AMBIENTI DELTIZI



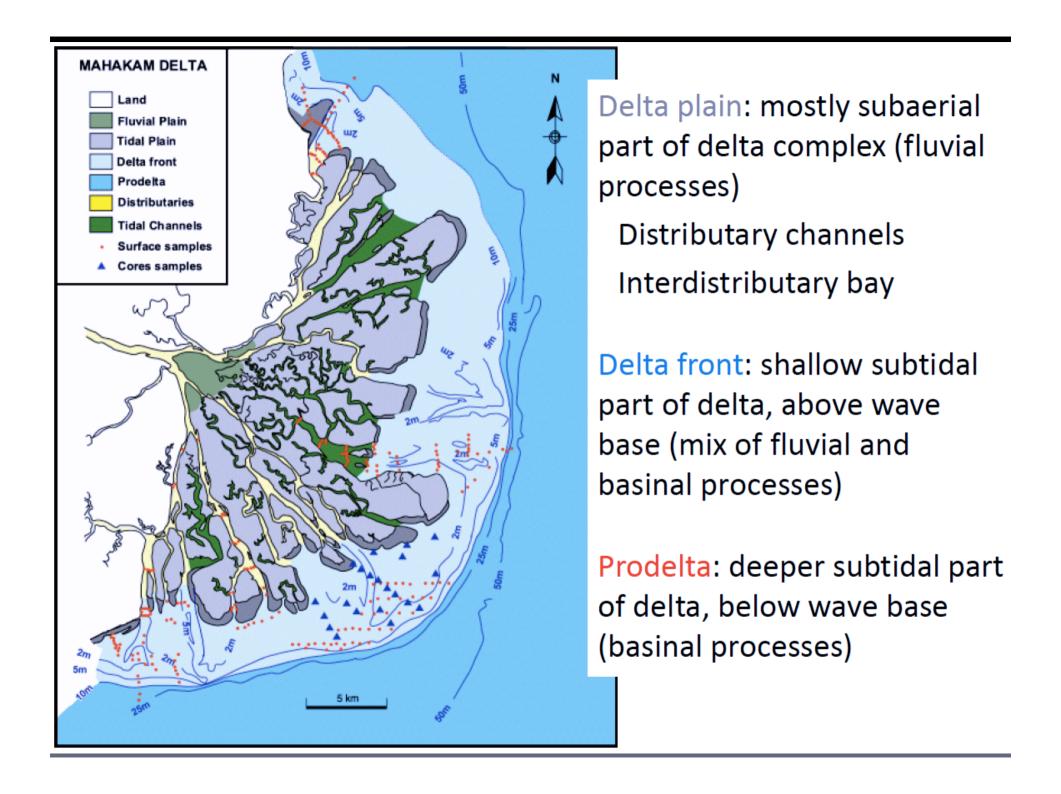
Delta del Nilo

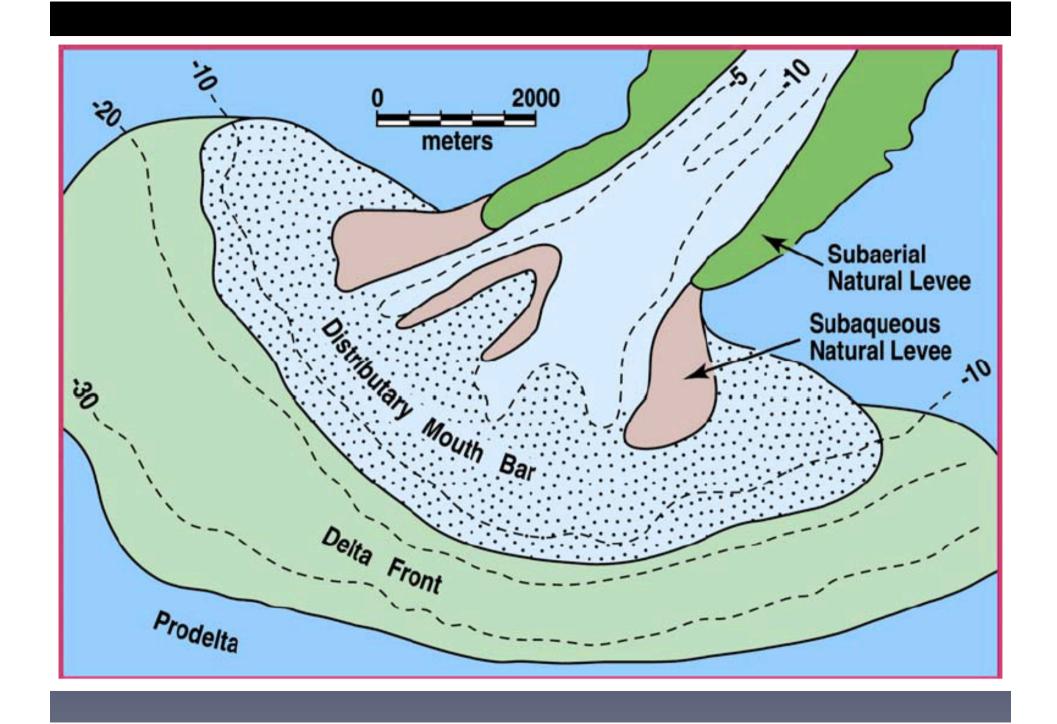
Delta del Mississippi

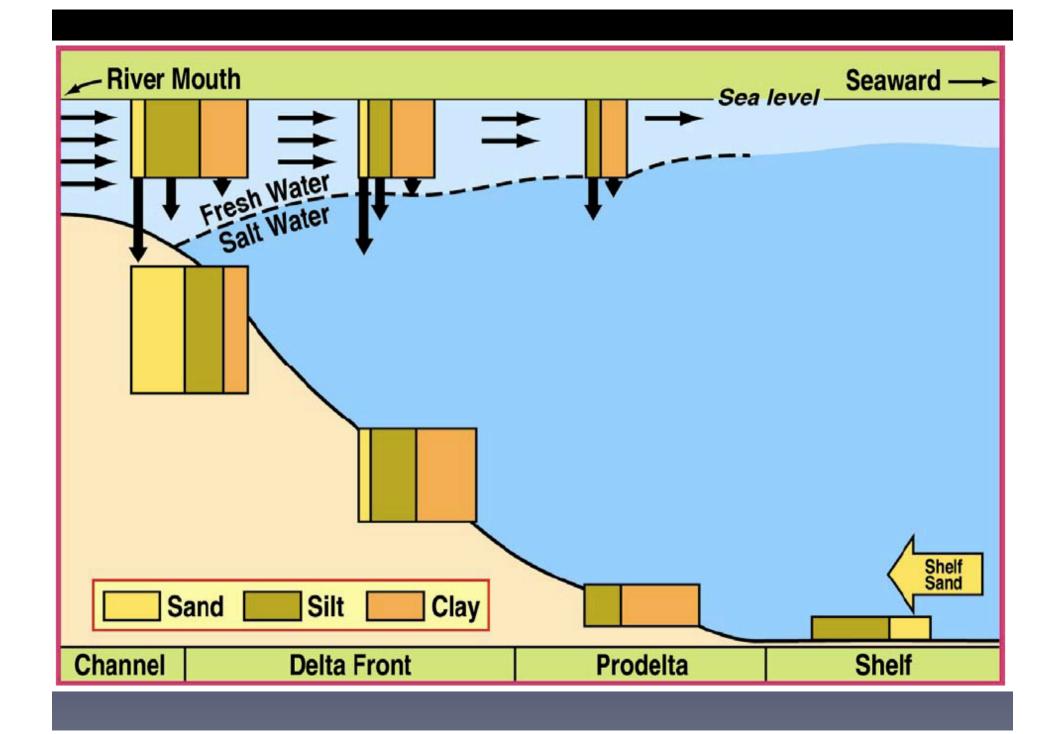


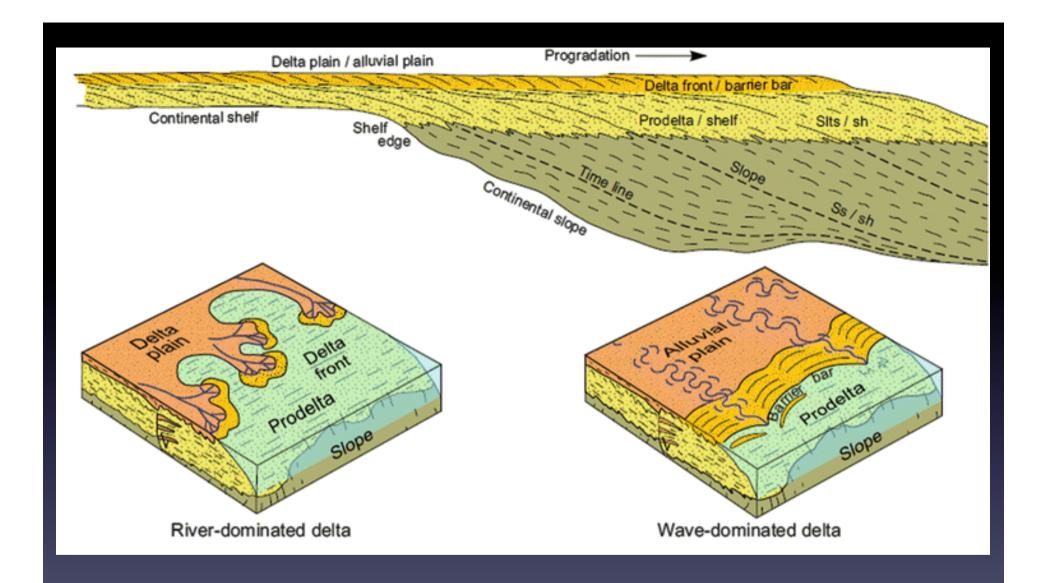


Discrete shoreline protuberances, partly subaerial, built by rivers into a body of permanent water

