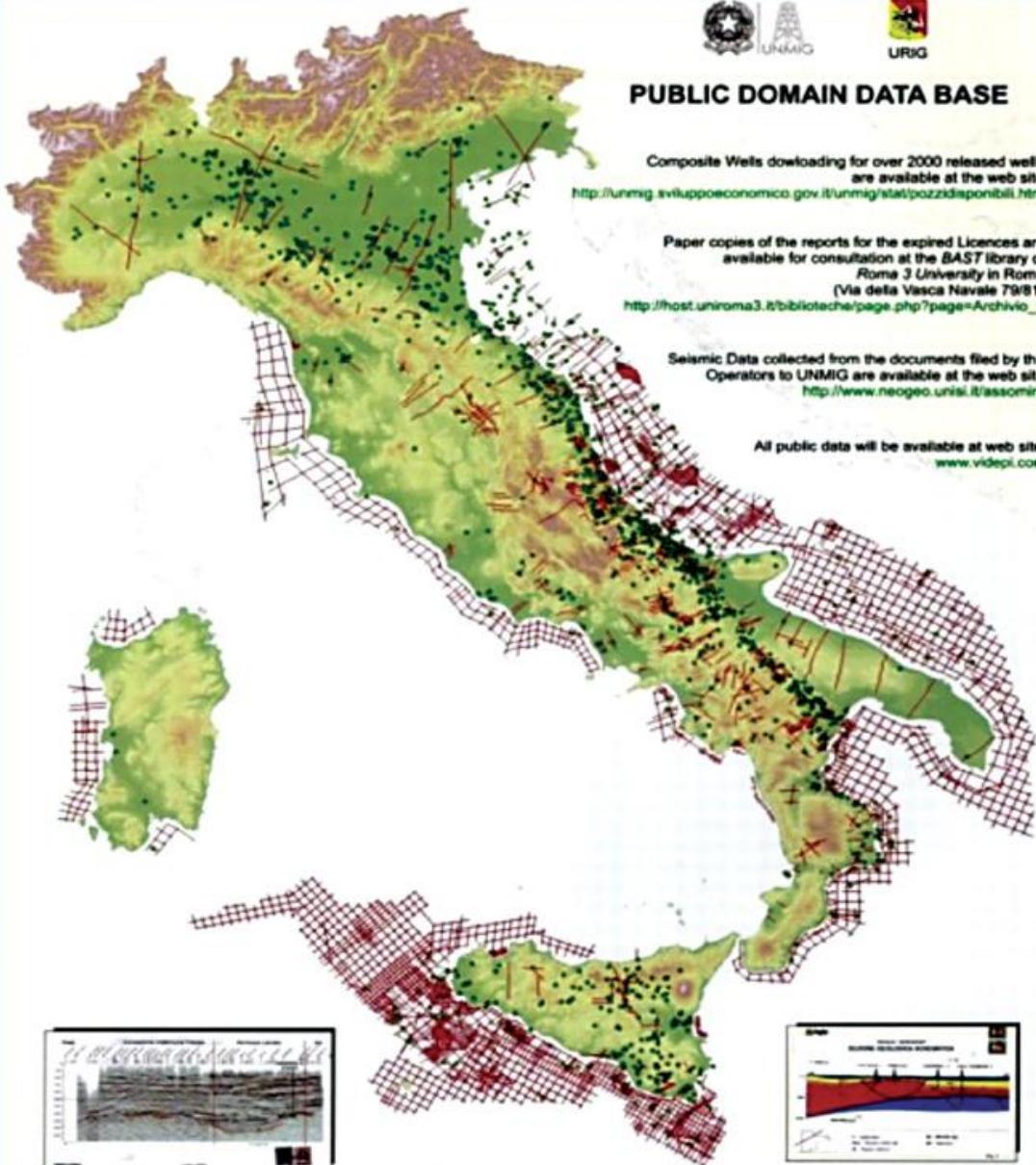
## PUBLIC DOMAIN DATA BASE

Composite Wells dowloading for over 2000 released wells  
are available at the web site  
<http://unmig.sviluppoeconomico.gov.it/unmig/stat/pozzi/disponibili.htm>

Paper copies of the reports for the expired Licences are  
available for consultation at the *BAST* library of  
Roma 3 University in Rome  
(Via della Vasca Navale 79/81)  
[http://host.uninoma3.it/biblioteche/page.php?page=Archivio\\_9](http://host.uninoma3.it/biblioteche/page.php?page=Archivio_9)

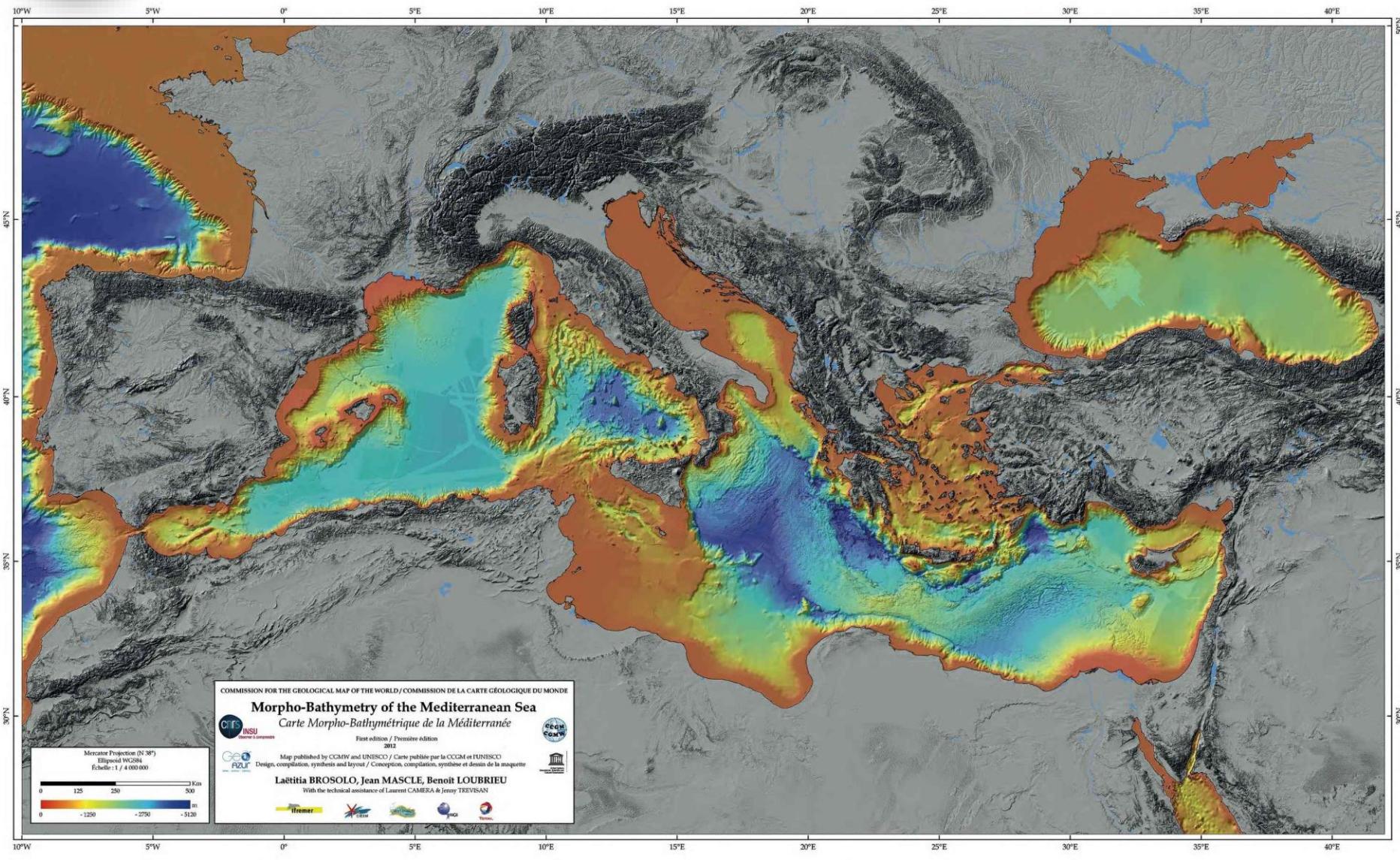
Sismic Data collected from the documents filed by the  
Operators to UNMIG are available at the web site  
<http://www.neogeob.uni.it/assomin/>

All public data will be available at web site:  
[www.videpi.com](http://www.videpi.com)





# CCGM Morpho-Bathymetry of the Mediterranean sea



**Légende**

La cartographie systématique par bathymétrie multibancs du bassin méditerranéen et de ses marges continentales, initiée dès le milieu des années 1990, a profondément modifiée notre connaissance de la morphologie de la Mer Méditerranée ainsi que des différents processus géologiques actifs (sédimentaires, tectoniques, volcaniques, bio-chimiques) qui contribuent au modèle du fond de mer.

**Legend**

The systematic swath bathymetry of the Mediterranean Sea in the nineties. This mapping has modified our understanding of the morphology of the Mediterranean Sea and of the various active geological processes (sedimentary, tectonic, geochemical processes) which control the seafloor morphology.

### Origine des dommages

Cette carte a été construite à partir de différents modèles numériques de terrain.

• Le MINT de Brest, à 500m, provient d'une coopération anglo-baltique de la Méditerranée publiée par Louchev, B., Maclacé, et Medema (2009) : *Baltic-bathymetry of the Mediterranean Sea*, CIESM/FIRER publication.

• Le MINT de Brest a été fondé sur la base des données de bathymétrie multilâser acquises depuis 1996 à 2000 m de profondeur et de sondages acoustiques et laboratoires de recherches (<http://www.ciesm.org/>).

• Ce MINT a été complété par grilles de bathymétrie multilâser mesurées précédemment par l'IFREMER, du SHOM, de la CGRAC, de l'OCGS (Trente), de l'UBA (Bordeaux), du CRG (Université de Bambouk), ainsi qu'à larges intervalles par les sondages acoustiques de l'IFREMER, du CRG, de l'OCGS, de l'UBA, du CRG, de l'OCGS. Ces dernières données complètent celles du CRG sur le site Web de la page européenne [www.mint-eu.org](http://www.mint-eu.org).

• Pour le golfe de Cadiz la carte a été fondée sur les données de bathymétrie multilâser acquises par l'IFREMER, du CRG, de l'OCGS, de l'UBA, du CRG, de l'OCGS.

• Pour le Golfe de Gascogne à 1000m, extrait de Shallow, J.C., Louchev, S., et al. (2004) : *Carte bathymétrique de l'atlantique nord-est et du Golfe de Gascogne*, CIESM/FIRER publication.

• Pour le Golfe de Gascogne à 1000m, extrait de Shallow, J.C., Louchev, S., et al. (2004) : *Carte bathymétrique de l'atlantique nord-est et du Golfe de Gascogne : implications écopalentielles*, Bull. Soc. Geol. Fr., a été utilisée.

#### Data sources

This map results from the compilation of the following DTMs and complementary data:

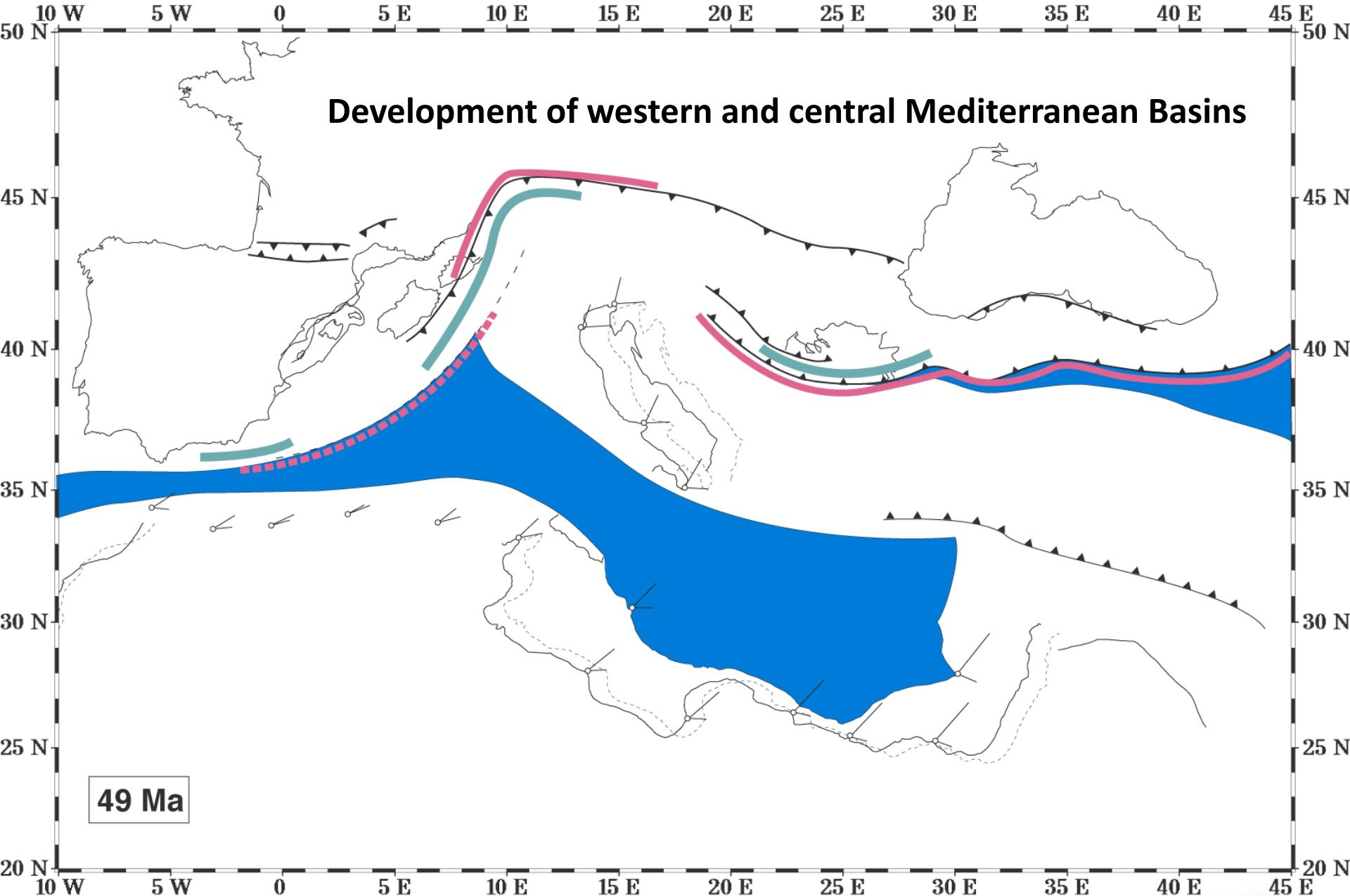
The basic TDS is used for the Morpho-bathymetric compilation of the Mediterranean Sea published by Lebedeva, L., Mackie, J. and Melisup Group (2007). *Morpho-bathymetry of the Mediterranean Sea - CIBSM/Werner special publication.*  
 The bathymetry results derived from various TDSs with swath bathymetry data sets recorded between 1995-2007 by several European oceanographic laboratories and Institutes ([see <http://www.ciscom.org/>](http://www.ciscom.org/)).  
 This TDS has been compiled by unpublished swath bathymetry data sets from IFREMER, SHOM, GEOMAR, OGZ-Tenete, EMAP-Europe, CEMRAC-Hamburg and, off Egypt, by a few 3D seismic data first arrivals from industry (BP-Egypt, ENI, ENI).

A few complementary data have been downloaded from the EMODNET European project portal ([www.emodnet-bathymetry.eu/](http://www.emodnet-bathymetry.eu/))

- For the Gulf of Cadiz the bathymetry data sets from IFREMER in 500 m depth swath data published by Zeffane, N., Garcia, E., et al., (2009) - Bathymetry of the Gulf of Cadiz, North-East Atlantic: the Spain-Moroccan compilation - Earth Science Frontiers, 18(1), 1-12.
- For the Black Sea, the swath bathymetry data sets were obtained during the European project ASSEMBLEAGE.

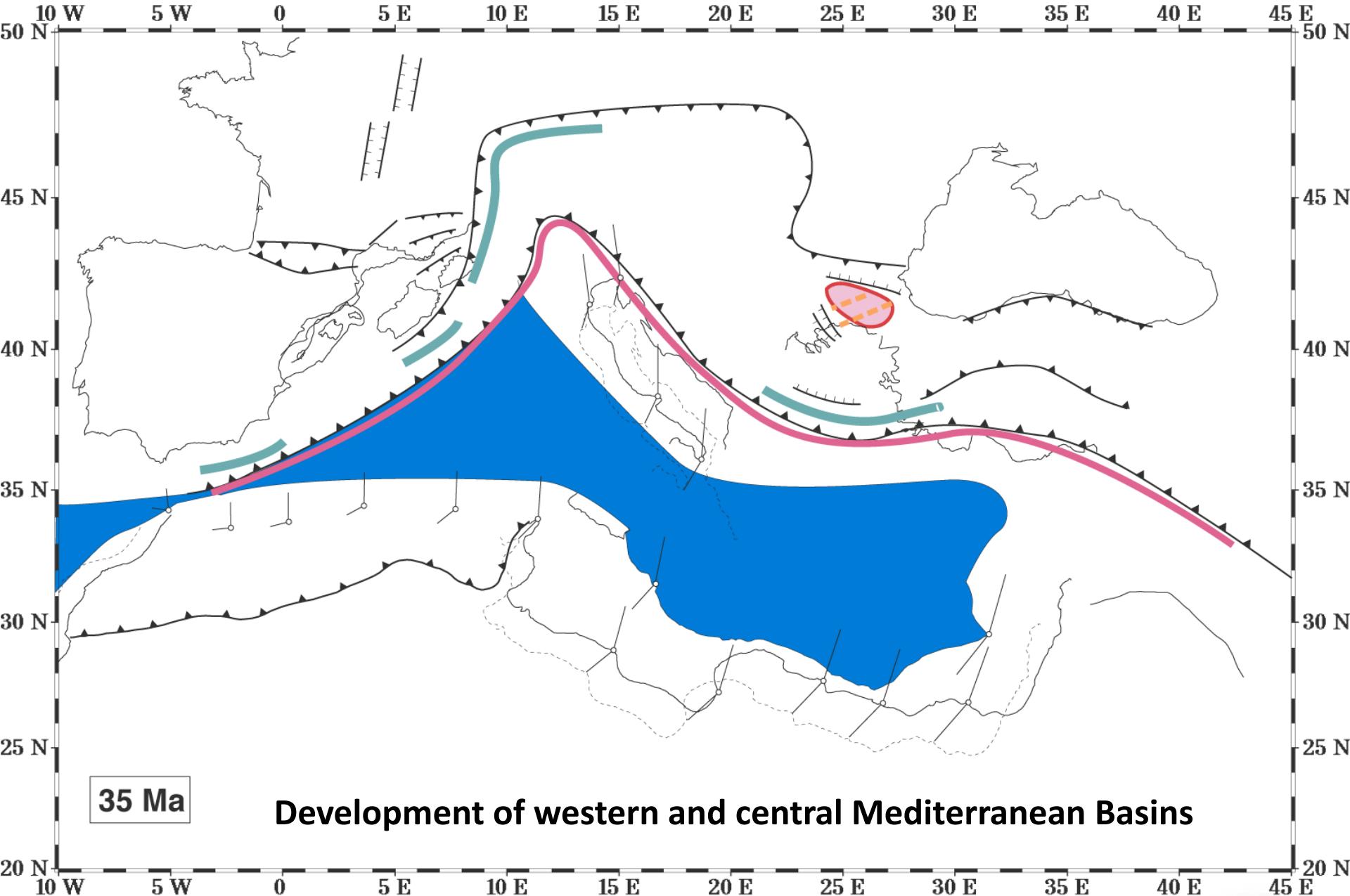
- The present-day geological configuration of the Mediterranean region is the result of the creation and ensuing consumption of two major oceanic basins: the Paleotethys and the Neotethys.
- The overall tectonic regime was (and it is) the regime of prolonged interaction between the Eurasian and the African-Arabian plates.
- The Mediterranean domain provides a present-day geodynamic analog for the final stages of a continent-continent collisional orogeny. Over this area, the oceanic lithospheric domains originally present between the Eurasian and African-Arabian plates have been subducted and partially obducted (ophiolitic terranes), except for the Ionian basin and the south-eastern Mediterranean.
- The modern marine basins of the Mediterranean Sea are variably floored by: (i) remnants of the Tethyan oceanic domains (Ionian, E Mediterranean); (ii) Neogene oceanic crust (Algero-Provençal basin and Tyrrhenian Sea); (iii) extended continental lithosphere (Alboran Sea, Valencia Trough, Aegean Sea), and (iv) thick continental lithosphere (Adriatic Sea).





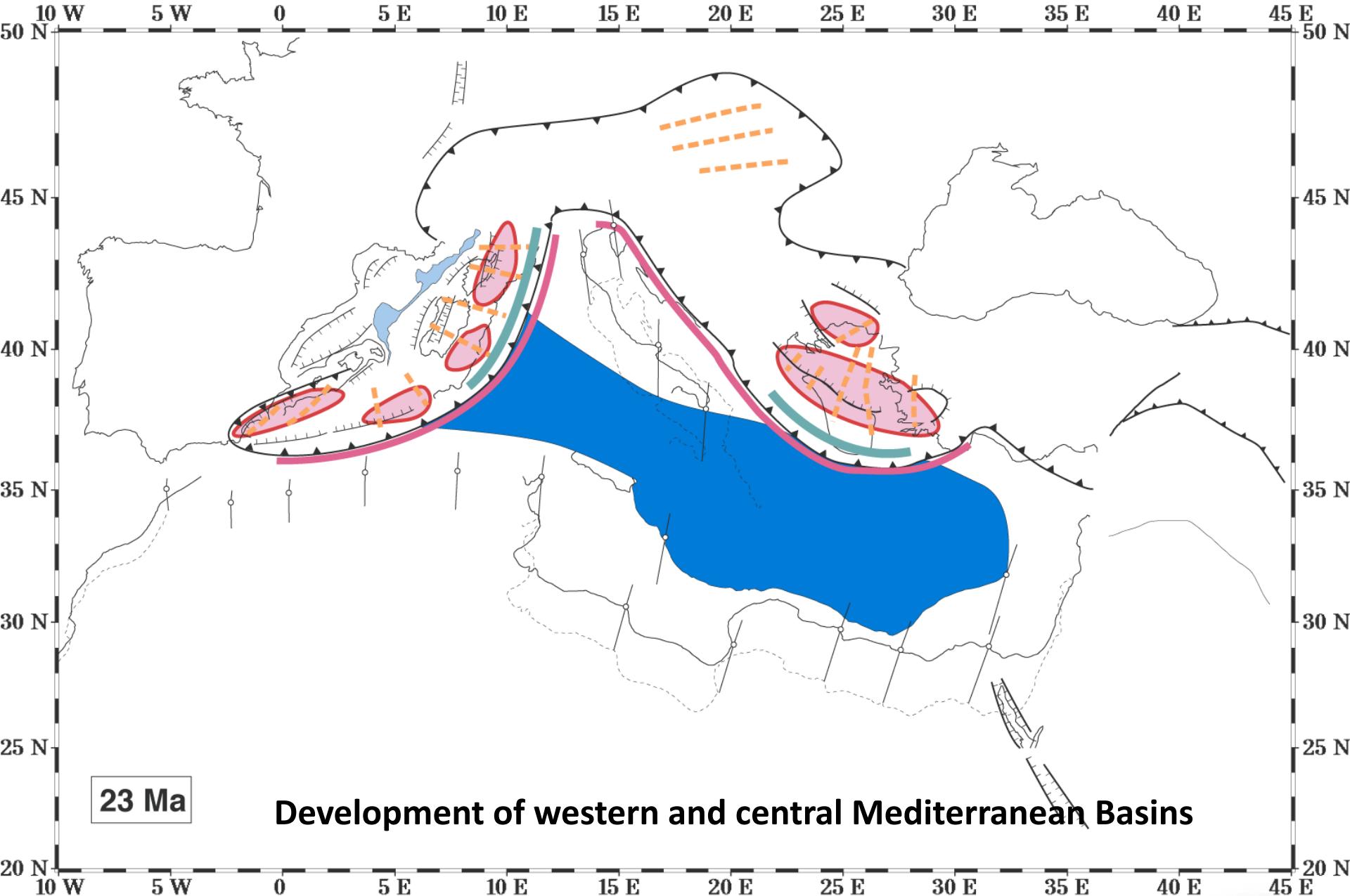
Jolivet et al., 2003., Kinematic data after Dewey et al.





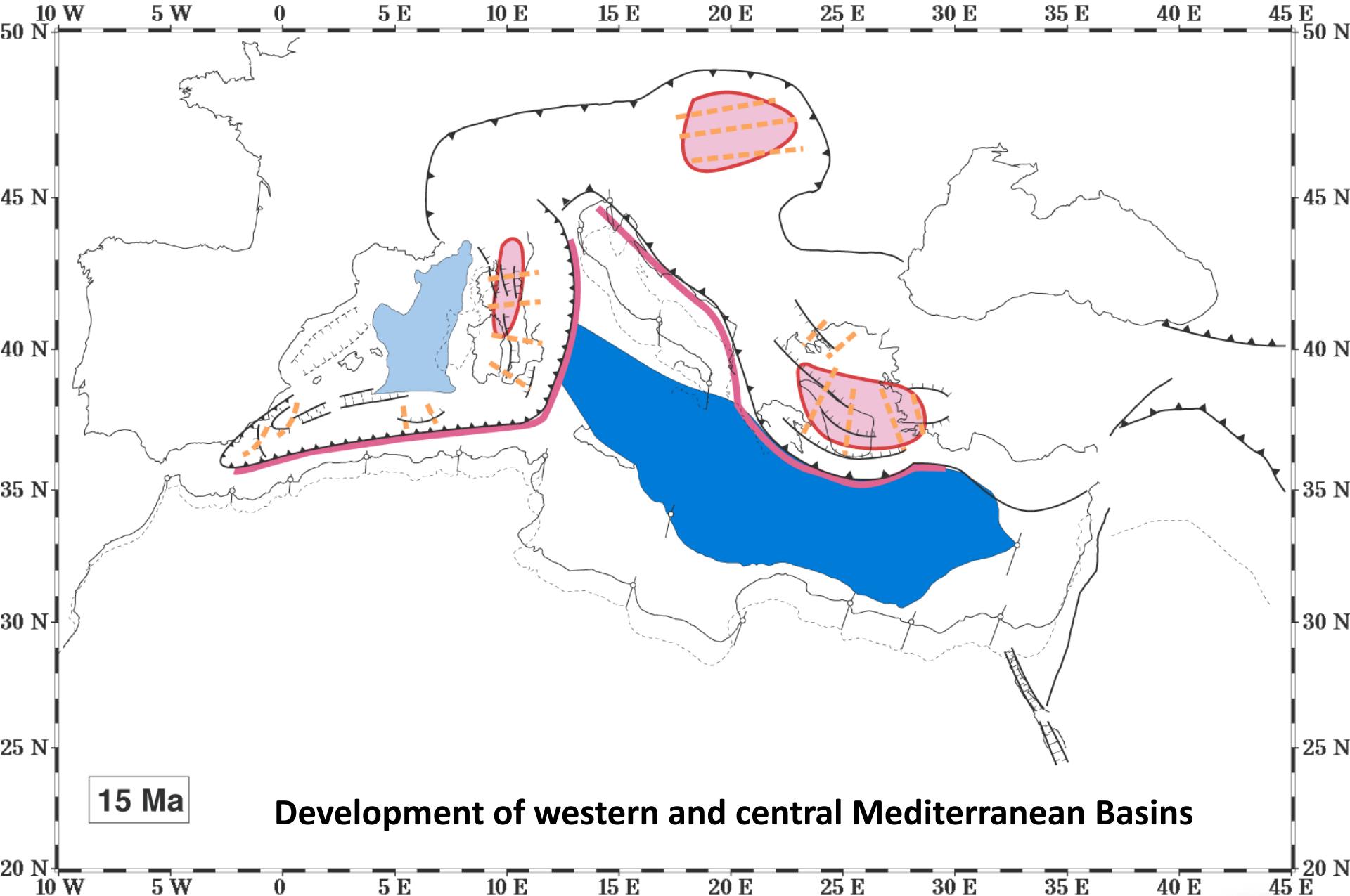
Jolivet et al., 2003., Kinematic data after Dewey et al.





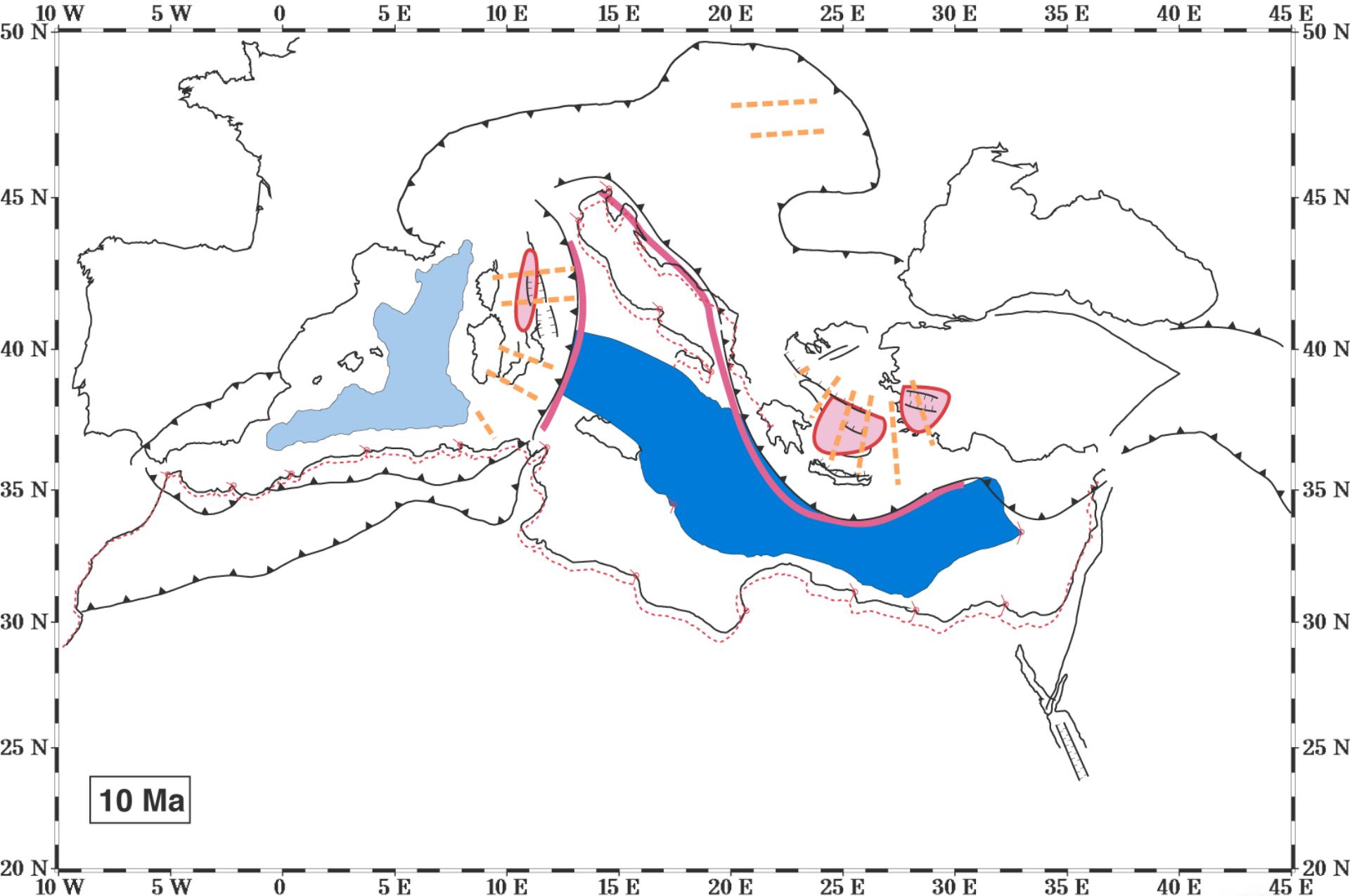
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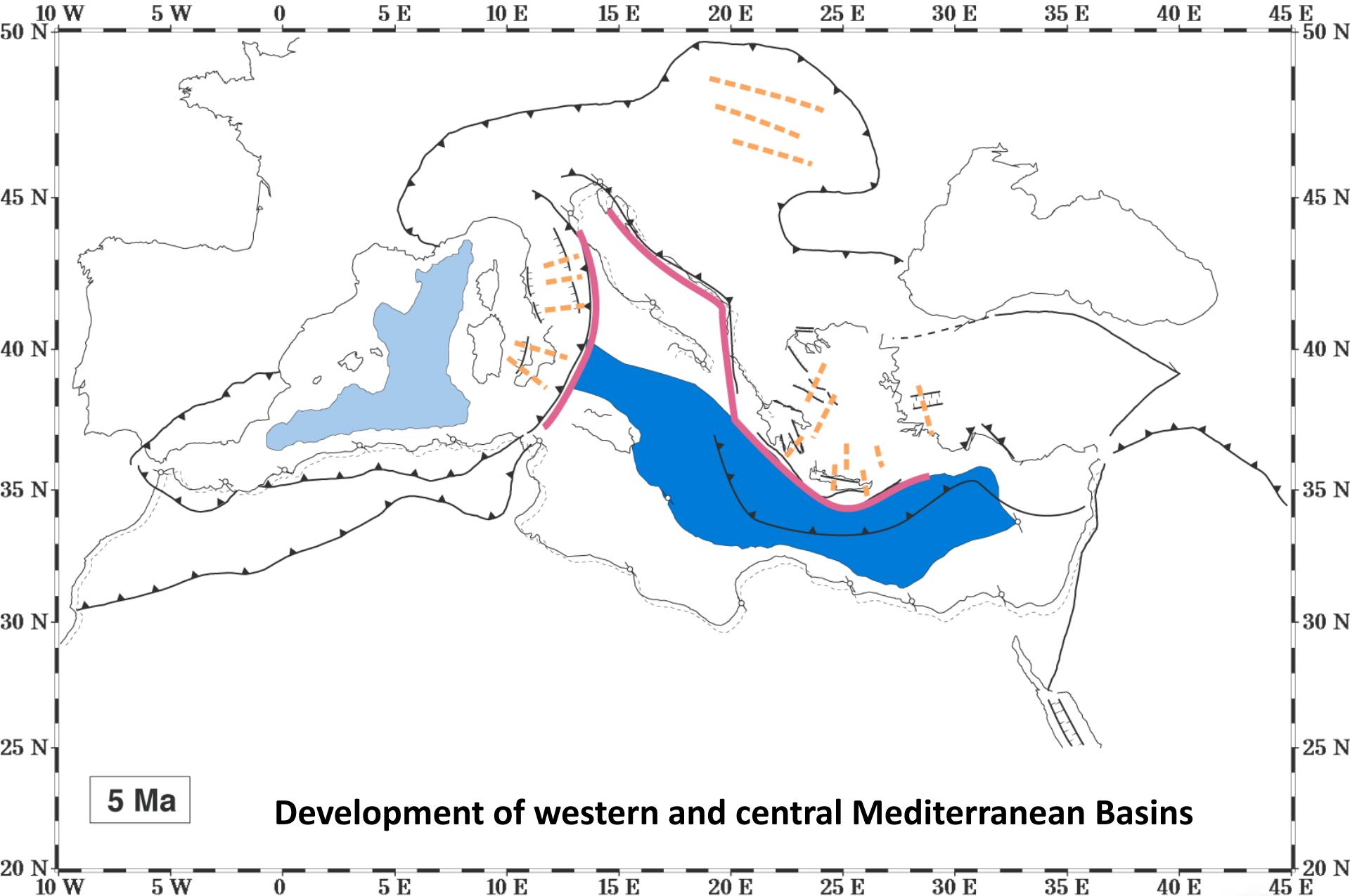
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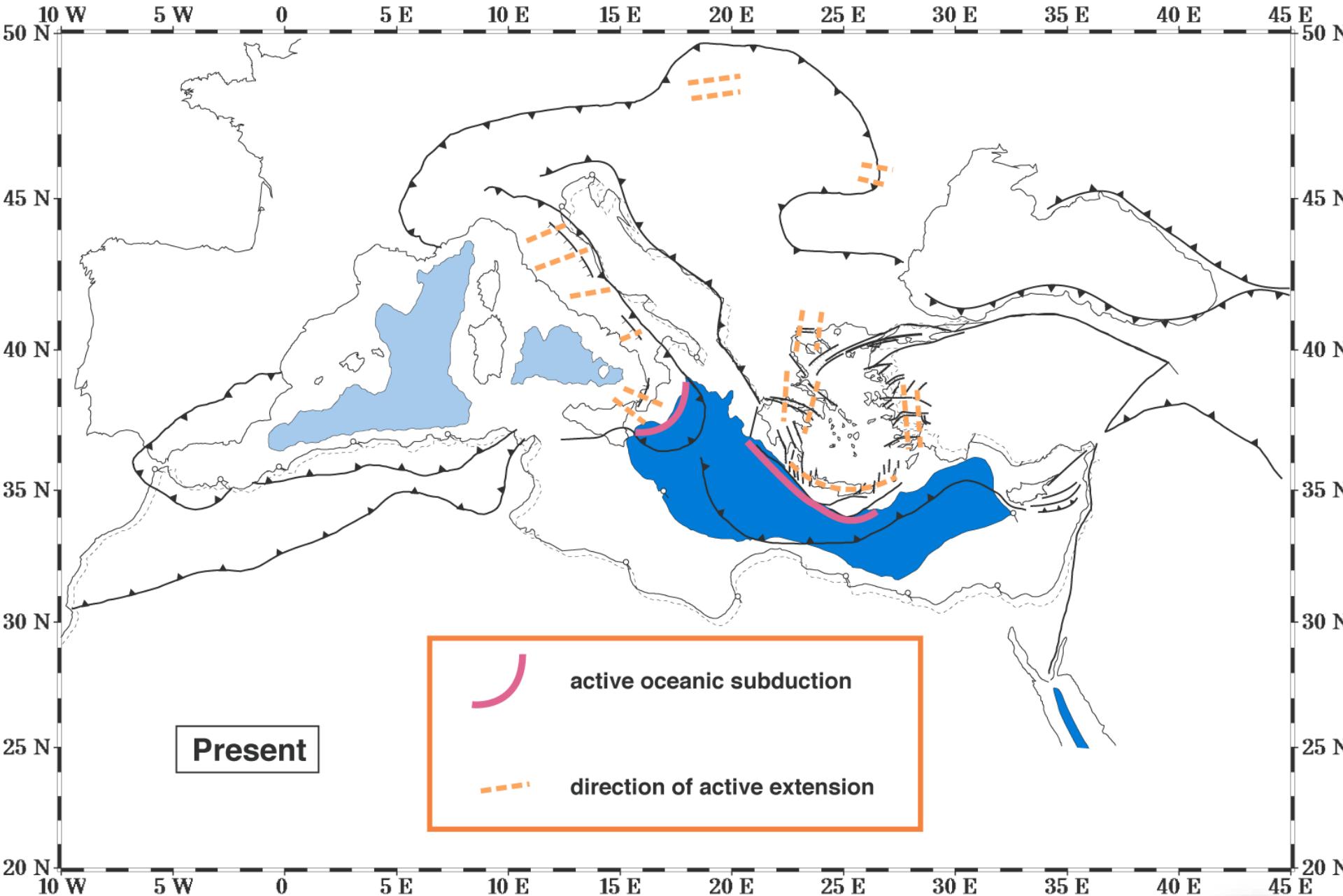
Jolivet et al., 2003., Kinematic data after Dewey et al.