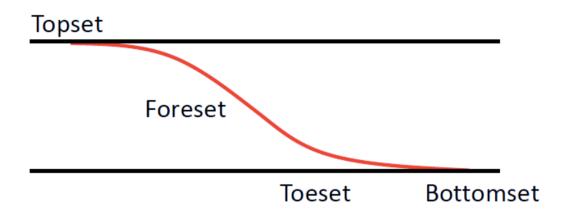


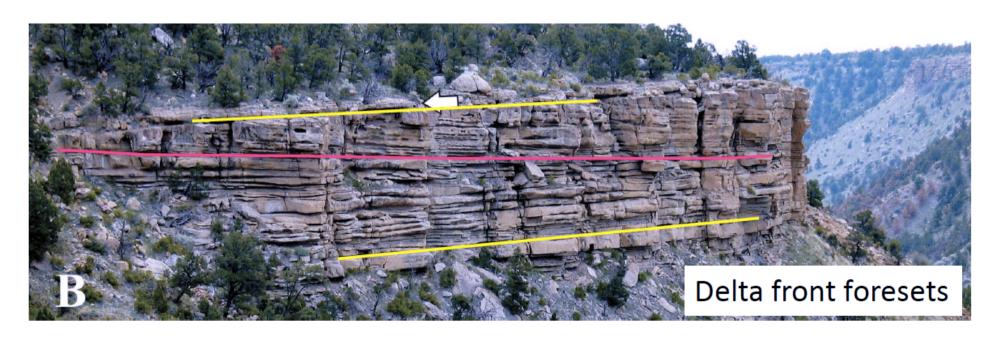


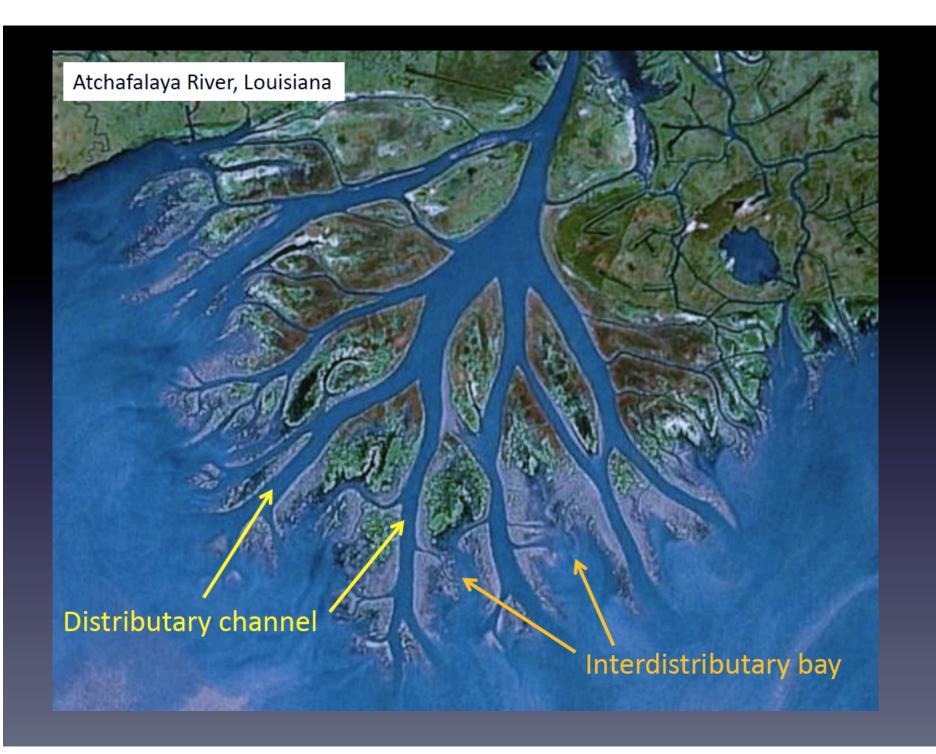


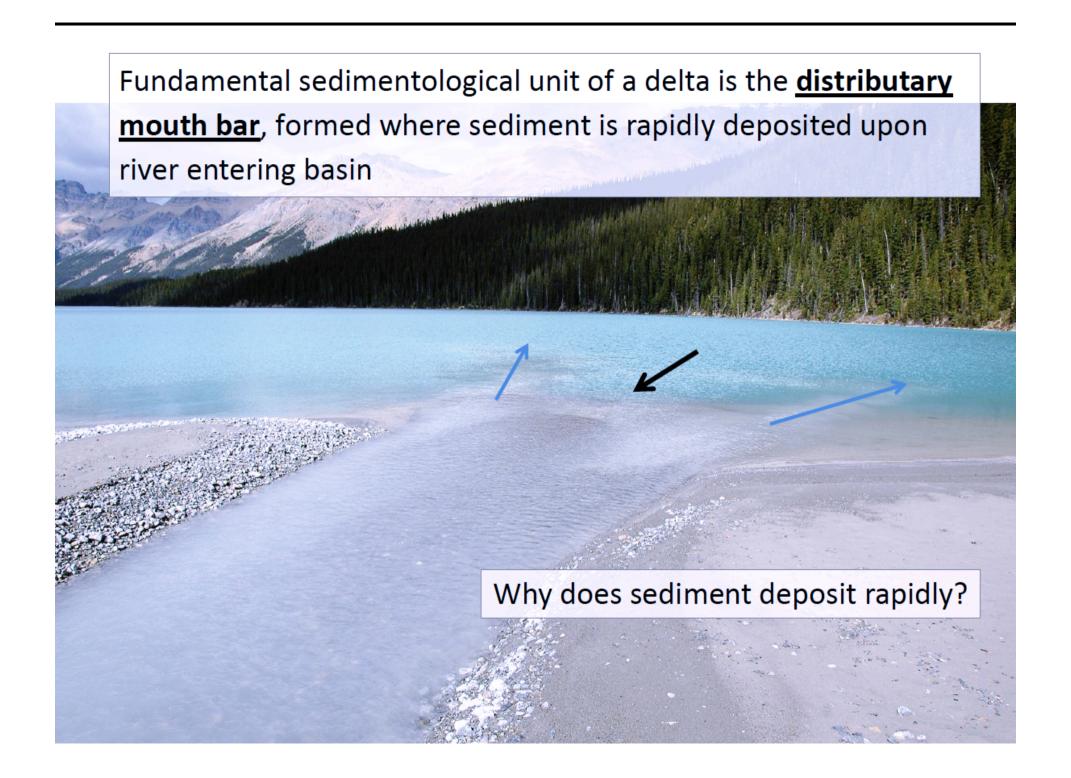


Progradation of relatively steep delta front (1-10°) produces a bed geometry called <u>clinoforms</u>



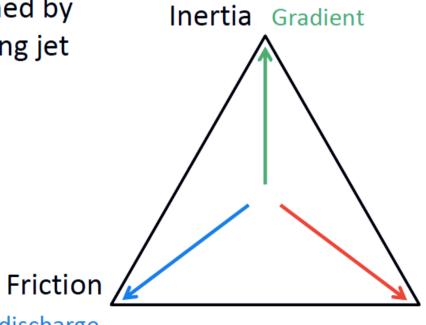








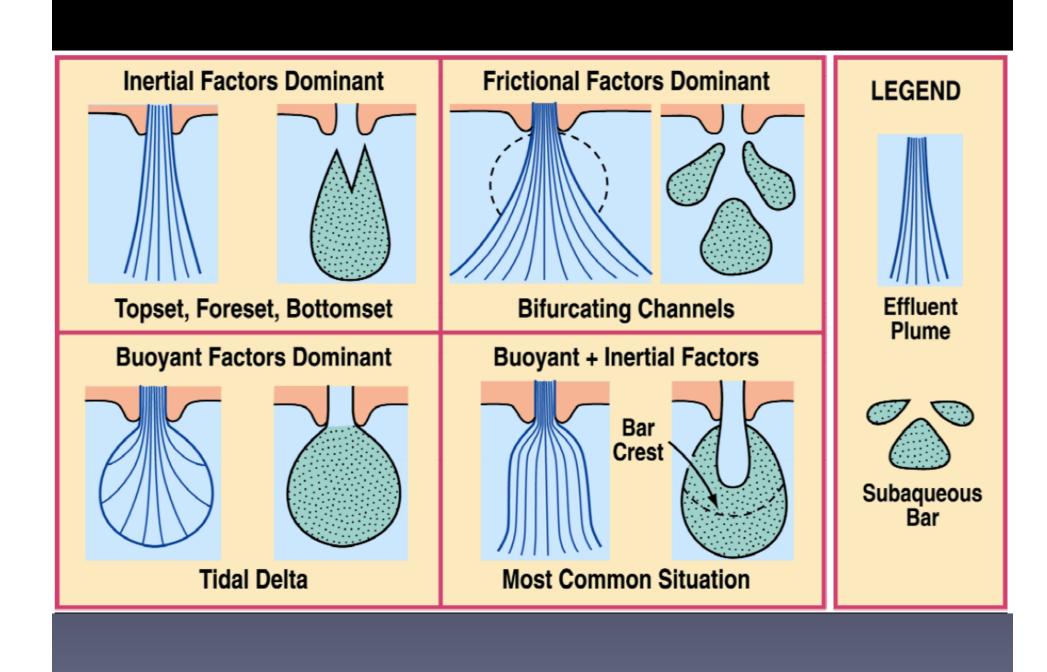
Mouth bar governed by physics of spreading jet of river water



Buoyancy

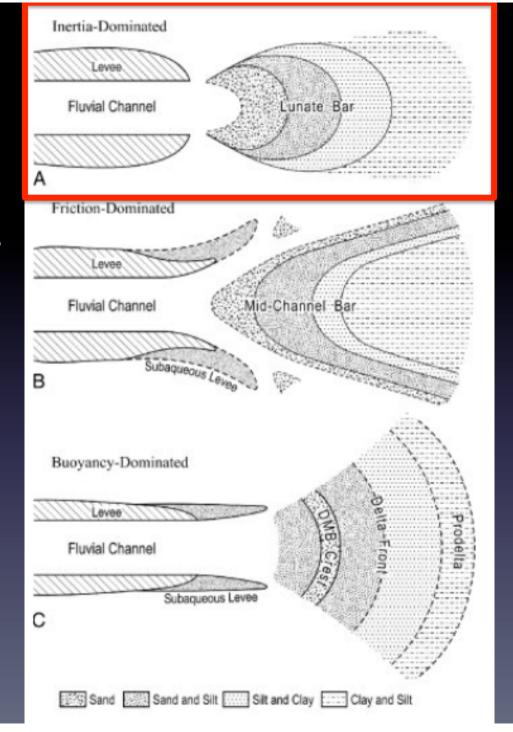
Fluvial discharge

Density contrast



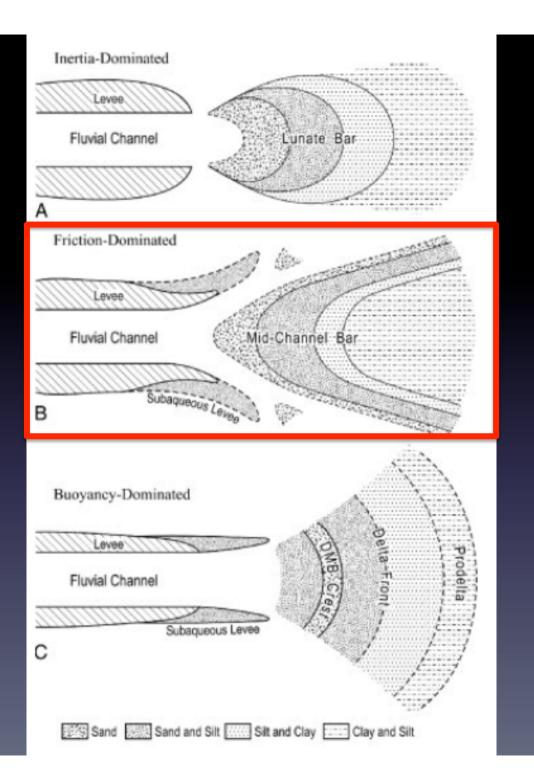
Inertia-dominated deltas

are considered to be an uncommon dominance for deltas. They are associated with high flow velocities and large amounts of turbulence. Sediments are deposited close to the main flow of the channel as it enters the basin. In other words, the deposition of sediments in inertiadominated deltas does not have a large lateral component.



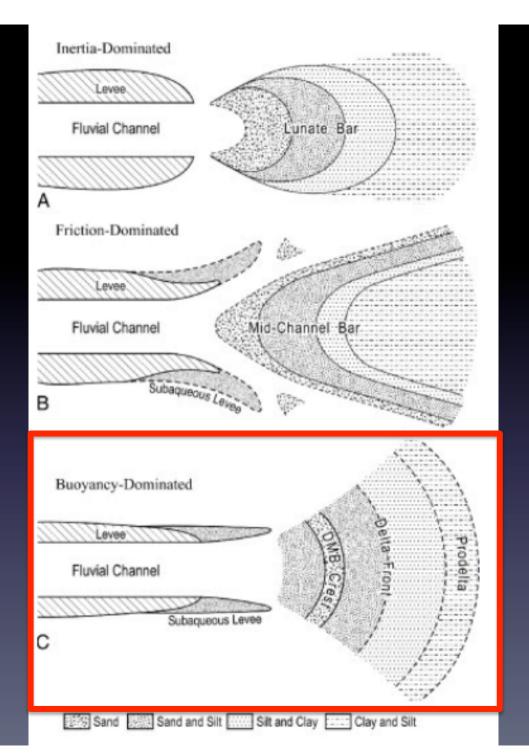
Friction-dominated deltas

are associated with high bed friction and shear stress. These stresses cause the flow to slow down rapidly and deposit sediment with a wider lateral extent than the inertia dominated deltas. The river/basin setting for this type of situation is more common. It consists of a shallow water level where the river flow and basin meet. This shallow area is a major factor in producing a friction-dominated delta. Features associated with this type of delta include subaqueous levees, middle ground bars (fining seaward), and bifurcated channels.



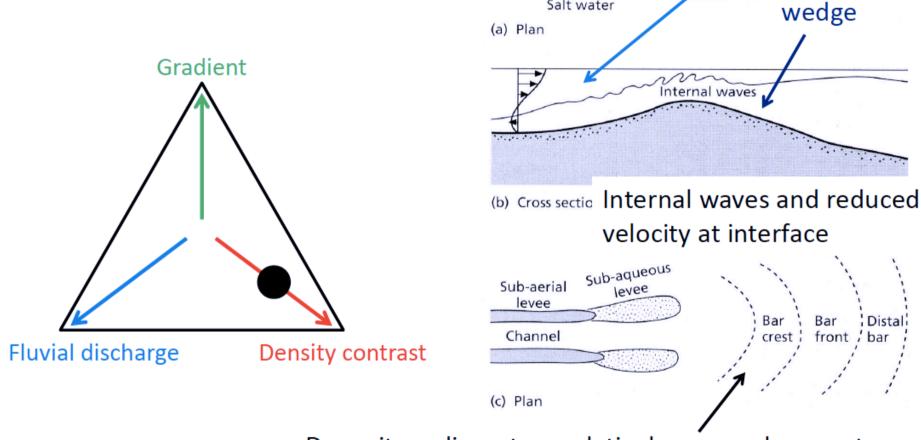
Buoyancy-dominated deltas

occur where the river/basin setting consists of a deeper water level when compared to the friction-dominated delta. This situation leads to the formation of subaqueous levees (parallel banks extending out from the channel). Other deposits common with this type of delta are distributary mouth bars (grading seaward), bar finger sands, distal bars, and prodelta clays.