Jolivet et al., 2003., Kinematic data after Dewey et al.

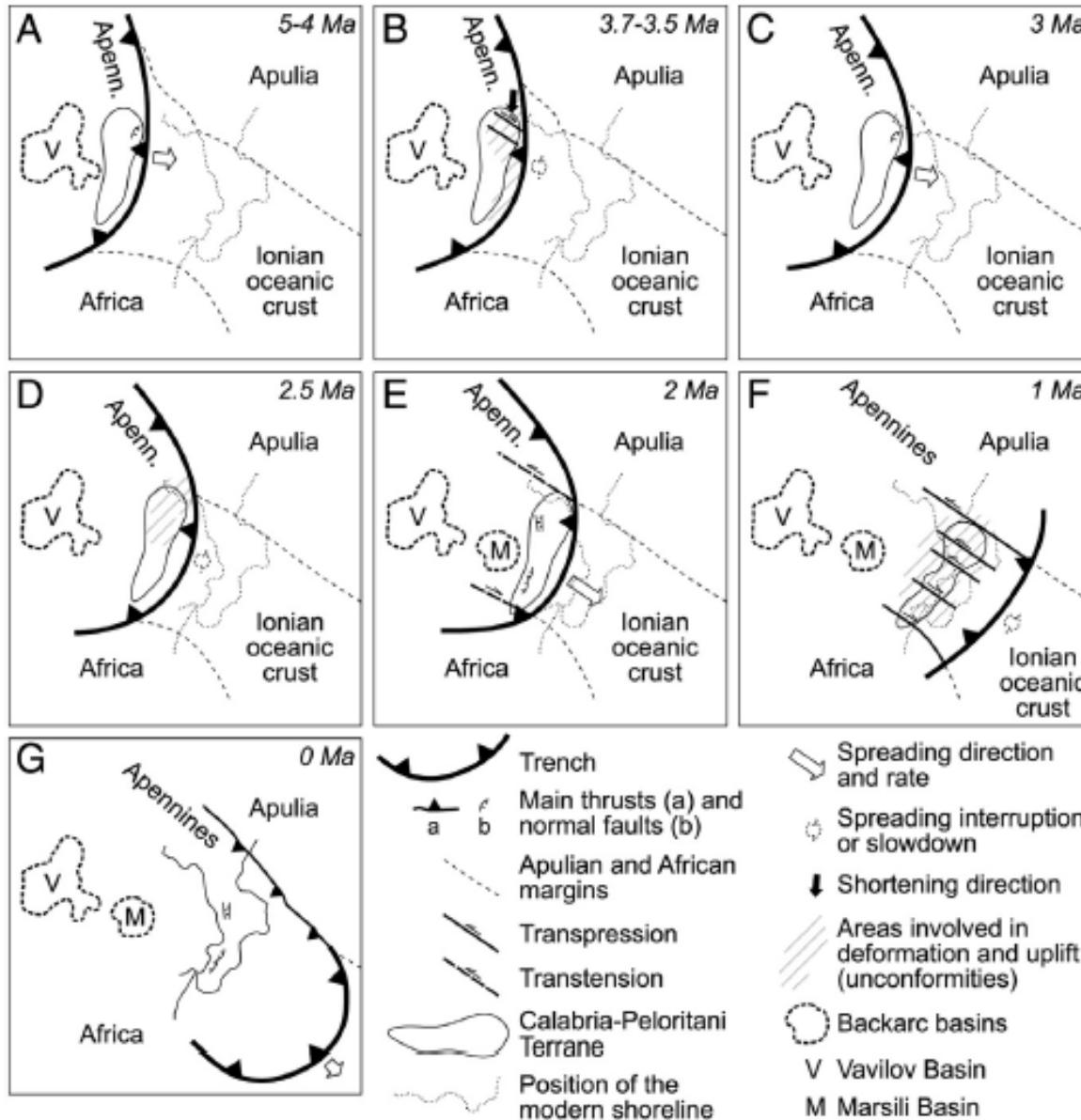


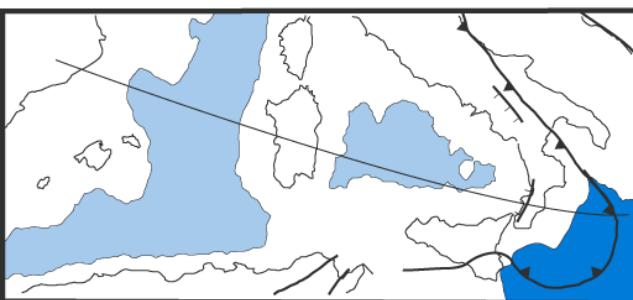


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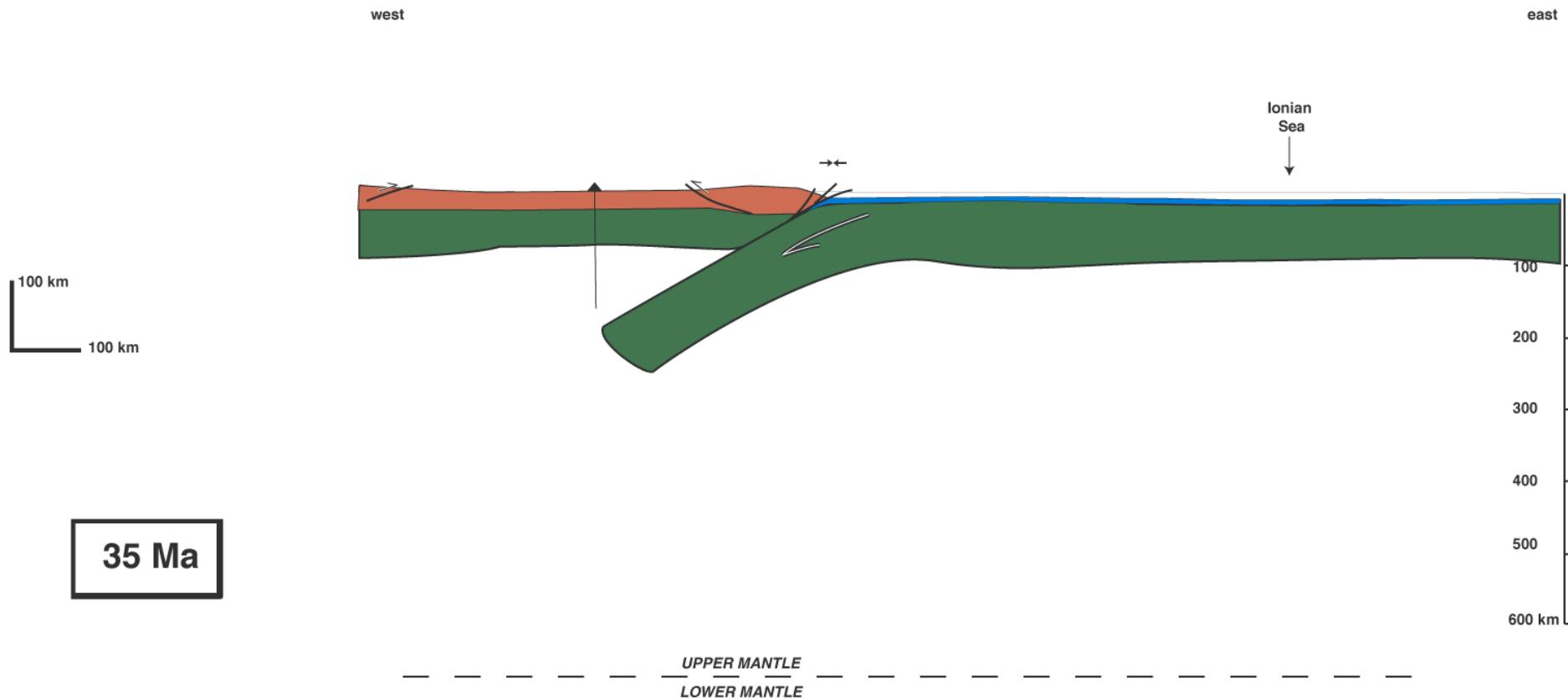