Buoyancy-dominated deltas (density contrast)

River water is usually less dense than ocean water (hypopycnal)



Deposits sediment on relatively narrow bar crest

Fresh water

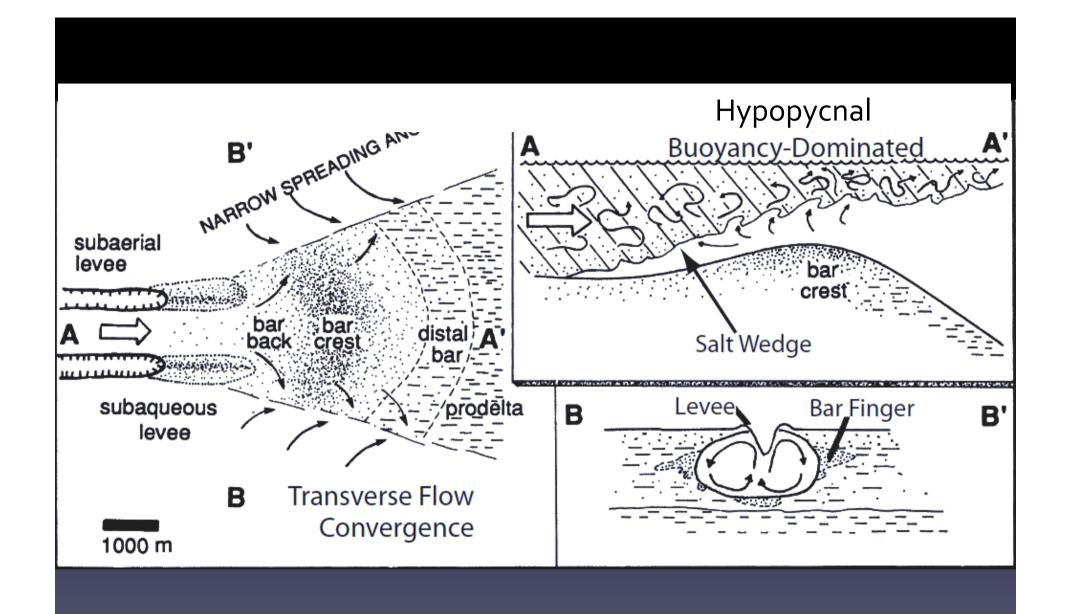
Salt water

Freshwater

lens above

saltwater

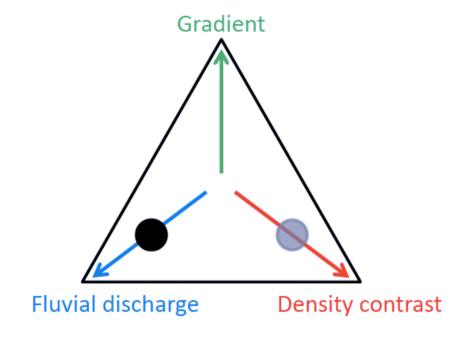
What is the proximal-distal grain size trend?



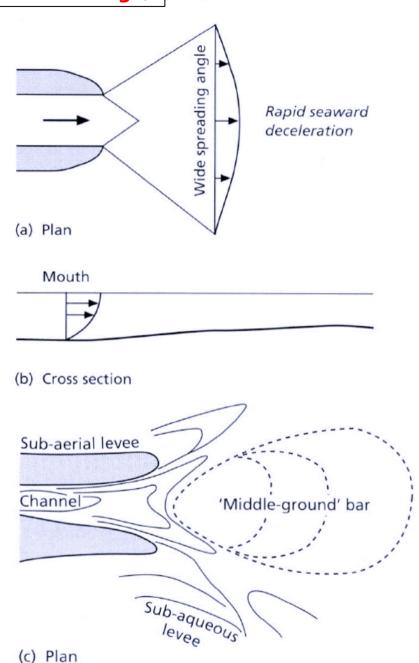
Examples of mouth-bar processes in river-dominated deltas (from Reading and Collinson, 1996, after Orton and Reading, 1993) incorporating ideas of Bates, 1953, Wright (1977) and others.

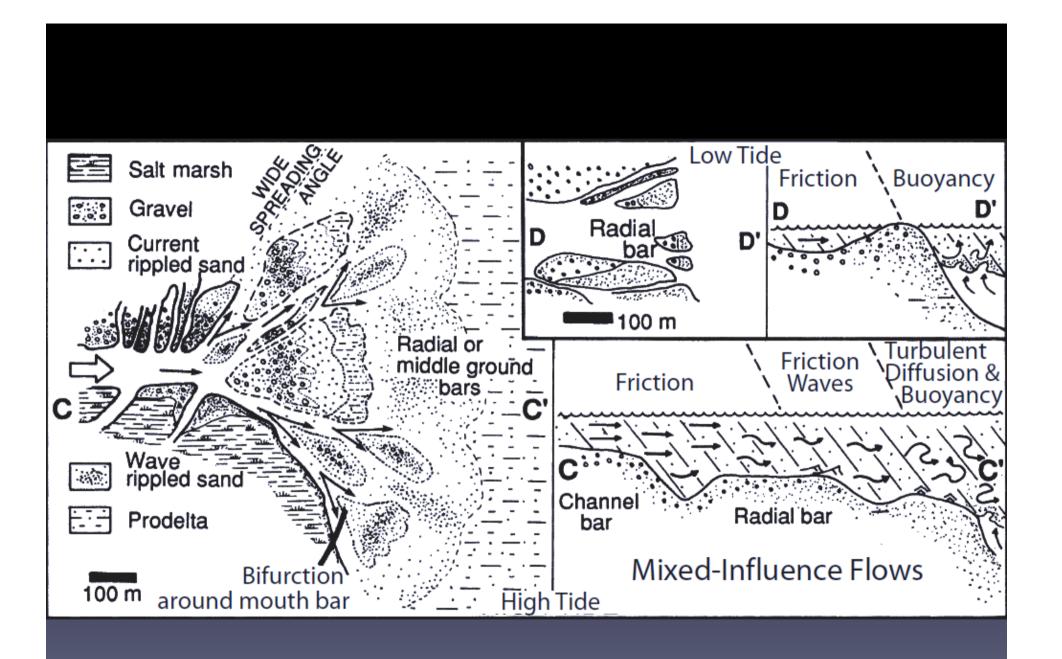
Friction-dominated deltas (fluvial discharge)

Flood stage pushes salt wedge out More sediment = lower ρ contrast Bottom friction dominates



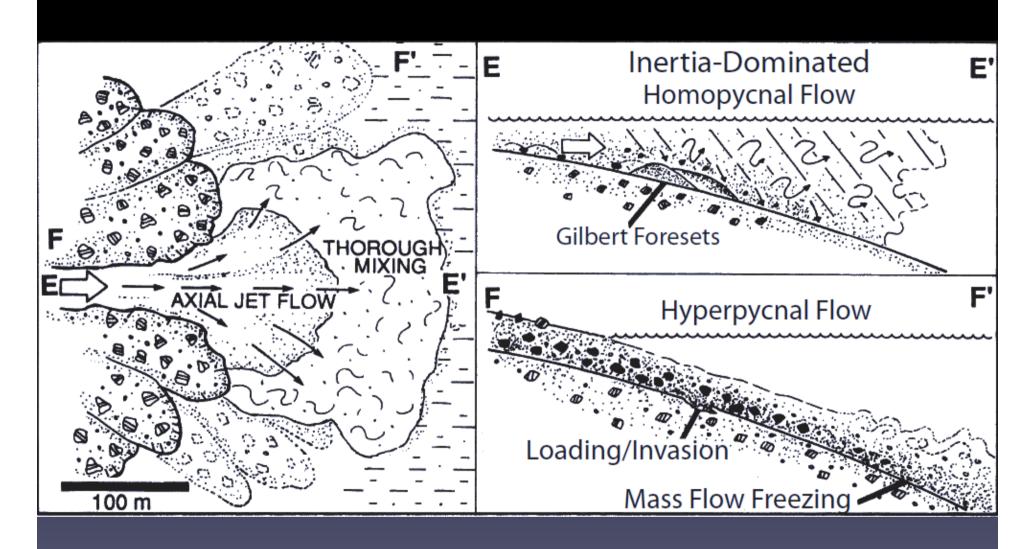
Mouth bar aggrades and progrades (up to 100 m!), forming y-shaped channel bifurcation



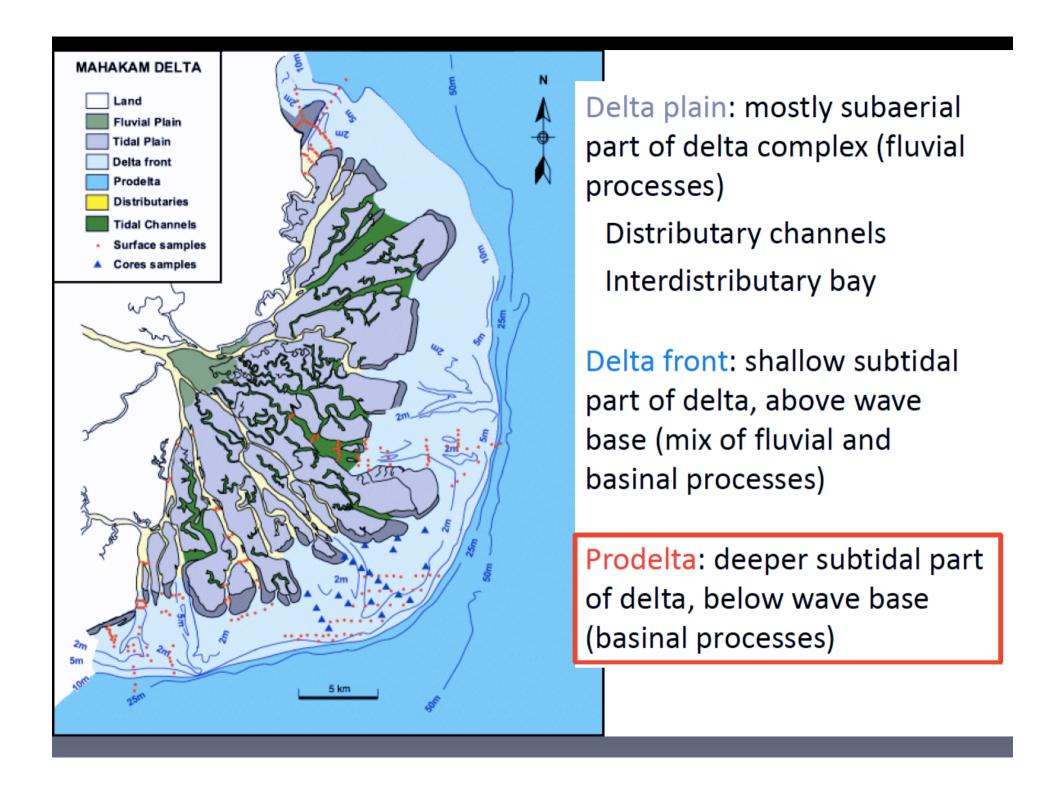


Examples of mouth-bar processes in river-dominated deltas (from Reading and Collinson, 1996, after Orton and Reading, 1993) incorporating ideas of Bates, 1953, Wright (1977) and others.

Inertia-dominated deltas (homopycnal flow)



Examples of mouth-bar processes in river-dominated deltas (from Reading and Collinson, 1996, after Orton and Reading, 1993) incorporating ideas of Bates, 1953, Wright (1977) and others.

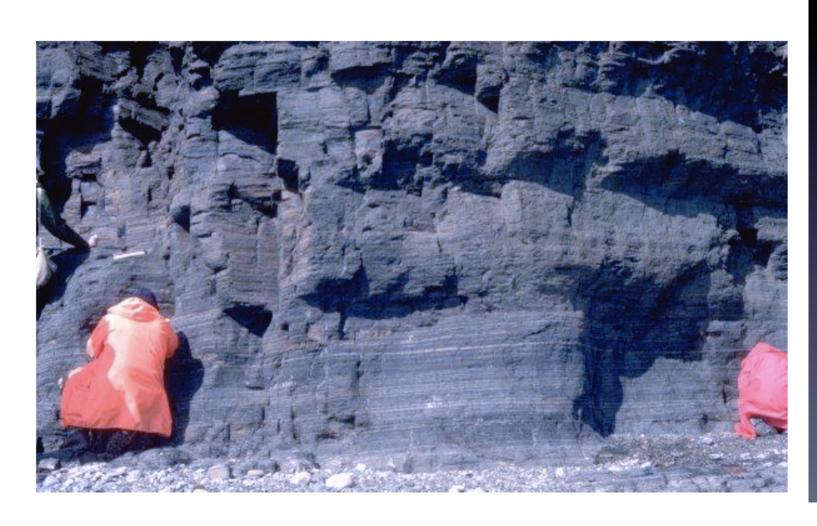


What features would you predict in prodelta sediments?

Grain size?

Sedimentary structures?

Bioturbation?



Hypopycnal plume

Density < seawater

Relatively continuous

Critical

Time

concentration

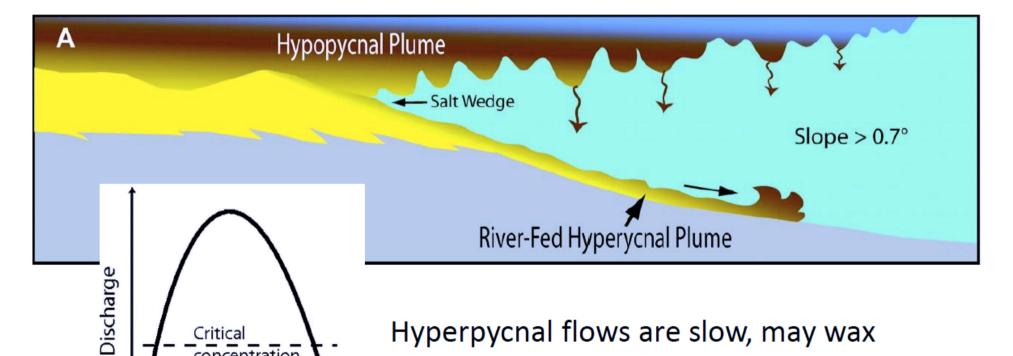
Deposition from suspension (=hemipelagic sedimentation)

Hyperpycnal plume

Density > seawater

Episodic, lasts hours-days

Deposition from suspension, modified by traction



Hyperpycnal flows are slow, may wax and wane during river flood stage

Deposits called "hyperpycnites"

Waxing/waning flow Inverse-normal grading



Hemipelagic sediments
Finely laminated or
bioturbated mud



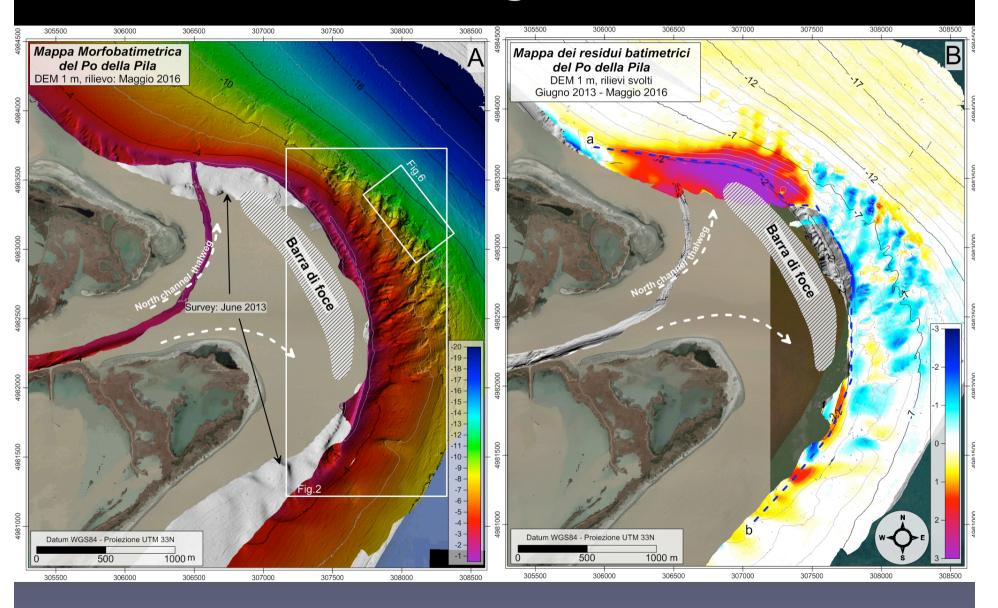
Suspension deposition Normal grading

Prodelta slopes are comparatively steep (and have rapid sedimentation rates)