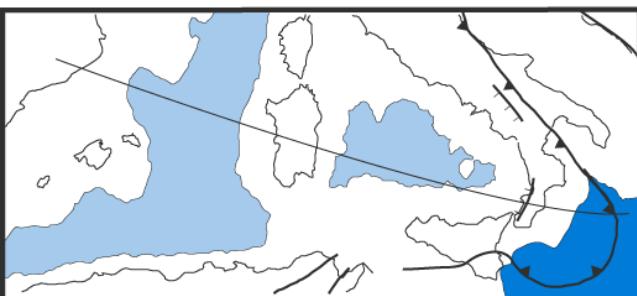
# The Calabrian subduction



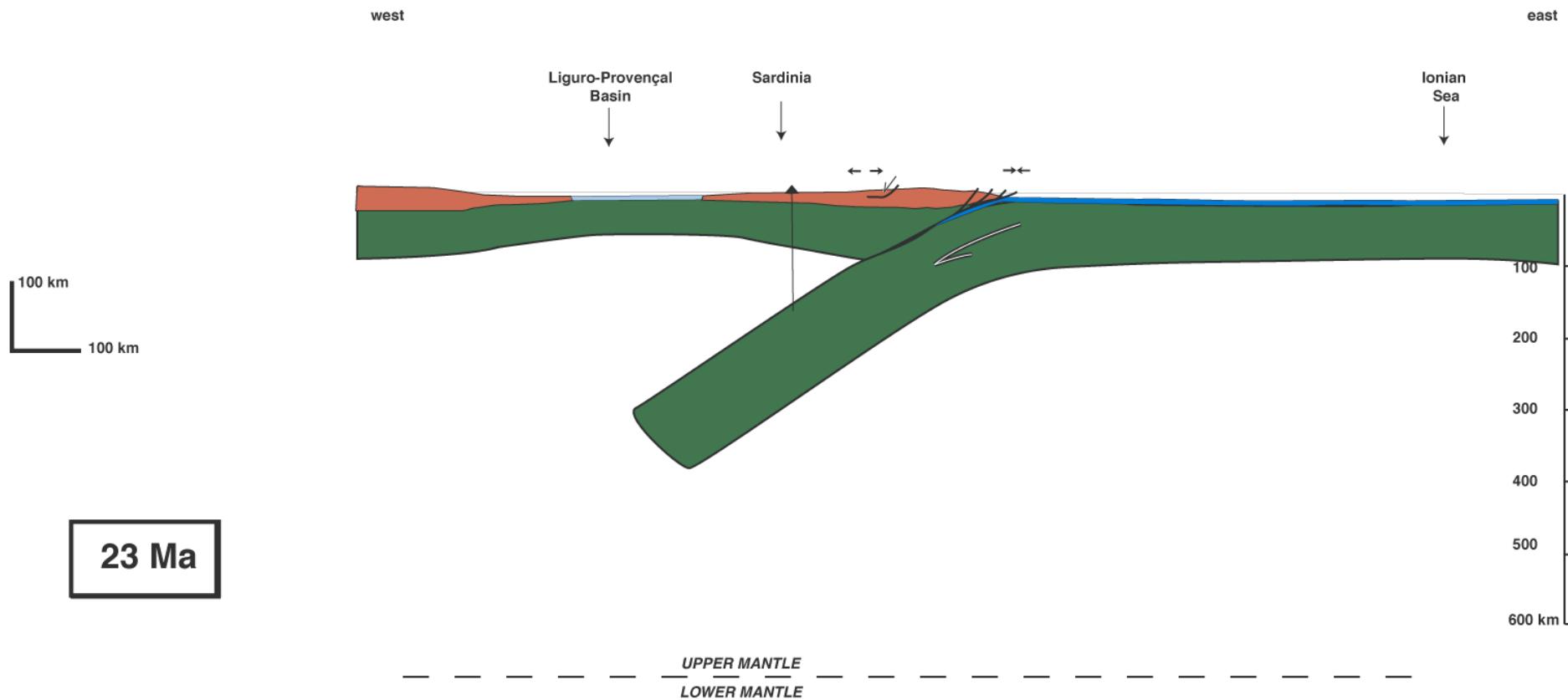


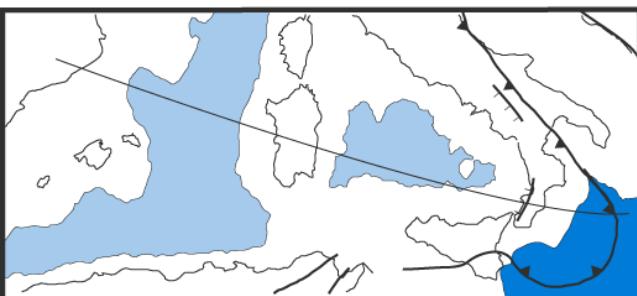
# The subduction of the Ionian Sea



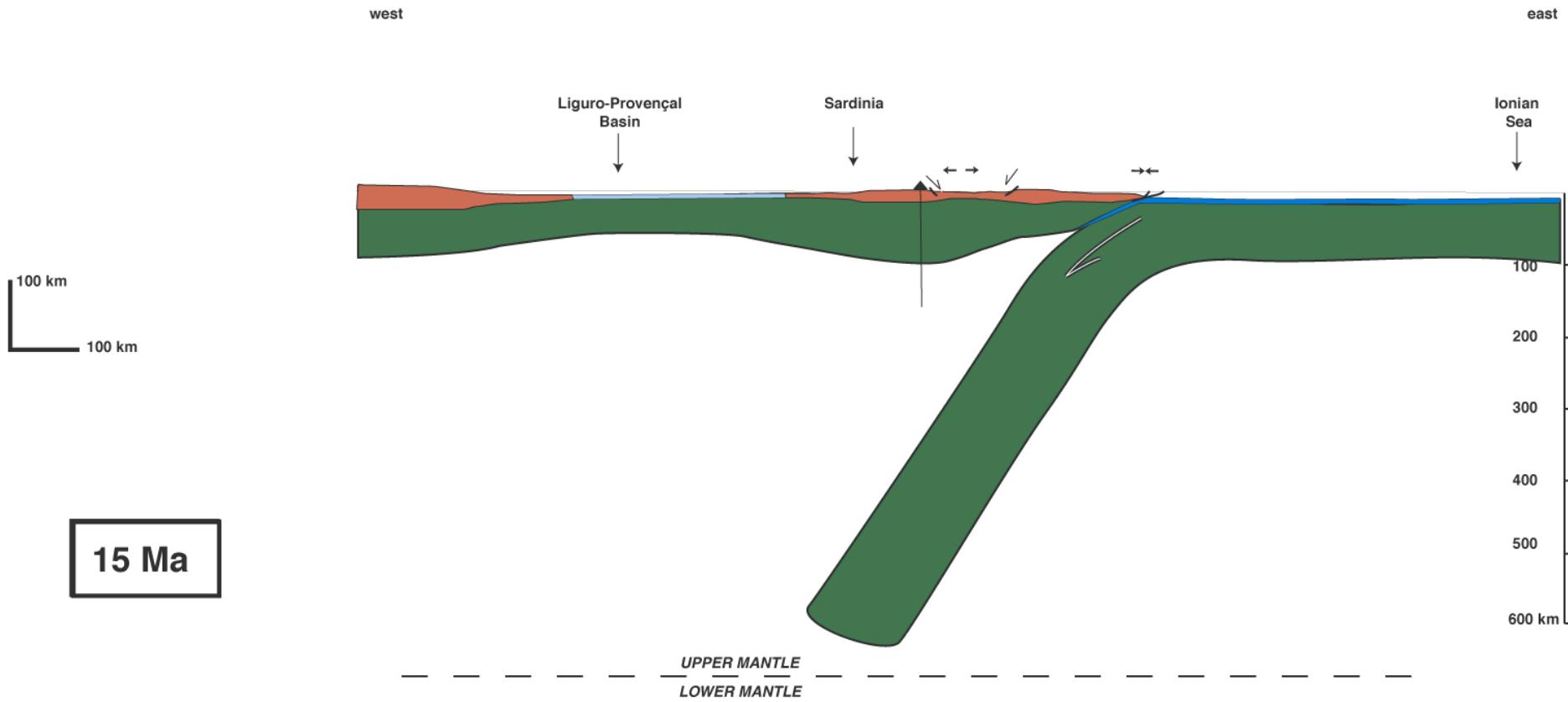


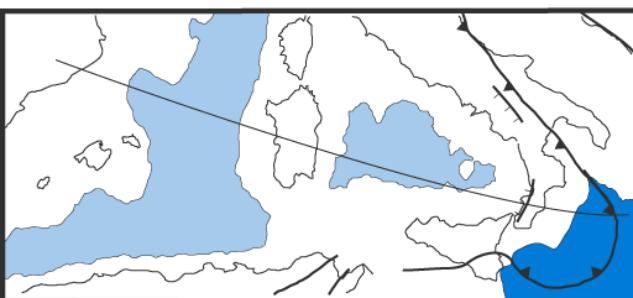
# The subduction of the Ionian Sea



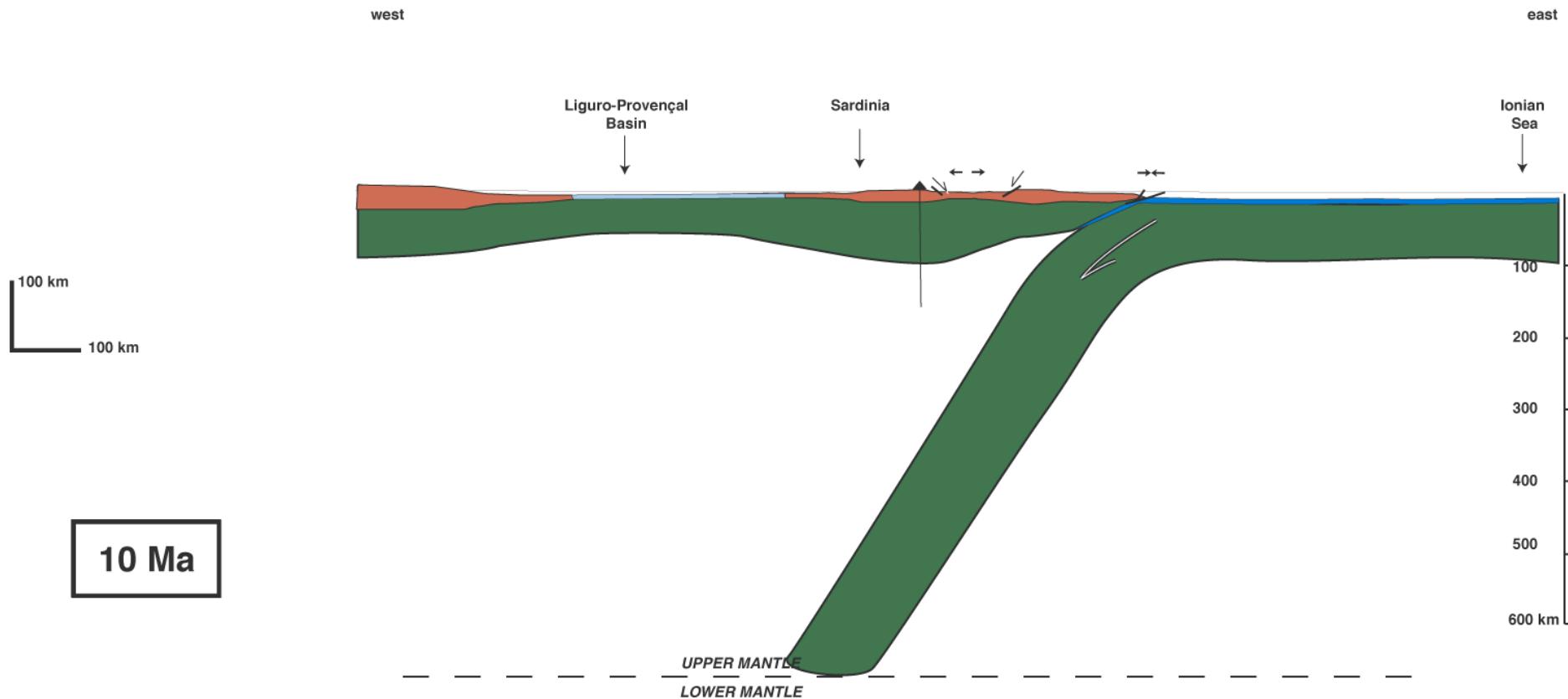


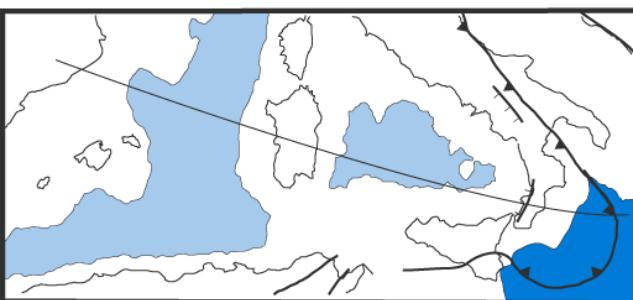
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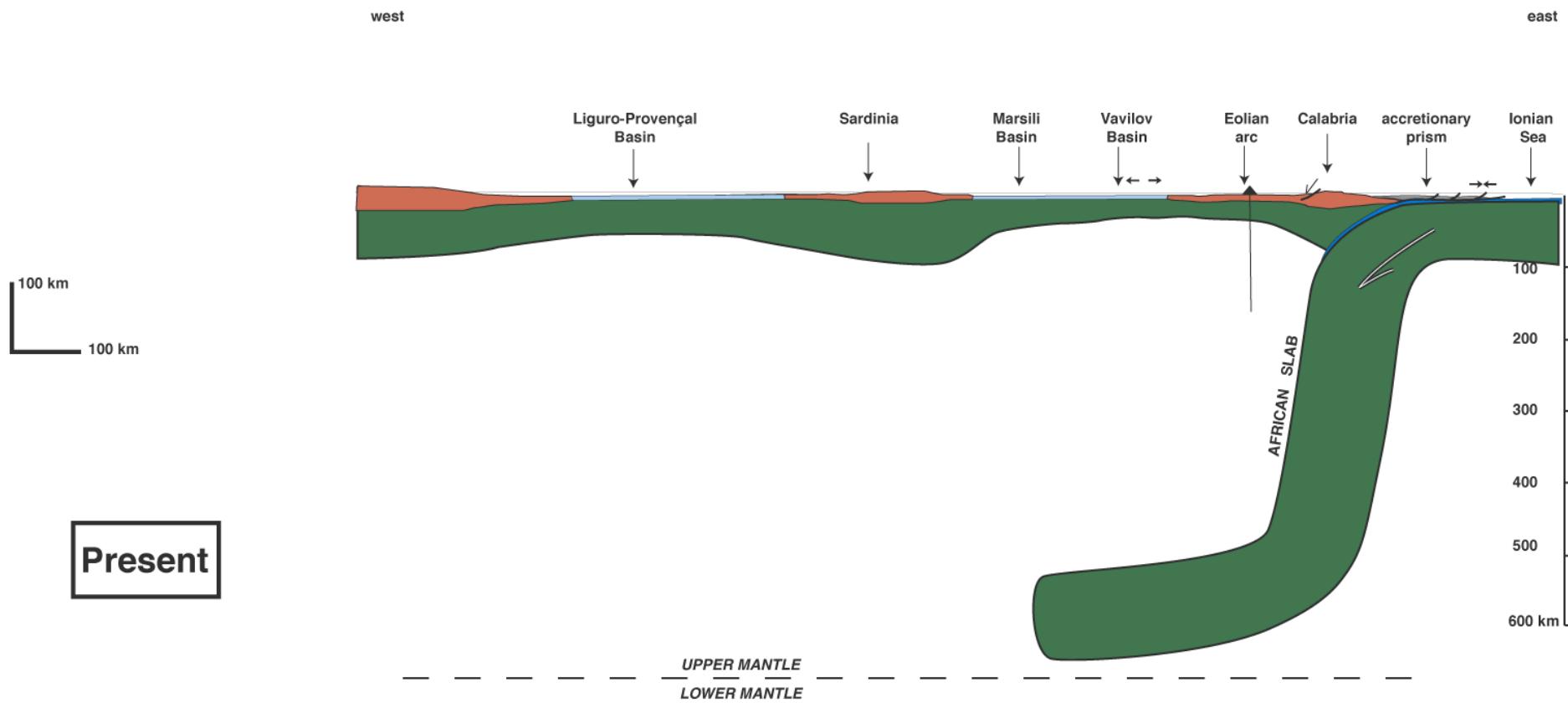


# The subduction of the Ionian Sea



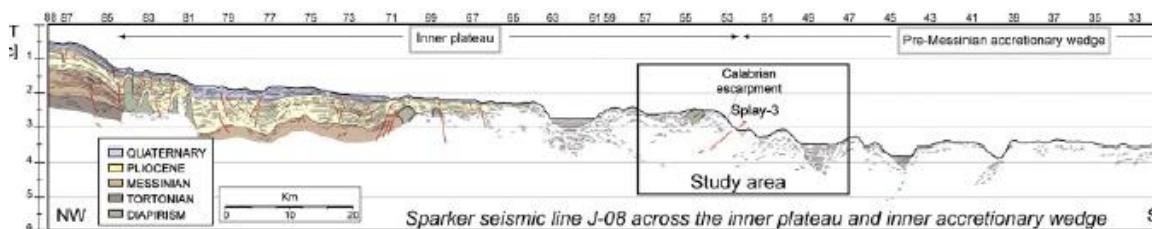
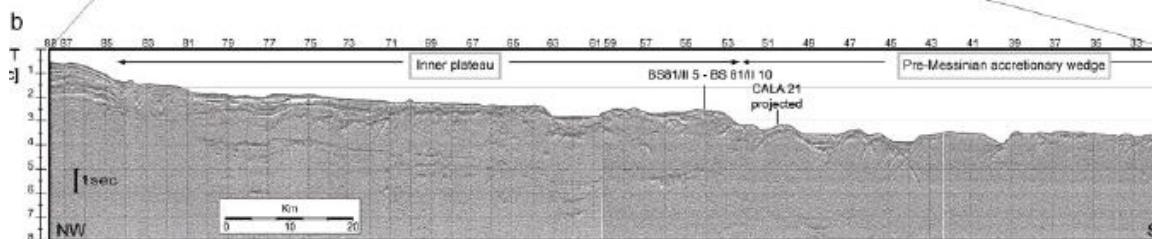
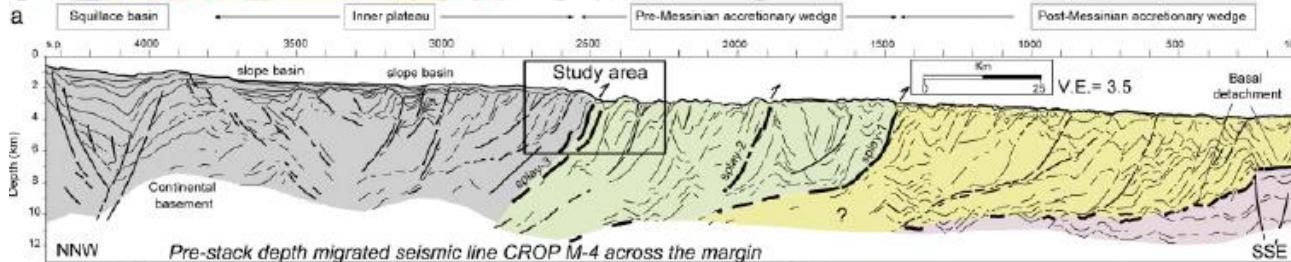
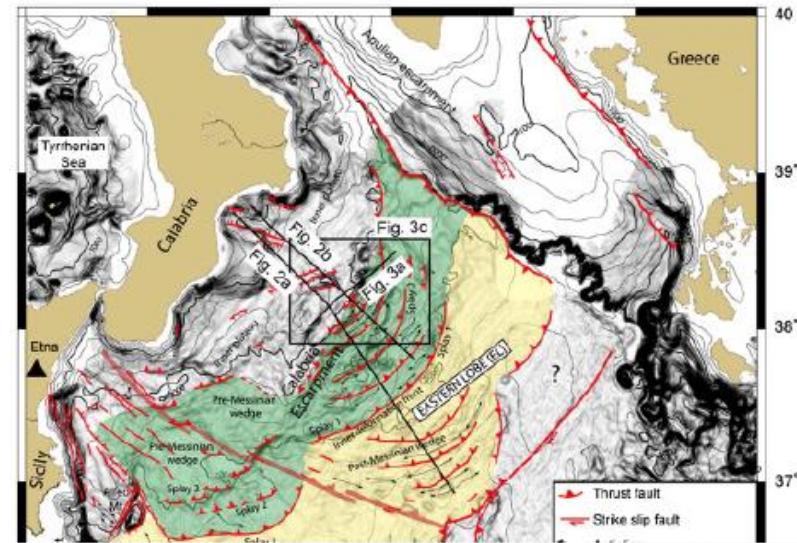


# The subduction of the Ionian Sea



# Calabrian accretionary Prism

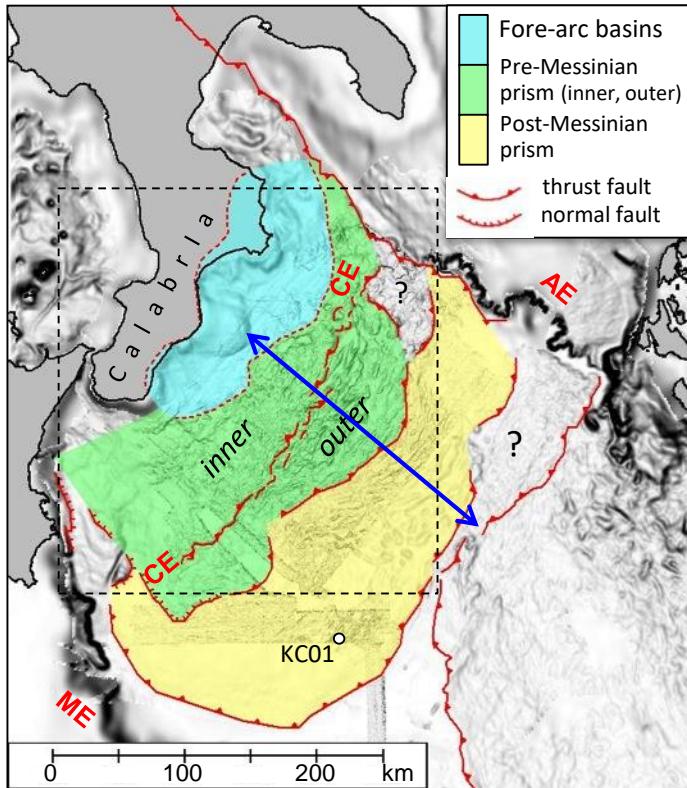
Deformation is related to an imbricate fan within the post-Messinian salt-bearing accretionary wedge (yellow domain), out-of-sequence thrust faults in the pre-Messinian wedge (green domain) and normal faults in the Inner plateau (gray domain).



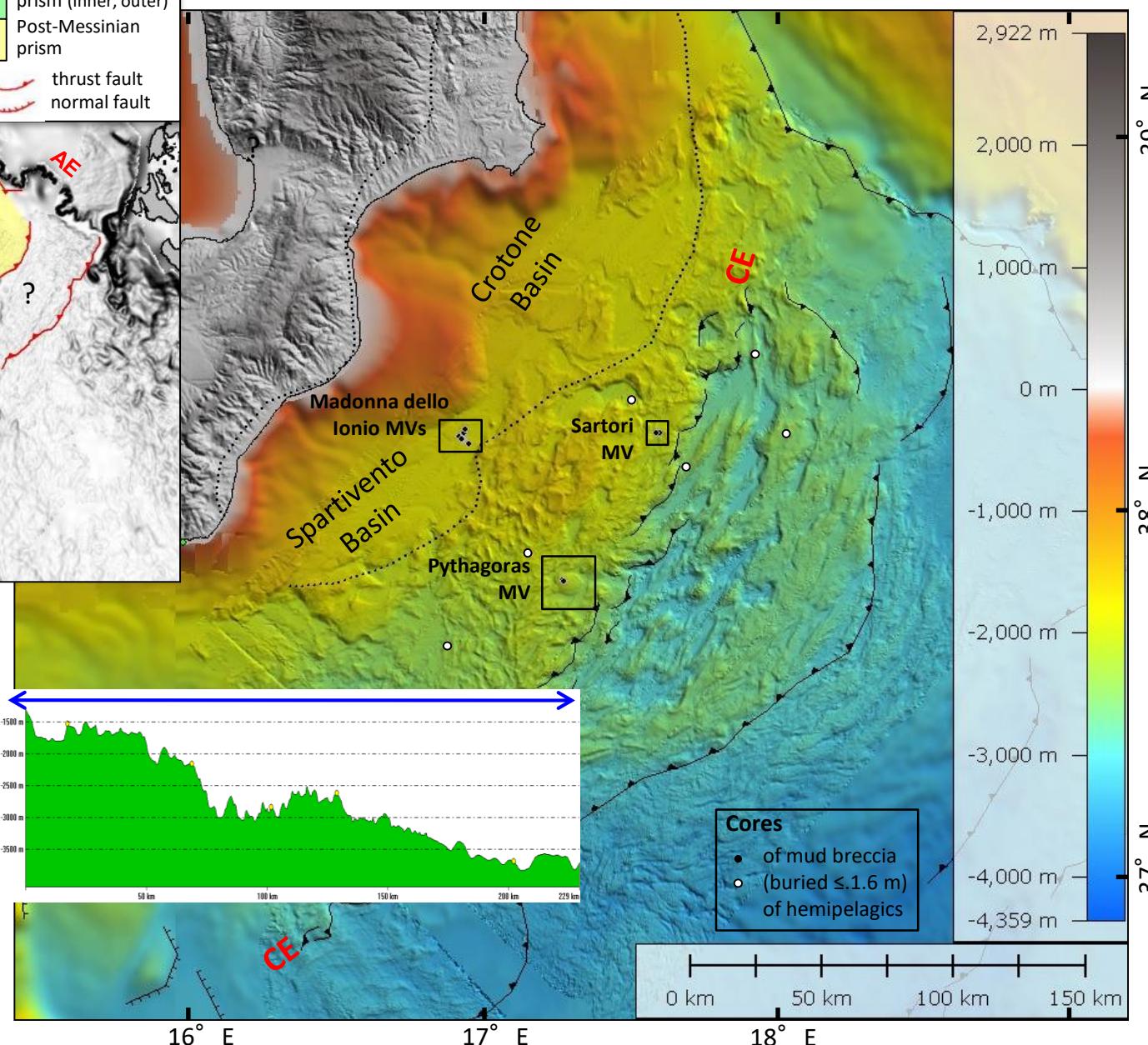
Polonia et al. 2011



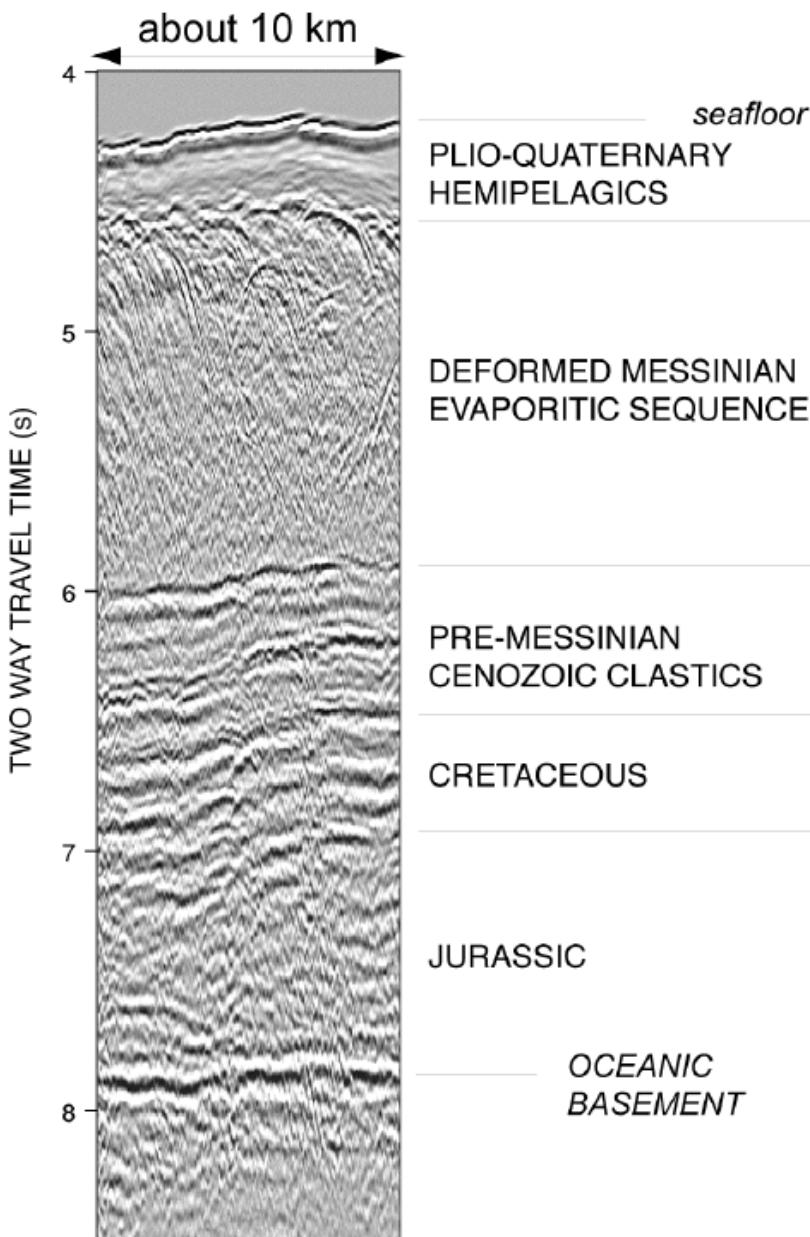
# Bathymetry of Calabrian accretionary prism



**Calabrian Escarpment**  
up to 750 m high, 200 km long, 1<sup>st</sup> order feature of prism, separating inner plateau from steeper outer wedge



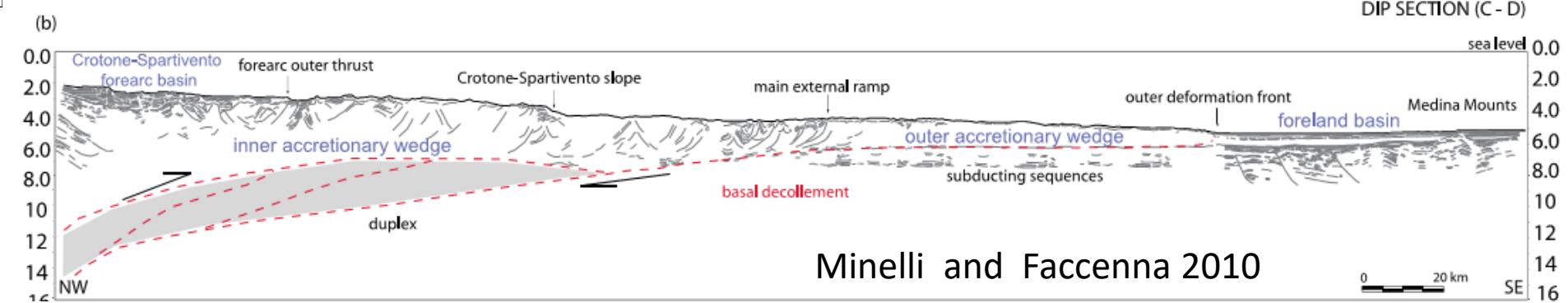
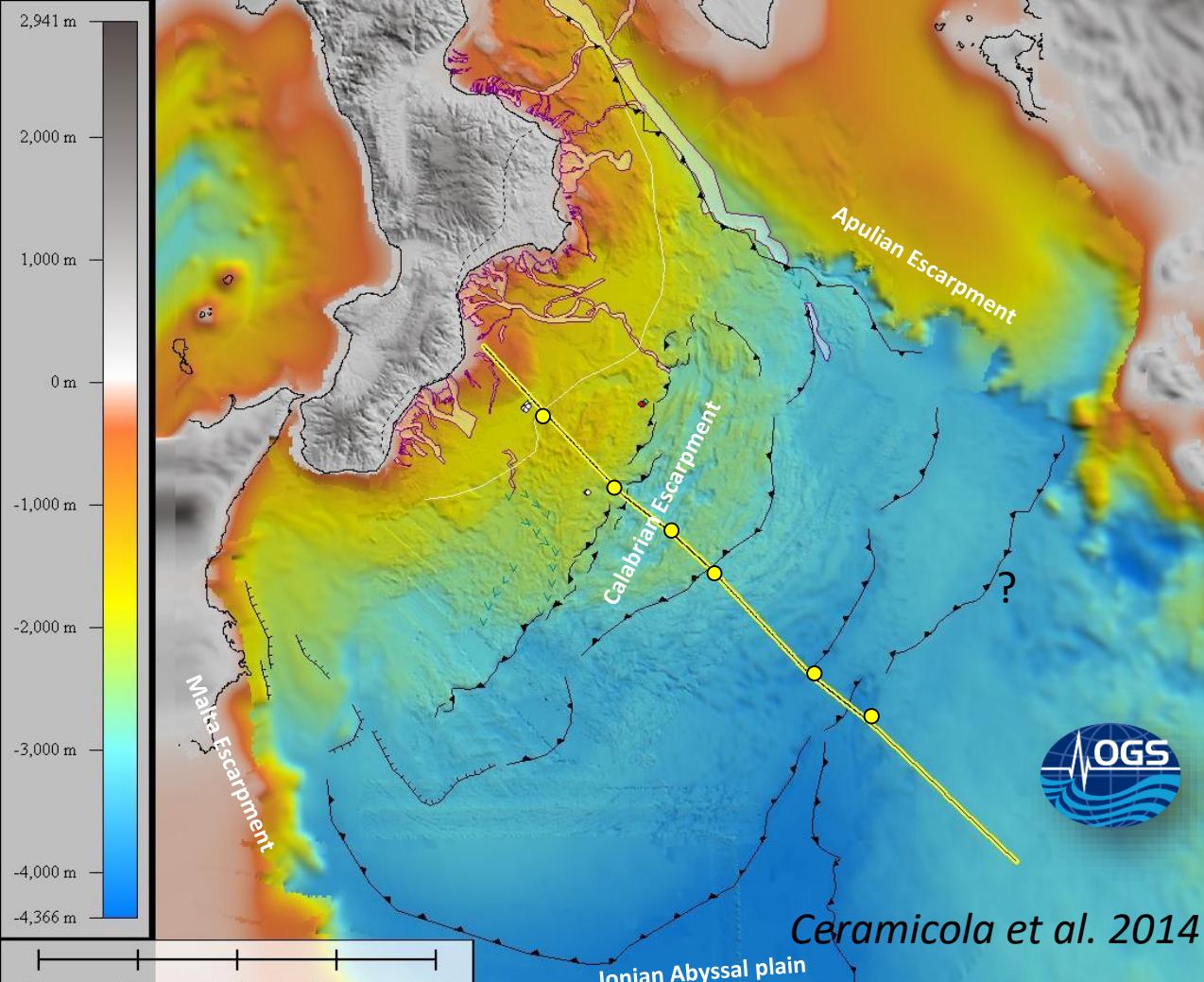
Ceramicola et al. 2014



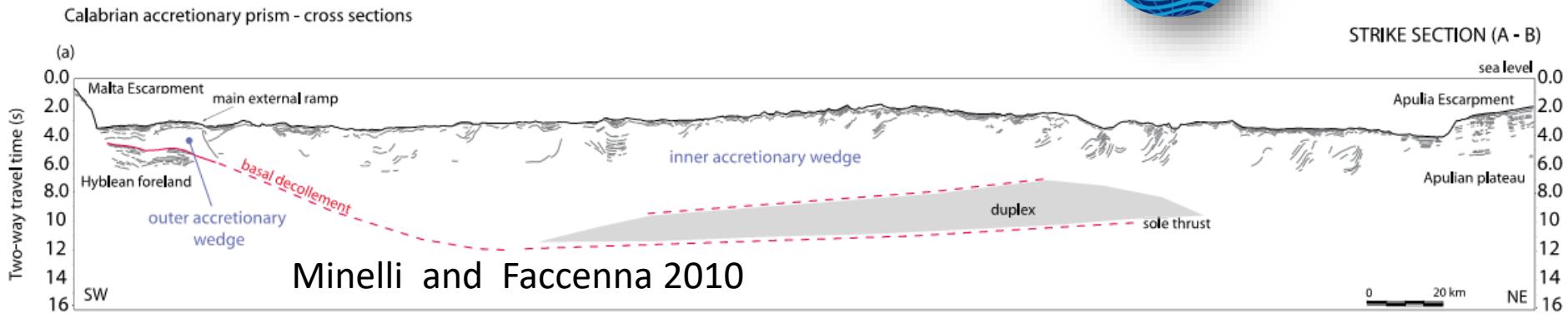
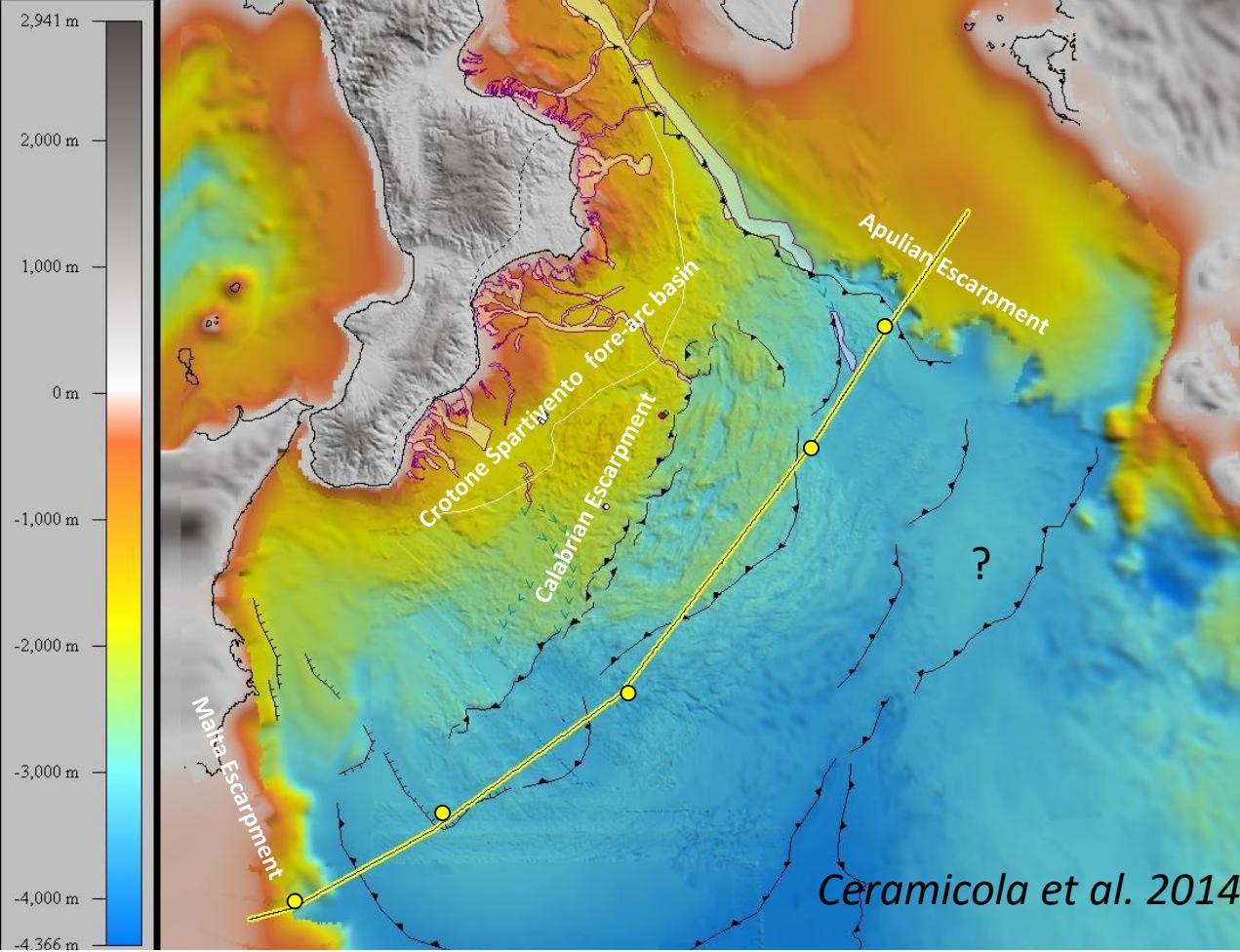
# Seismic stratigraphy of the Ionian basin

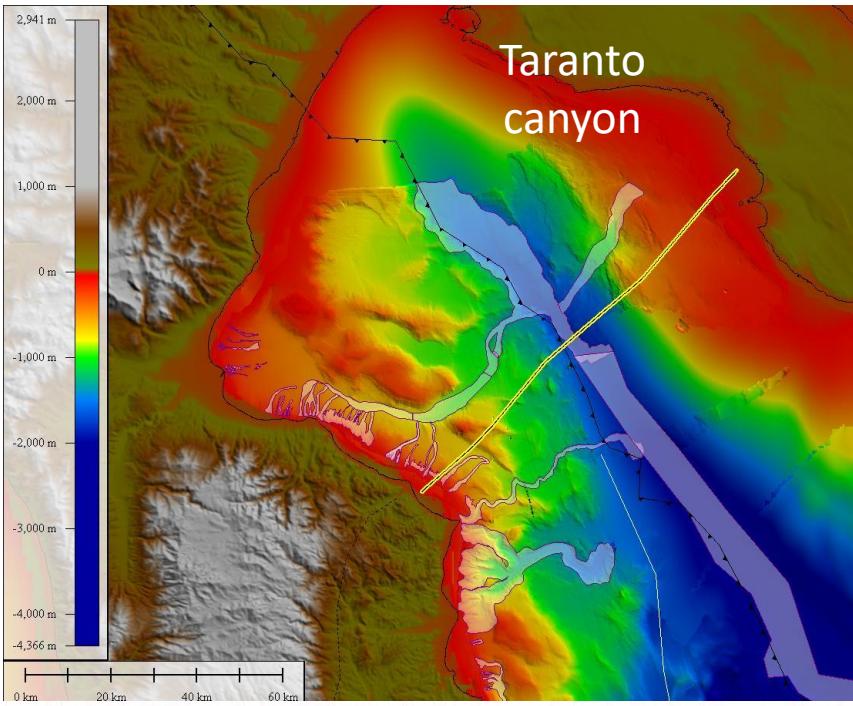
- As the last non-subducted sector of the Neo-Tethys ocean, the Ionian Sea turns out to be the oldest *in situ* ocean fragment of the world.
- It has been saved from subduction since locked within irregular S shaped continental margins of Africa and Eurasia.
- It is a 350 km wide x 600 km long abyssal plain lying at 3–4 km depth, locked between the continental platforms of northern Africa, Malta-Hyblean plateau, and Apulia, and active orogens of Calabria Arc and Hellenides.
- A thick package (5–7 km) of sediments overlying an extremely thin (8–11 km) crystalline crust.