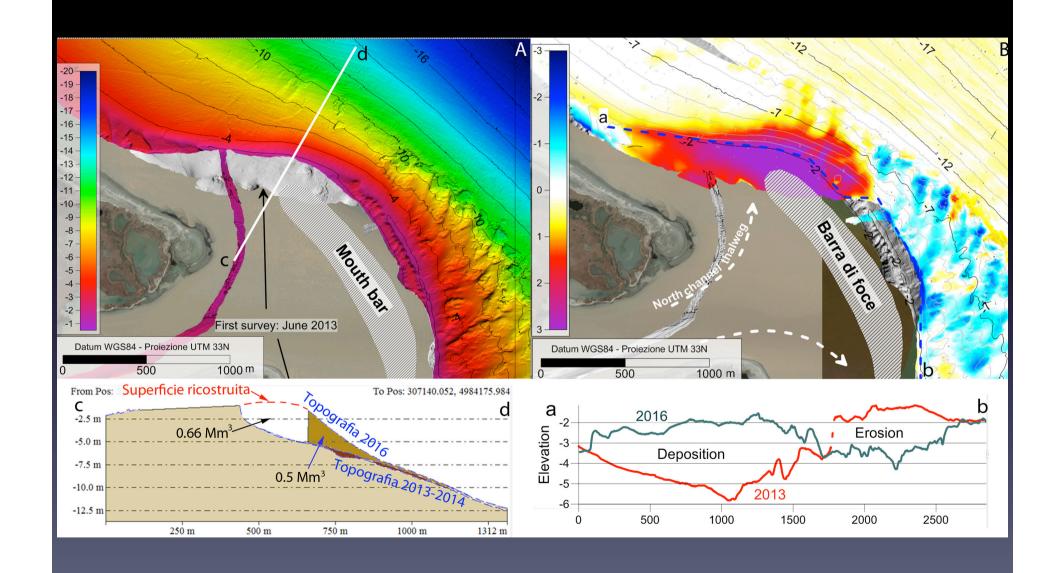
Slumps may be more common than in wave-dominated coasts



Instabilità gravitativa



DINAMICHE EROSIVO-DEPOSIZIONALI SHORT-TERM



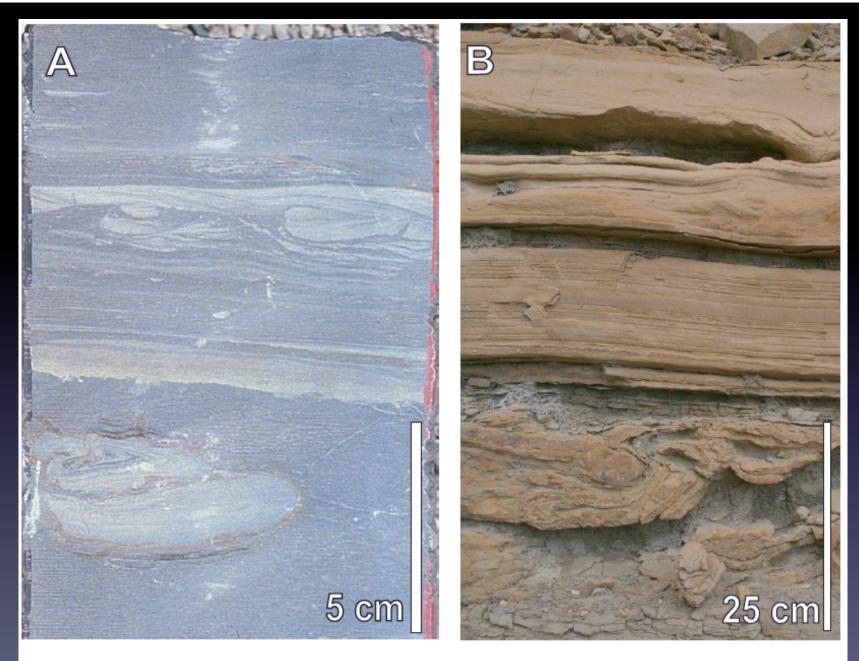
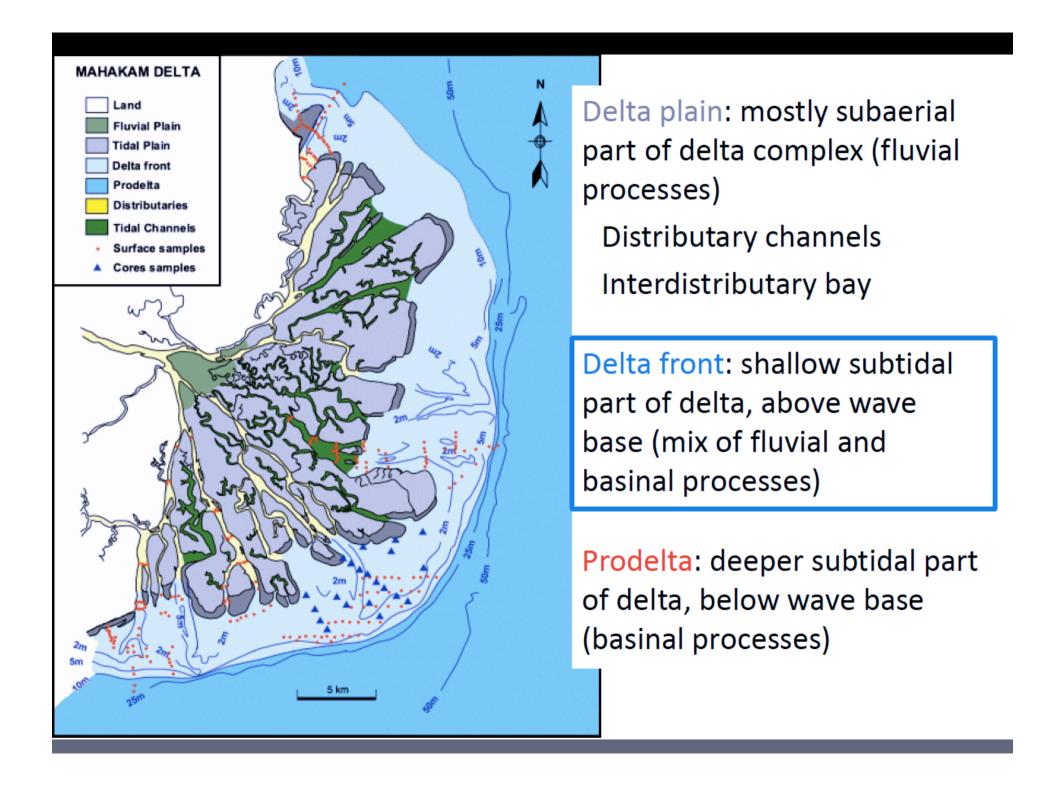


Fig. 29.—Deformation structures (load casts) in: A) prodelta mudstones of the Kavik Formation, Prudhoe Bay Field, Alaska, U.S.A.
B) Deformed sandstone bed overlain by parallel-laminated to rippled delta front splays interpreted as distal delta front, sediment gravity flow deposits, Cretaceous Ferron sandstone, Utah, U.S.A.

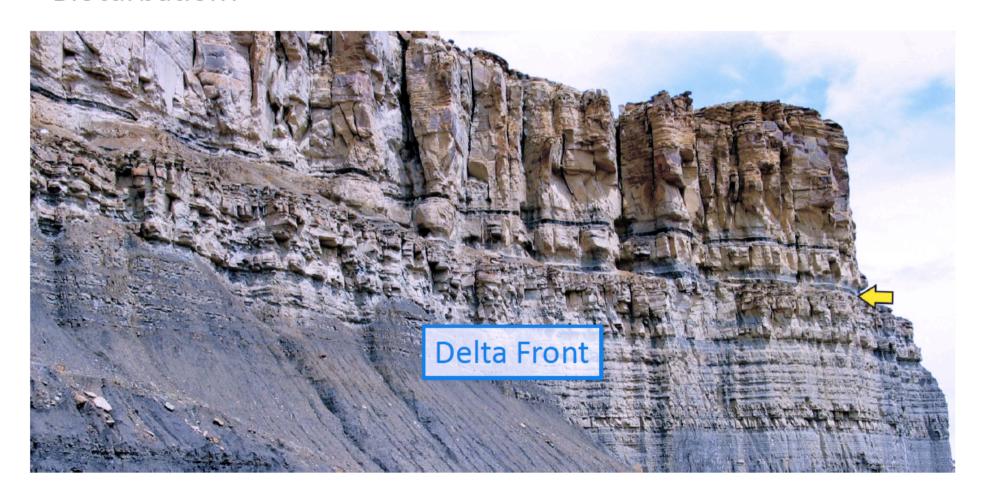


What features would you predict in delta front sediments?

Grain size?

Sedimentary structures?

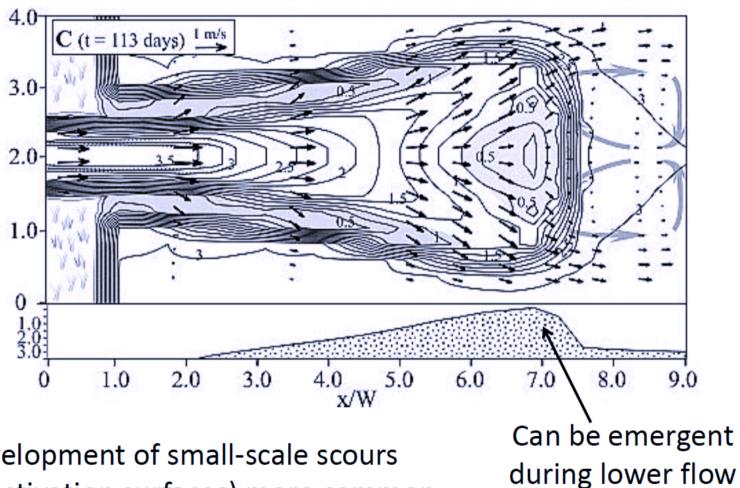
Bioturbation?



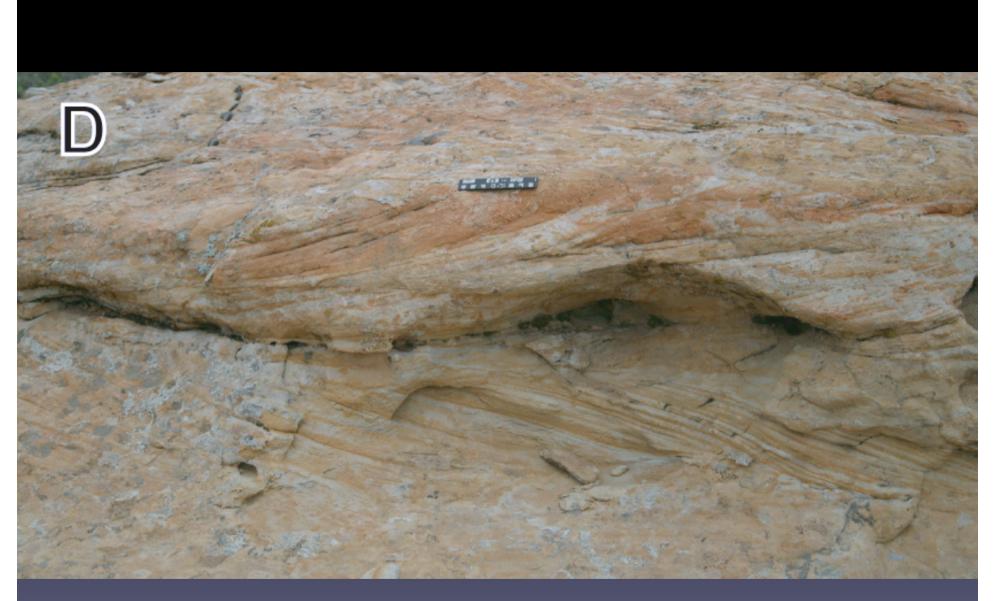
High sedimentation rates during river flood events



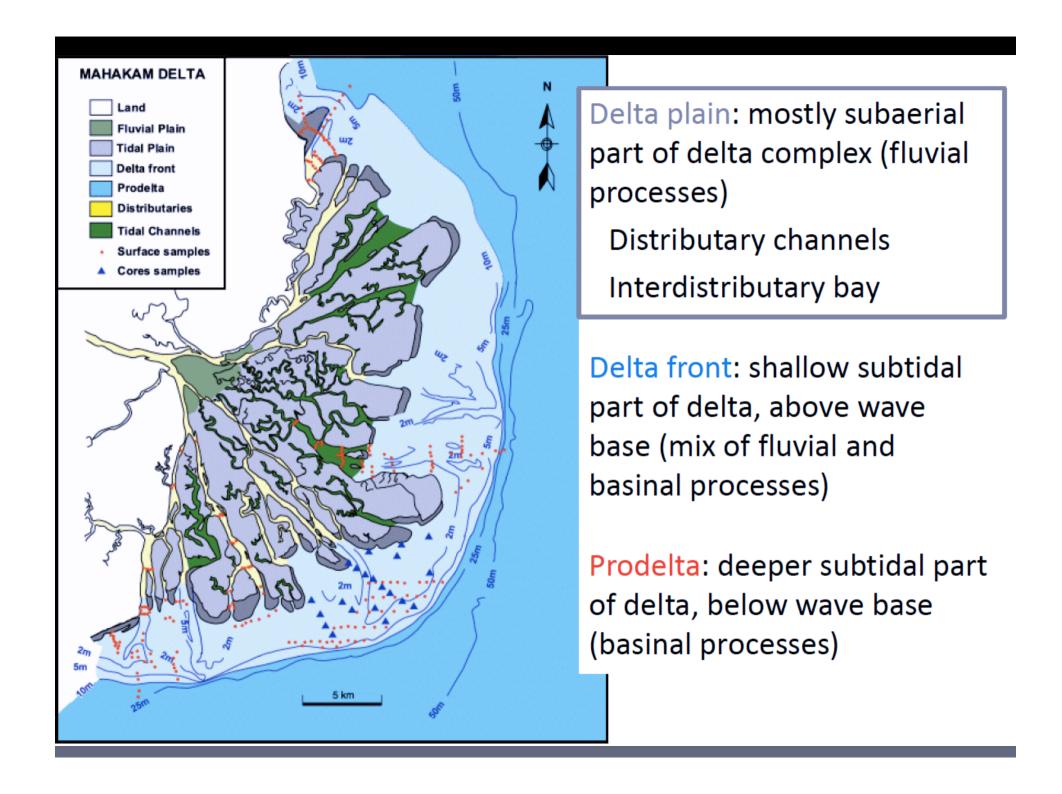
Mouth bar aggrades close to sea level during progradation



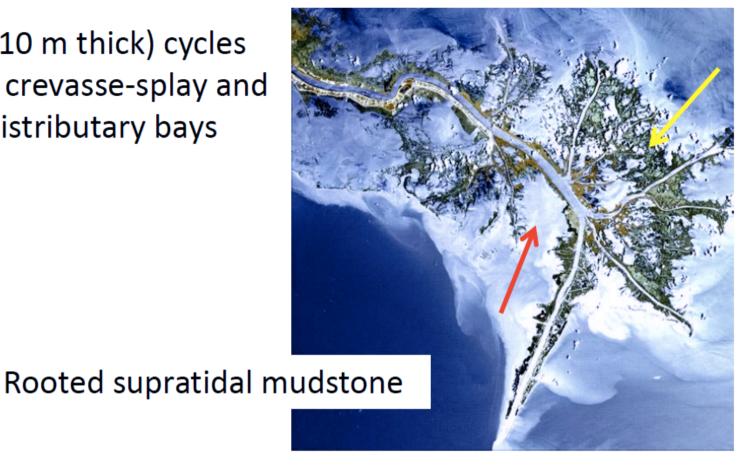
Development of small-scale scours (reactivation surfaces) more common near the top of the mouth bar



Bi-direction cross-bedding, upper shoreface/delta front



Small-scale (3-10 m thick) cycles formed due to crevasse-splay and filling of interdistributary bays



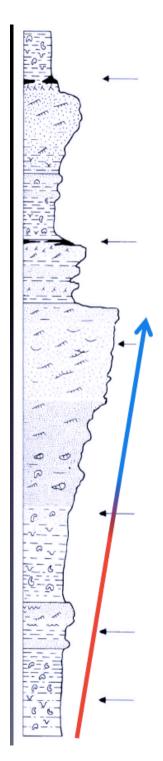
Increasing overbank sediment from river floods (rippled sandstone)