# THE IONIAN SEA

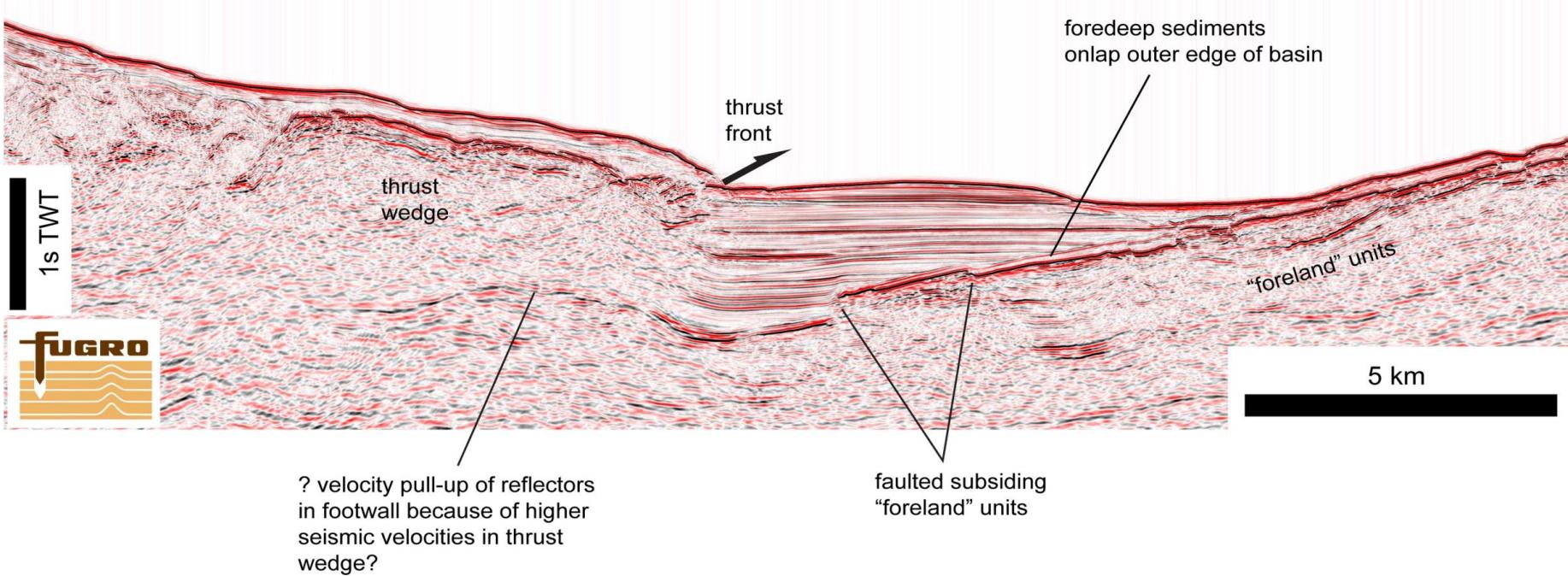


# THE IONIAN SEA



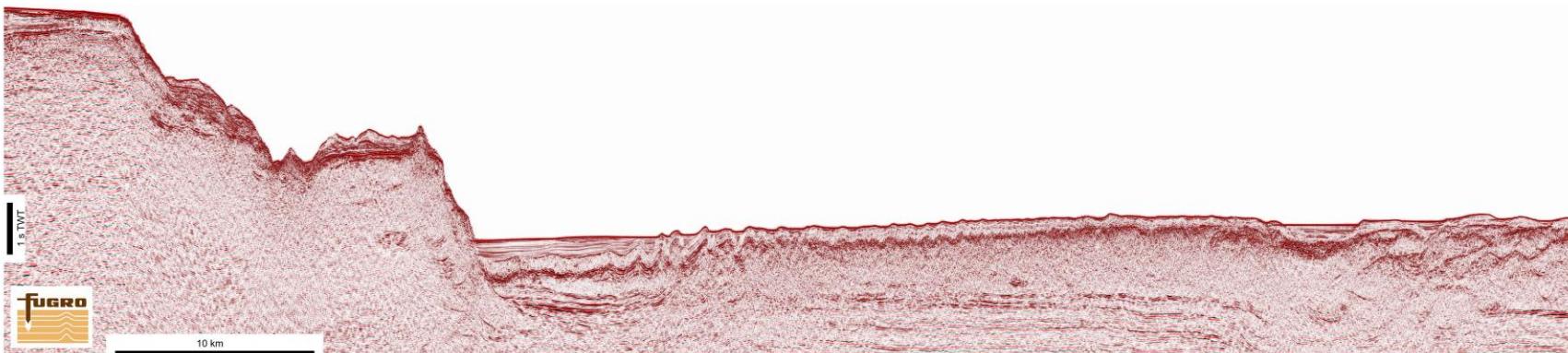
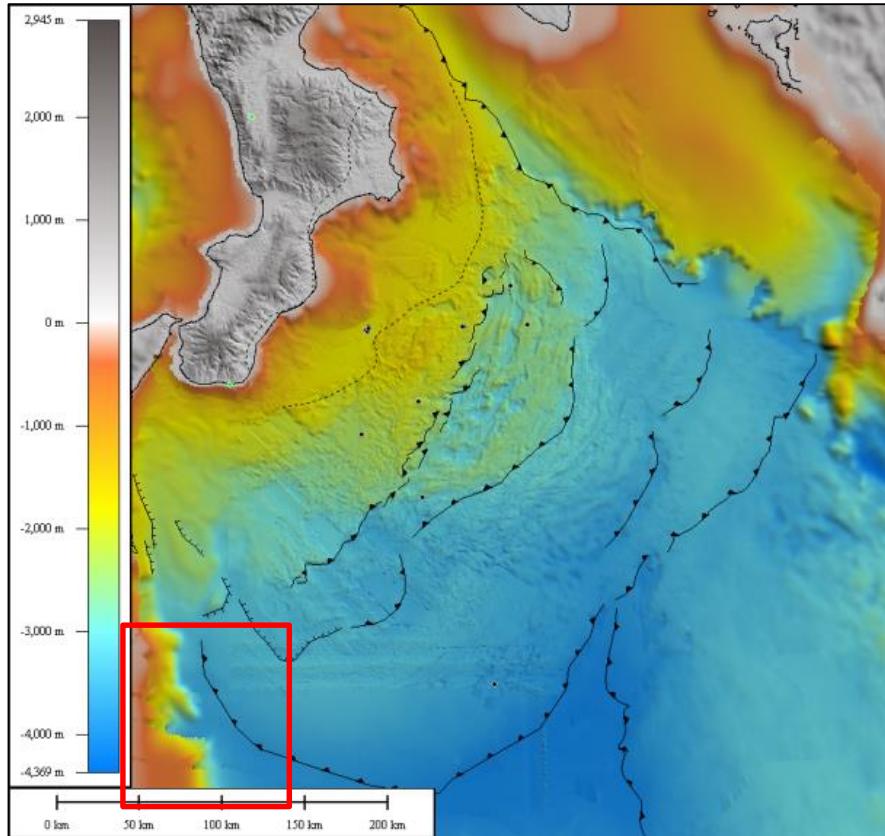


# GULF OF TARANTO

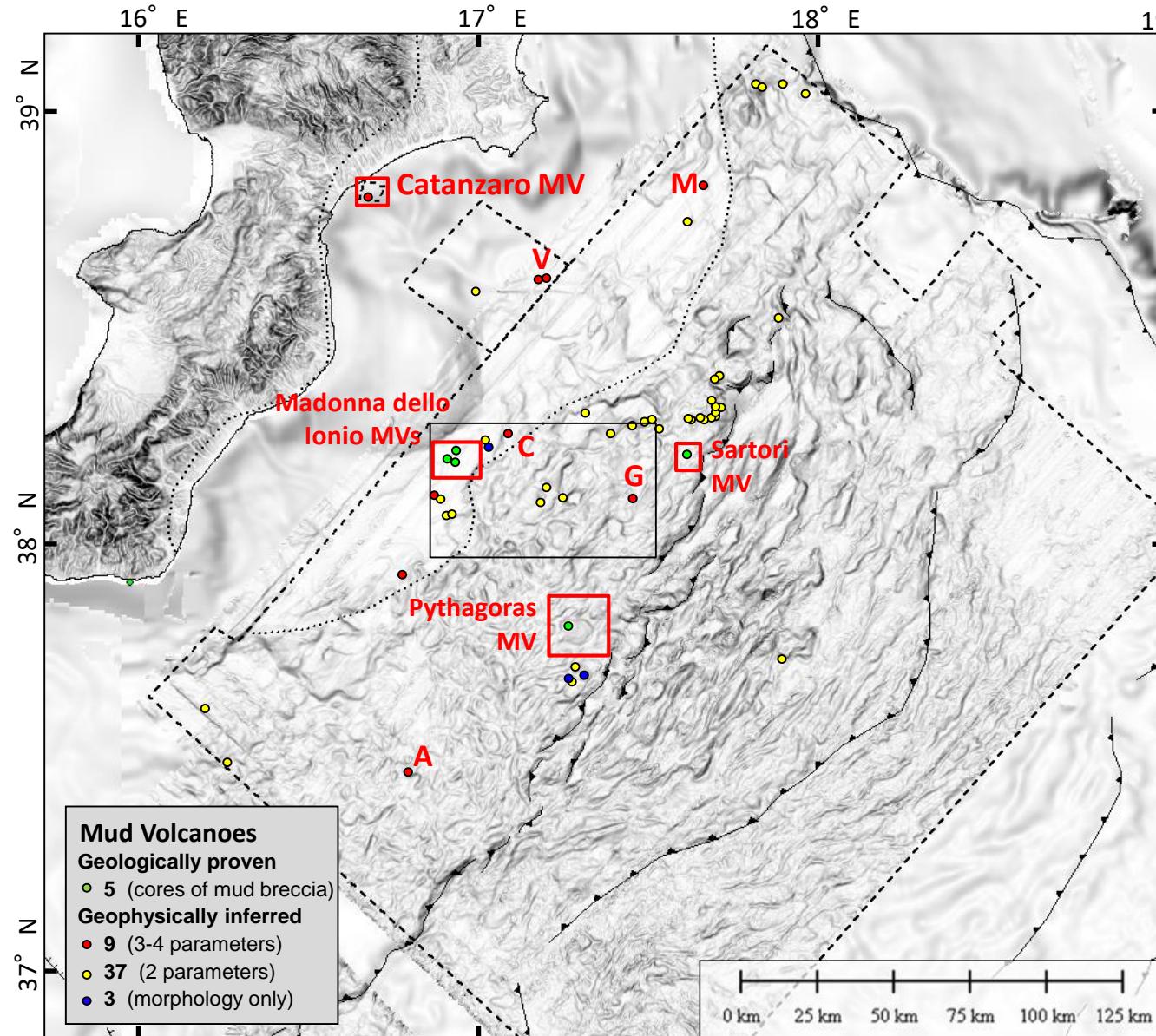




# MALTA ESCARPMENT



# The Calabrian mud volcano province

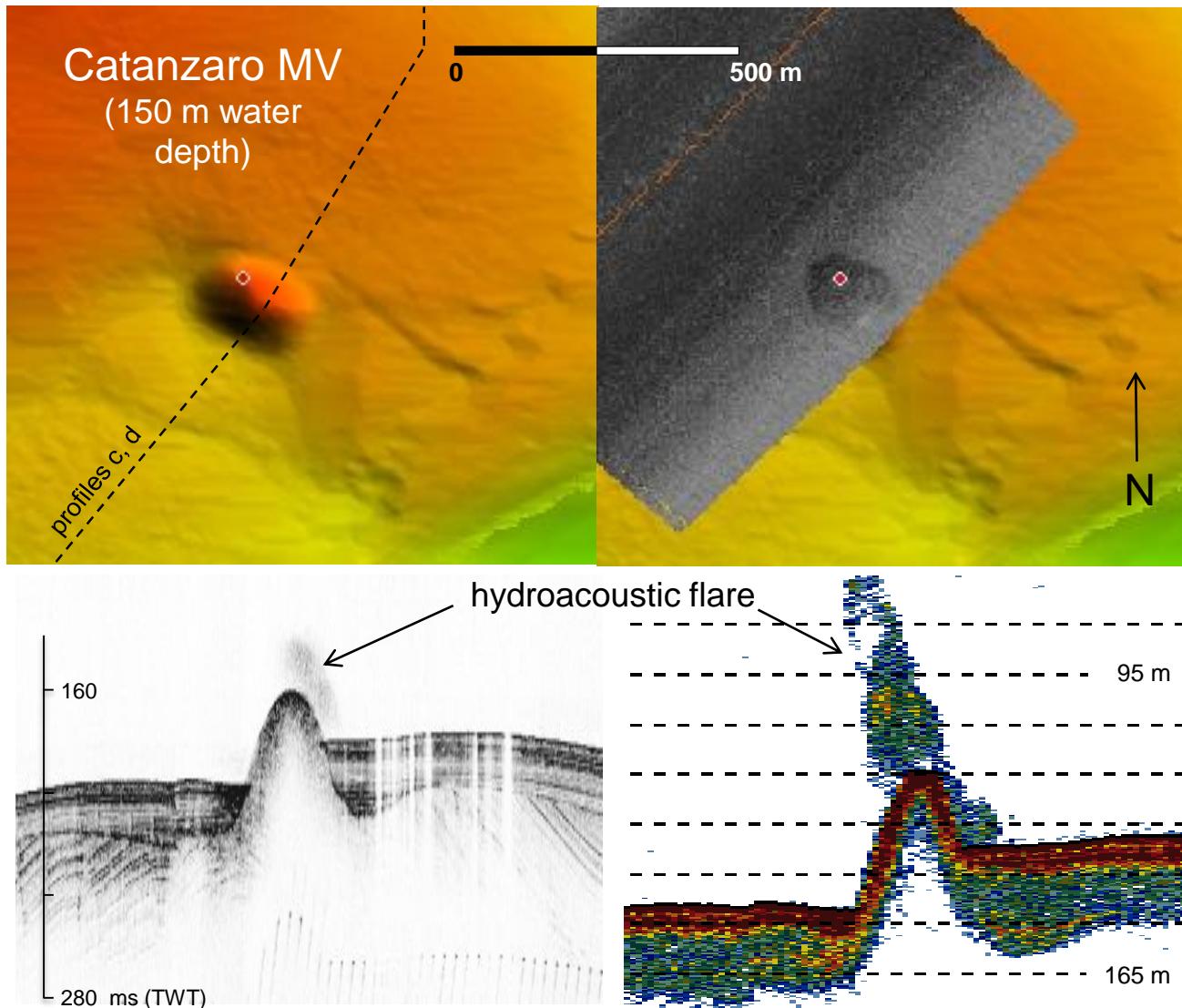


- Seabed mapping based on data acquired by OGS in 2005 and 2009
- At least 54 MVs (conservative criteria), all but one (possible) landward of Calabrian Escarpment
- 3 sites cored, others geophysically inferred

(Ceramicola et al. 2014b)



# Multi-parameter geophysical evidence of mud volcanoes

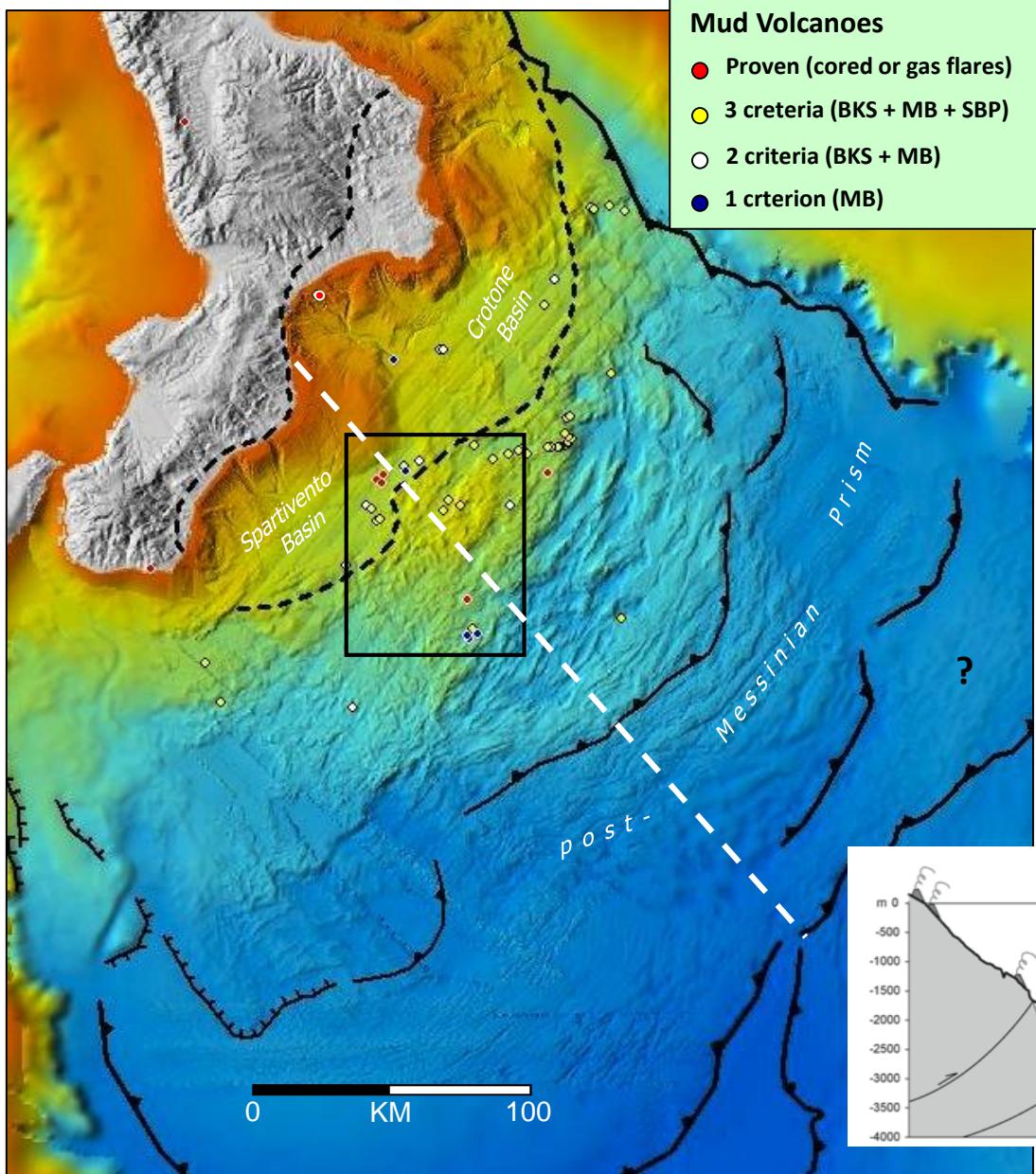


- 5 km off Calabrian coast, only MV in <1000 m water depth (100 kHz backscatter)
- Mud breccia extrusion above last glacial maximum unconformity (19-**23 ka BP**)
- Hydroacoustic evidence of gas venting to water column

Ceramicola et al. 2014b



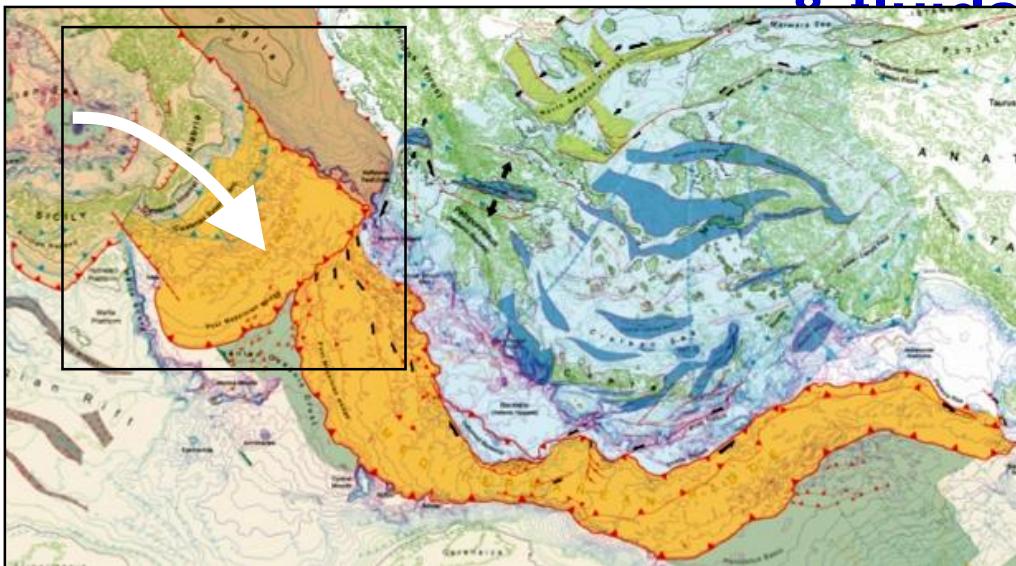
# Calabrian mud volcano province – in summary



- At least 54 MVs (+ one on land) distributed across fore-arc basins (thin salt) and inner prism (no salt) – lower density than Med Ridge, fluid sources at depth
- Local influence of faults, none in centres of fore-arc basins (rising fluids blocked by salt or mud?)
- Seabed extrusion of mud breccias over last glacial to post-glacial cycle (50/53), ongoing gas seepage (3 sites)
- All but one possible MV landward of Calabrian Escarpment

# Central and Eastern Mediterranean accretionary prisms

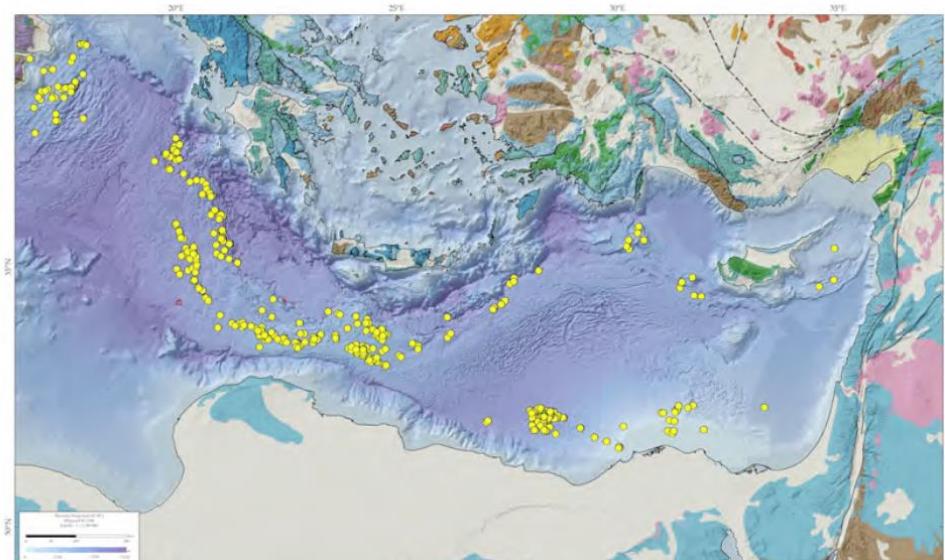
## • Shale



Chamot-Rooke et al. 2005

## Calabrian ‘arc’ :

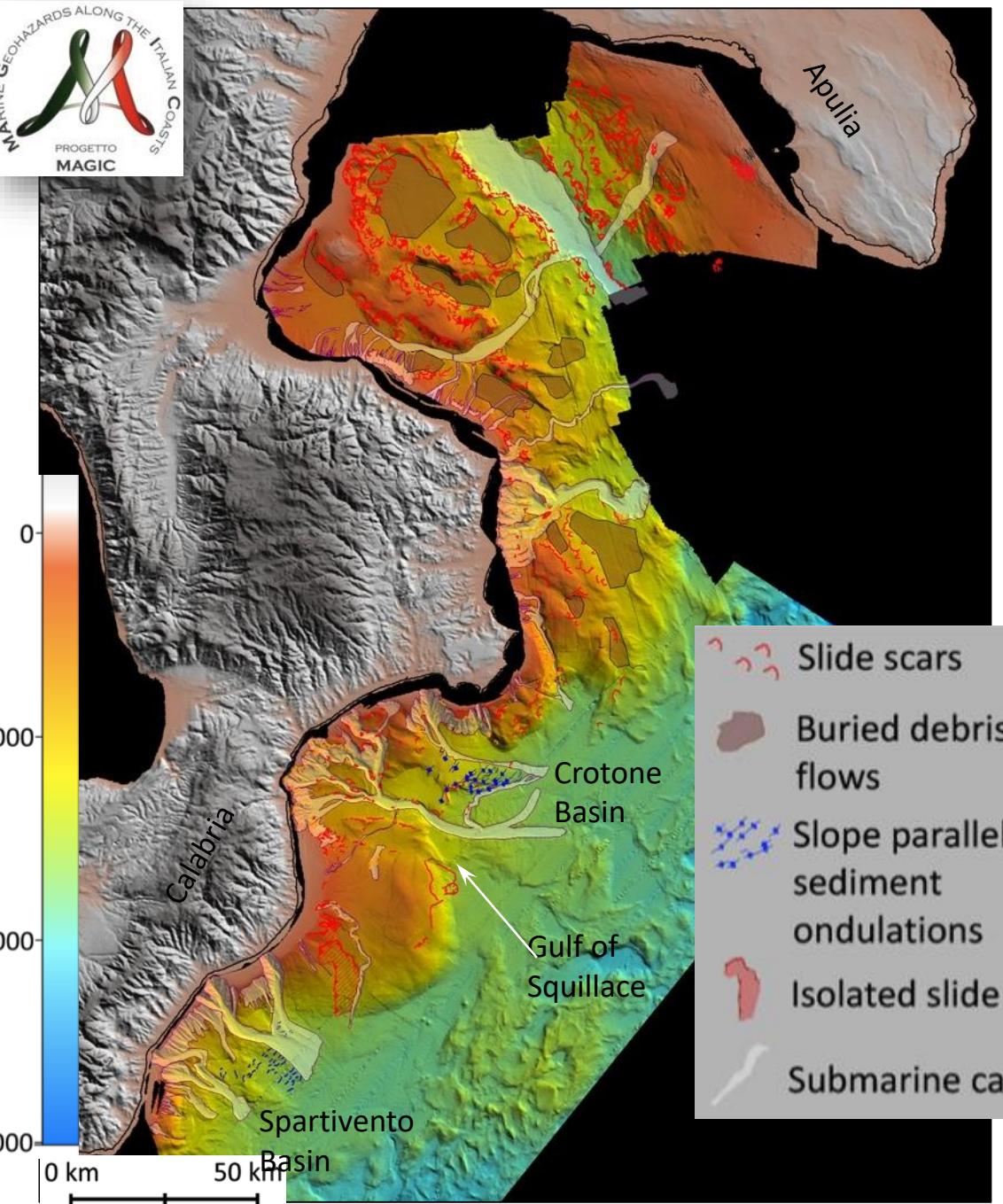
- Recent entry – rapid slab roll-back since the mid-Miocene c. 9 Ma → meets Med Ridge
- No seabed studies for >25 years since Rossi & Sartori (1981)
- Presence of long-lived mud volcanoes proven by OGS in 2005



Mascle et al. 2014

## Mediterranean Ridge etc :

- MVs 1st identified from mud breccias over 30 years ago (Cita et al. 1981)
- Mud breccia extrusion for >1 Ma (ODP 164, 1996)
- Hundreds – highest abundance on Earth? (Kopf 2002)
- Tectonically-controlled crestal belt

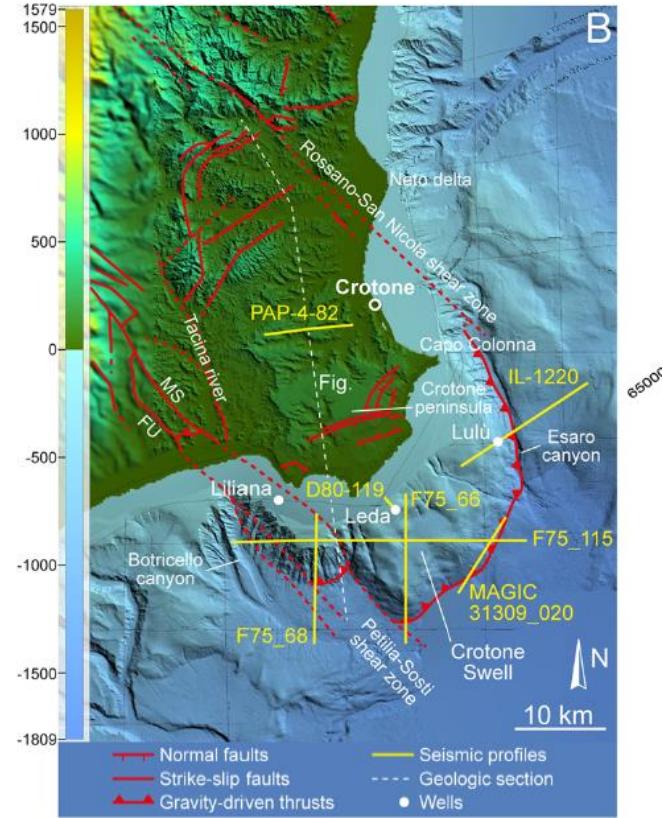
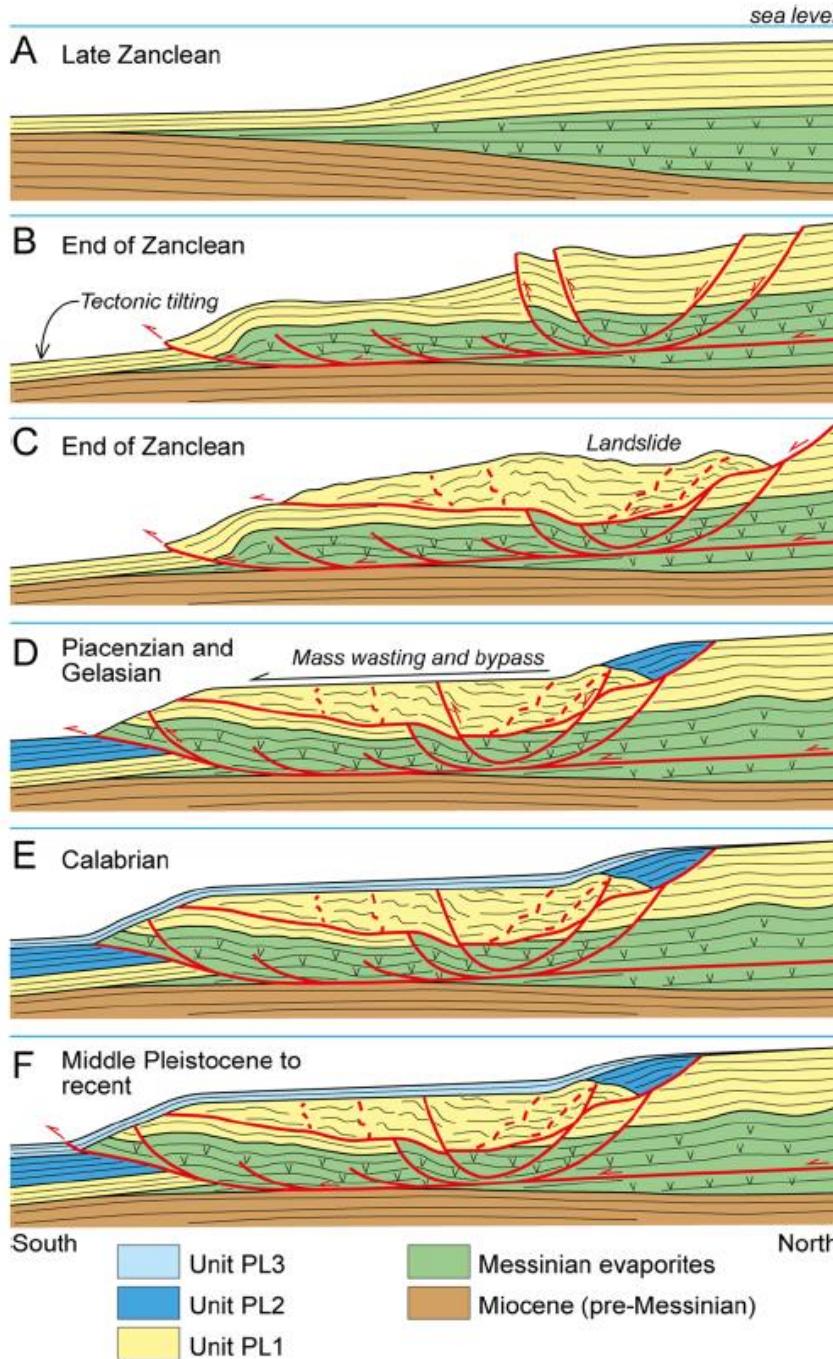


# Mass movements along the Calabrian margin

- 1 Mass Transport Complexes (MTCs) in intra-slope basins (northern Calabrian margin)
- 2 Isolated Slide Scars (ISSs) on open slopes (all margins)
- 3 Headwall & Sidewall Scars in Submarine Canyons (HSC) (Calabrian margin)
- 4 Slope-Parallel Sediment Undulations (SPSU) (southern Calabrian margin)

*Ceramicola et al. 2014a*

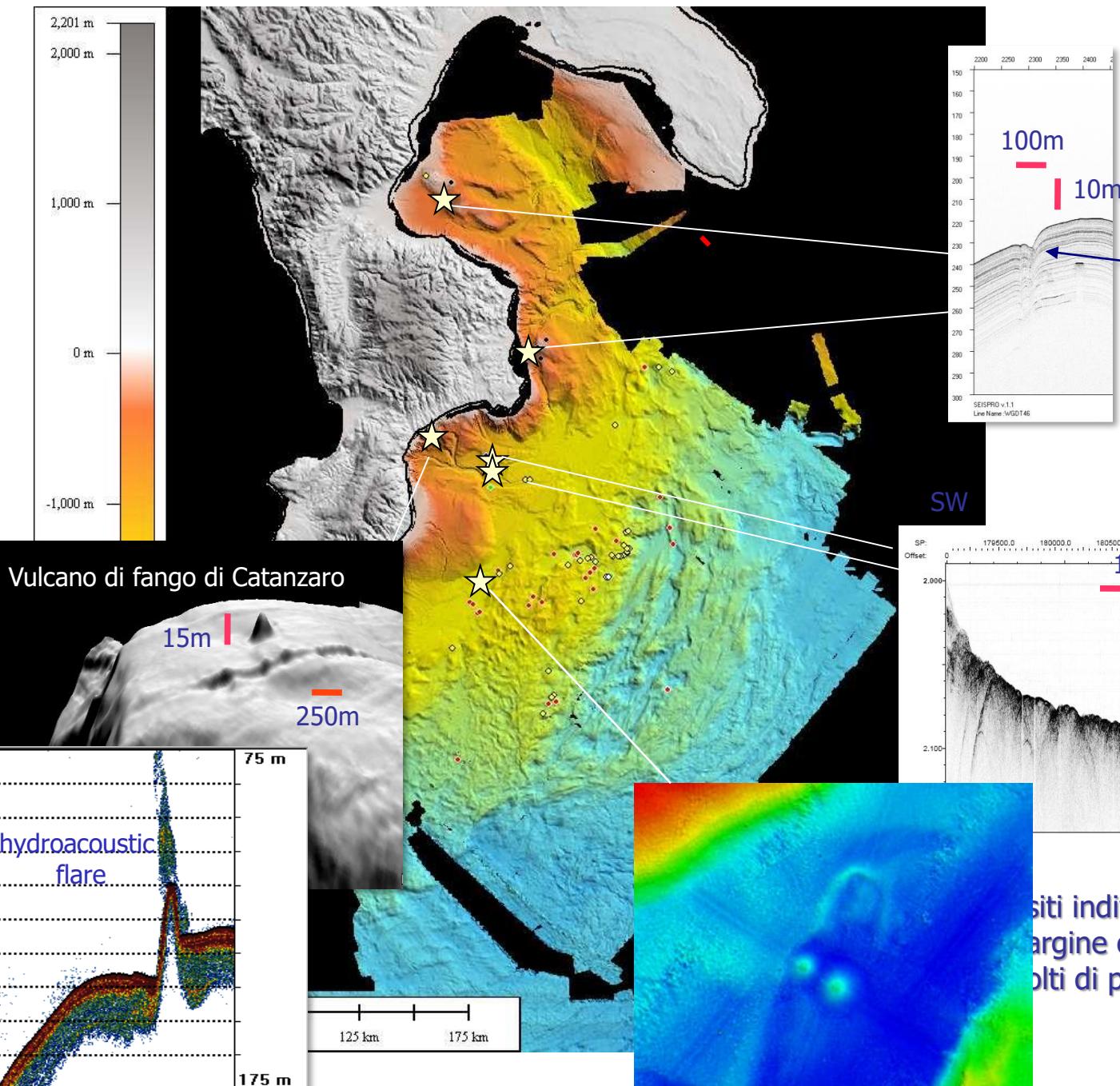
# The Crotone Megaslide



Zecchin et al. 2018



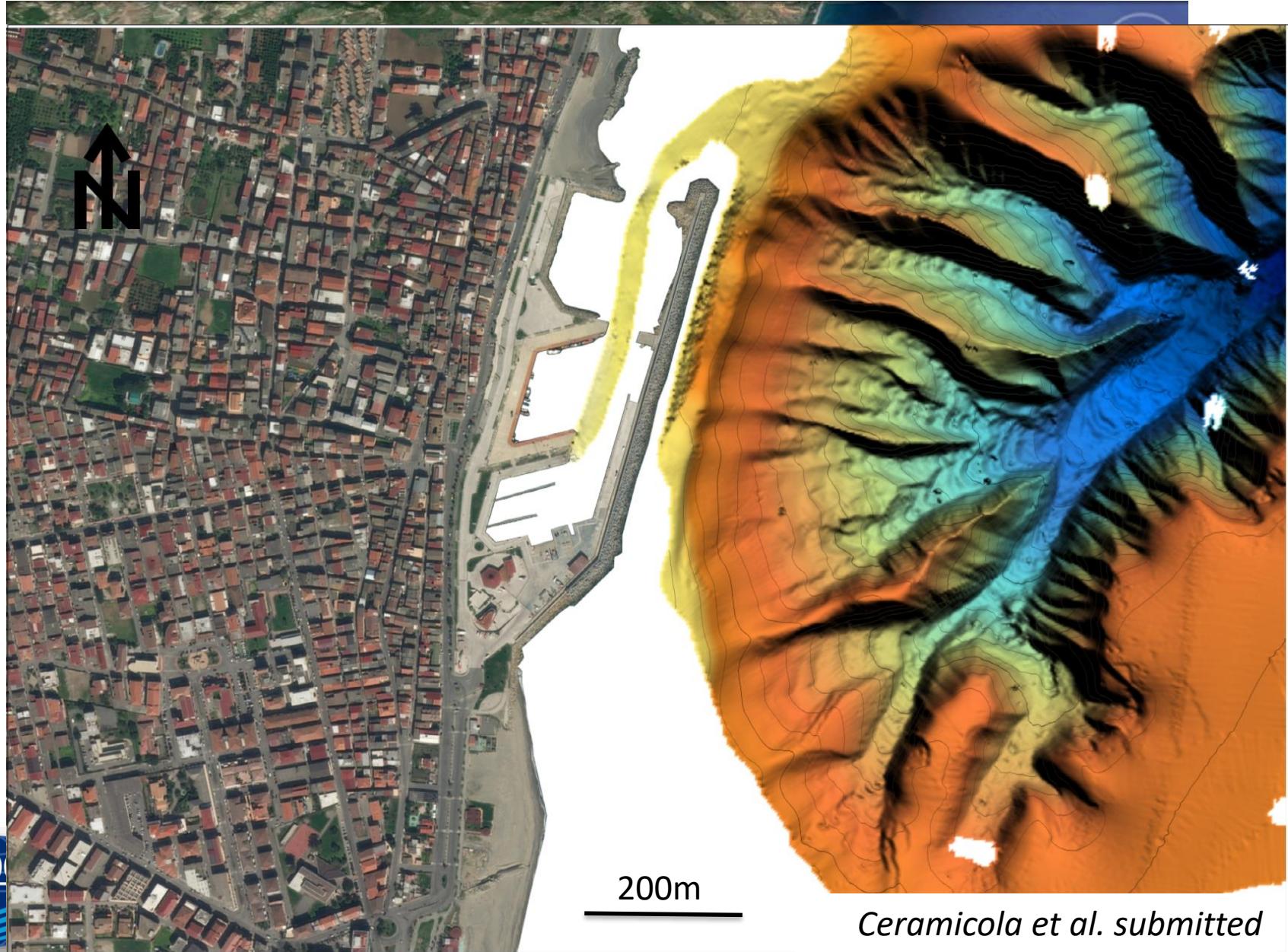
# Fluid seepage



Pockmarks



# Cirò marina submarine canyon and coastal hazard



*Ceramicola et al. submitted*

# Cirò marina submarine canyon and coastal hazard



Porto di Cirò Marina  
Ore 11:00 del 01 Dicembre 2013  
Foto by Sergio Marino

# Adriatic-Ionian Bimodal Oscillating System (BiOs) (Gacic et al., 2010)

The Ionian Sea is the deepest regional sea of the Mediterranean and plays an important role in the intermediate and deep thermohaline cell of the Eastern Mediterranean (EMed) conveyor belt (Gacic et al., 2010)

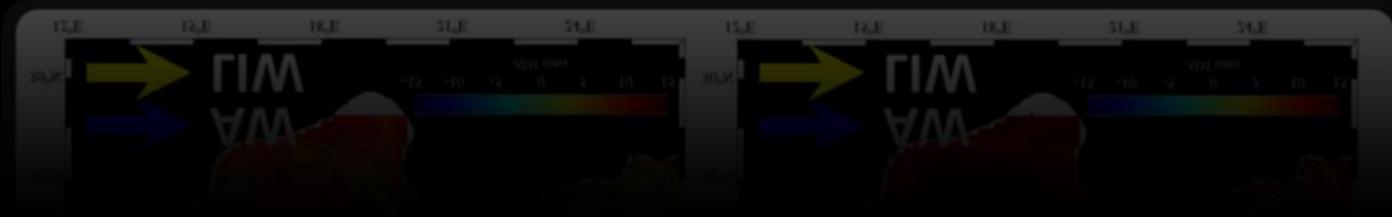
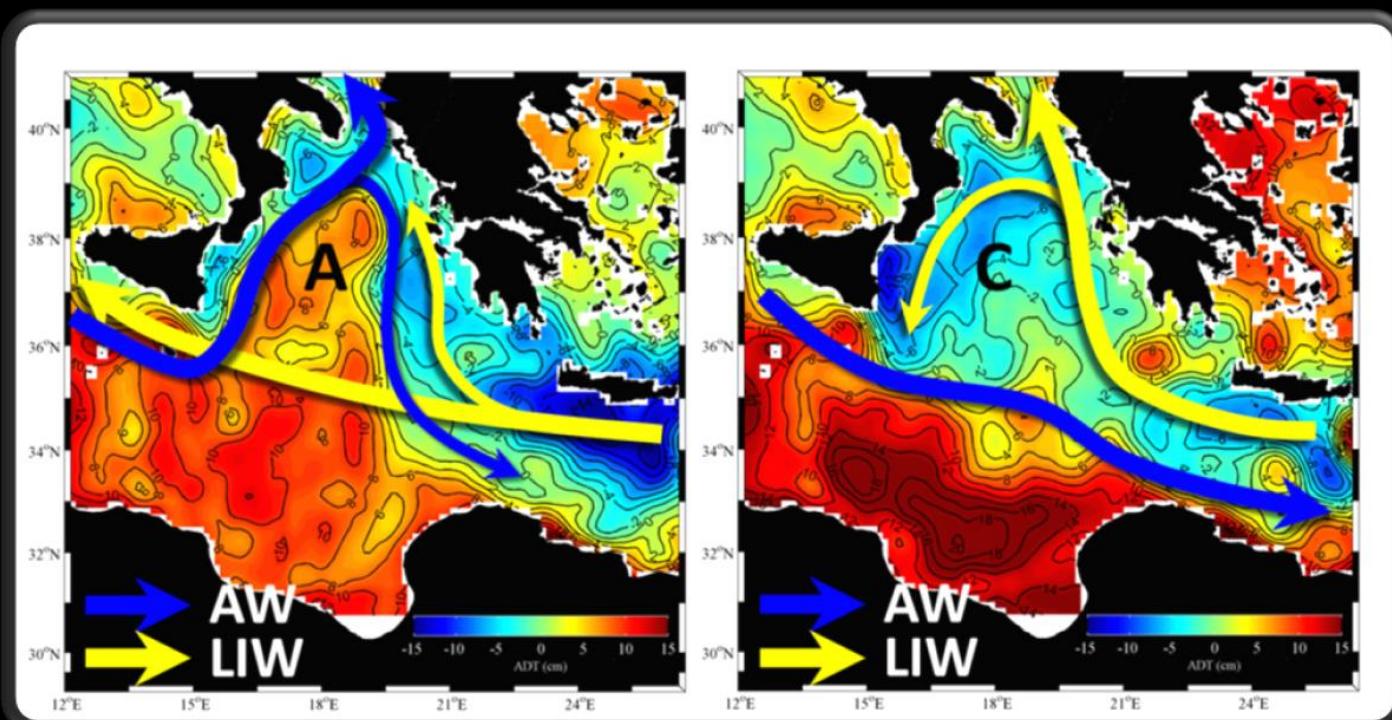
The vertical structure of a water column is formed by three layers : dense and oxygenated waters, mainly of Adriatic origin (in the bottom layer), salty and warm waters coming from the Levantine and Aegean basins (in the intermediate layer) , relatively fresh water of Atlantic origin (AW) propagating toward the Levantine basin over surface

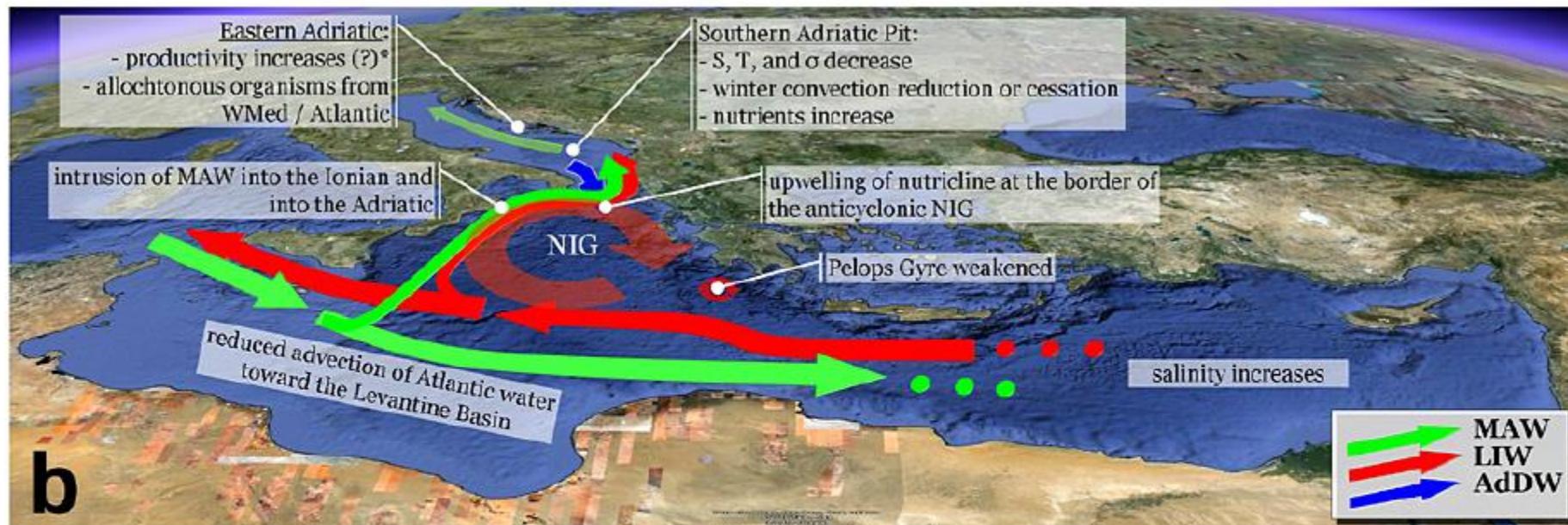
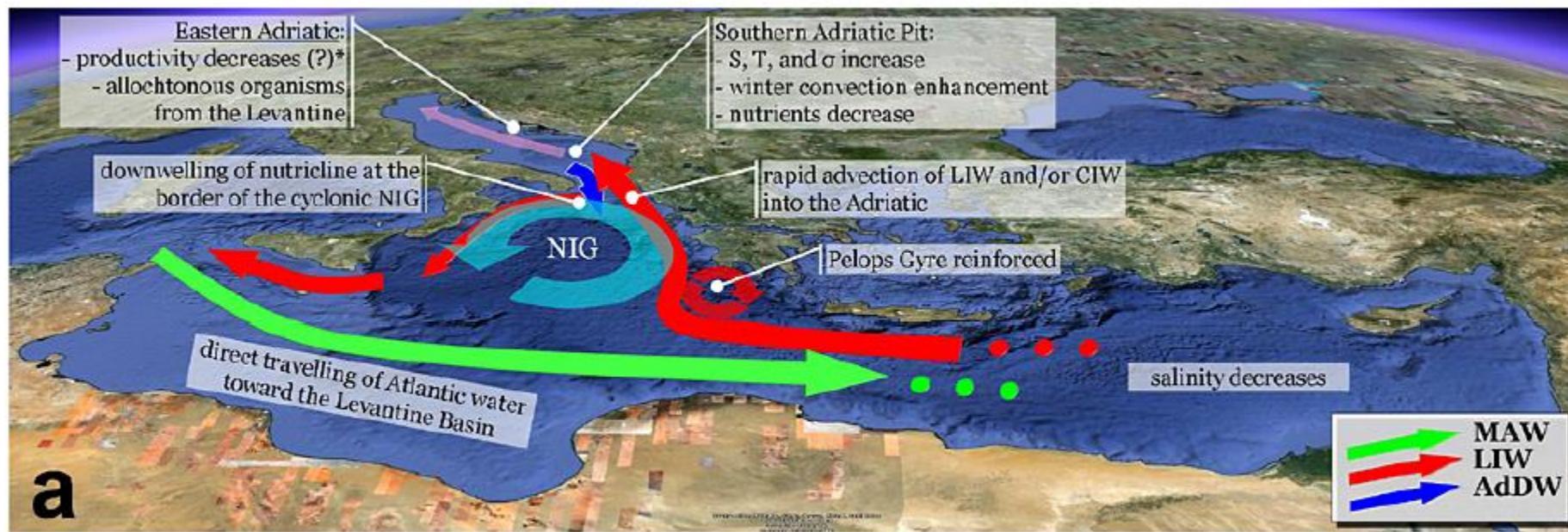
A reversal of circulation in the Ionian Gyre from anticyclonic to cyclonic has been observed in the middle of 1997.

Adriatic Ionian system behaves as a bimodal oscillating system when anticyclonic circulation are present in Ionian AW are deflected in the Adriatic leading to production of ADW of lowering density which spread in the Ionian producing a deepening of isopycnal surface and stretching of water column ! weaking of anticyclonic circulation produce the reverse of circulation to cyclonic when cyclonic circulation are present in Ionian salty LIW enters in the Adriatic leading the production of ADW of increasing density which produce a shallowing of isopycnal surface → weaking/inversion of cyclonic circulation



The Ionian Sea circulation ***reverses*** on decadal scale.  
Accordingly, Atlantic Water path is deviated, generating  
an ***alternate «dilution»*** of the  
North Ionian and Adriatic Sea (A), or Levantine (C)





**Fig. 5.** Summary of the main characteristics of the Adriatic-Ionian BiOS and its impact on the area. (a) cyclonic NIG; (b) anticyclonic NIG

\* For more detailed explanations, see main text (Sect. 4.2). For acronyms, see main text.