Interdistributary bay mudstone

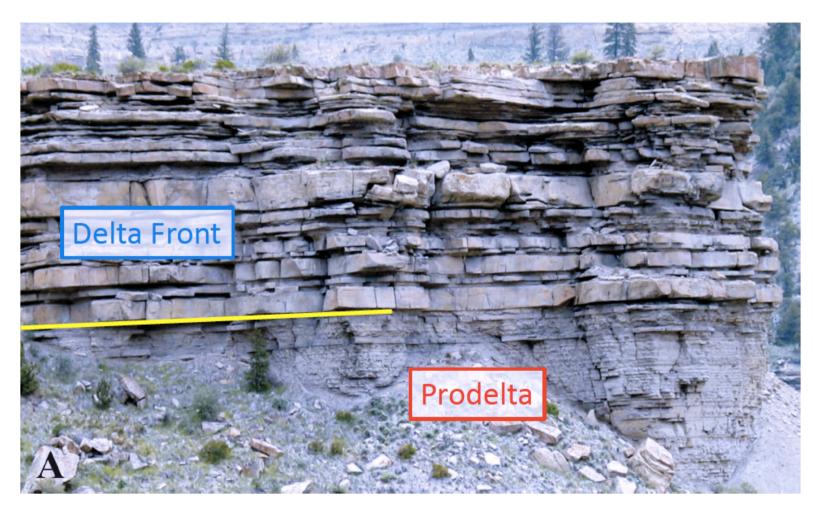
Transgressive lag

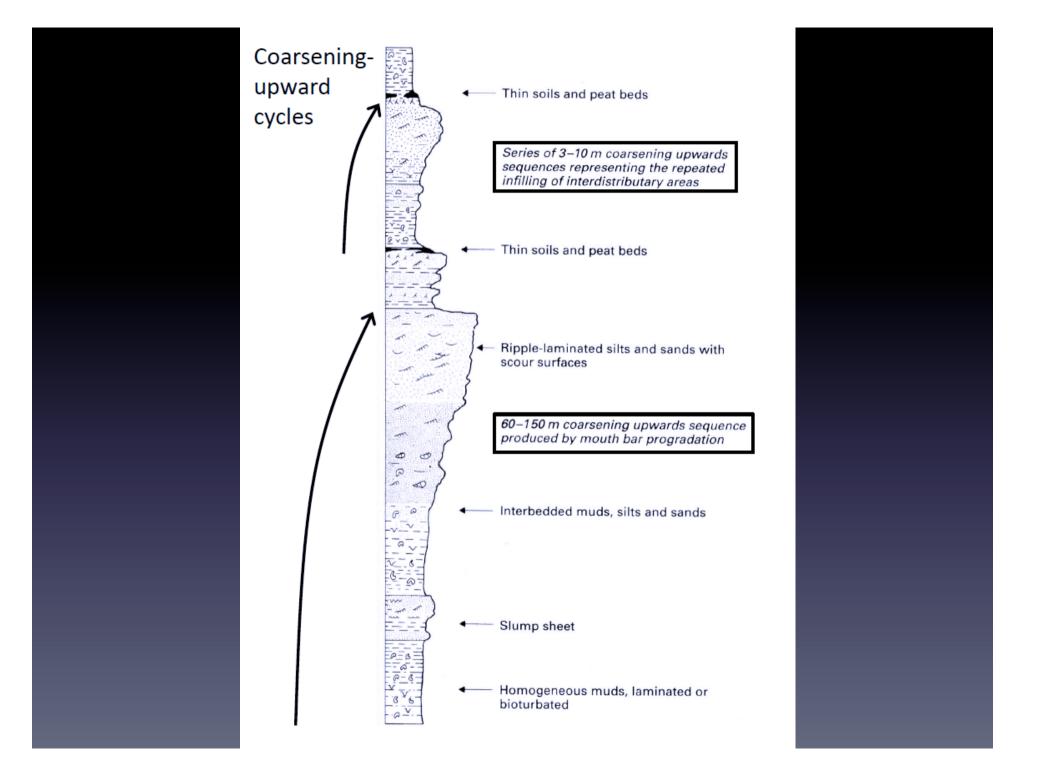


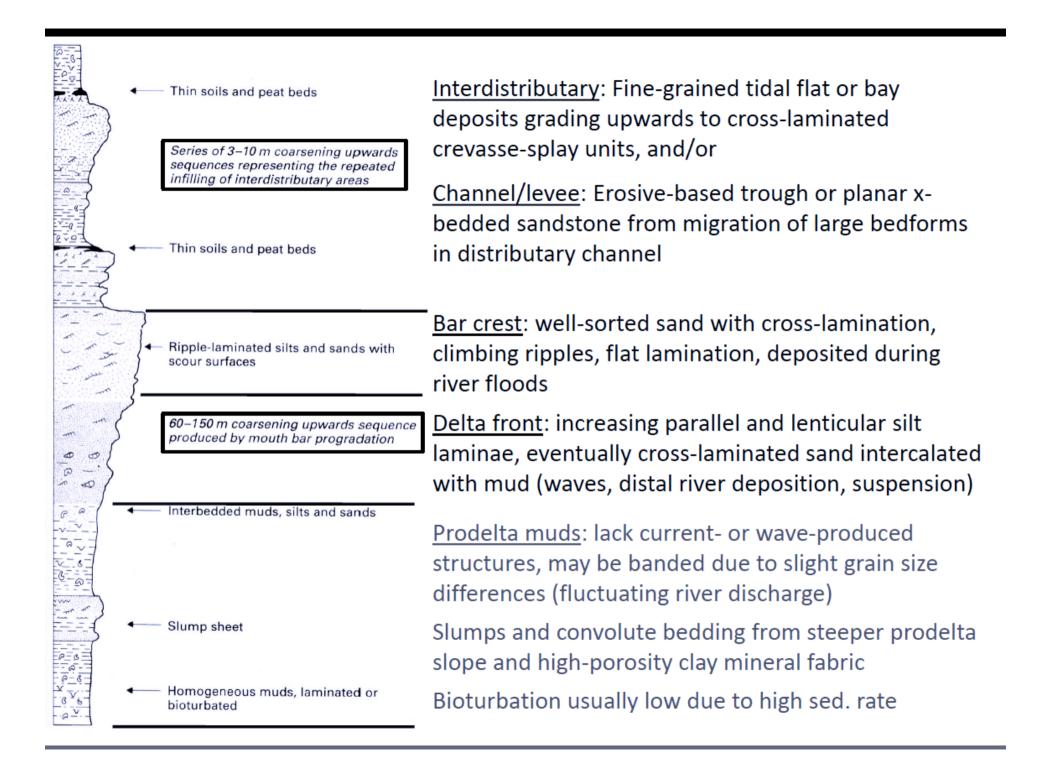
Delta Plain



Progradational delta facies coarsen upward from prodelta through delta front to (maybe) delta plain

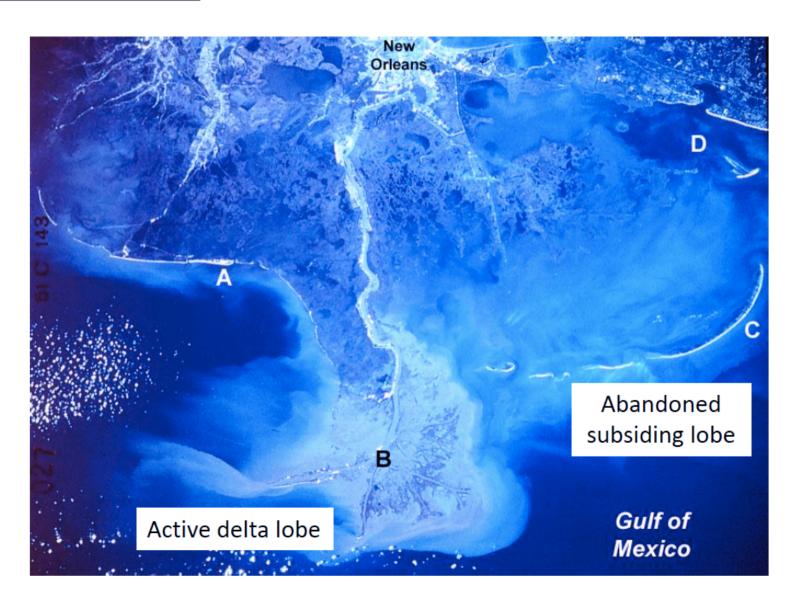




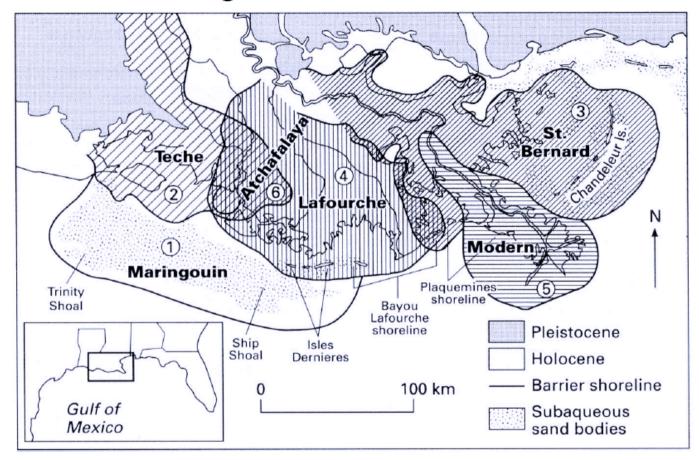


Cycles may be <u>autocyclic</u> (inherent due to delta processes) rather than <u>allocyclic</u> (due to external forces like base level)

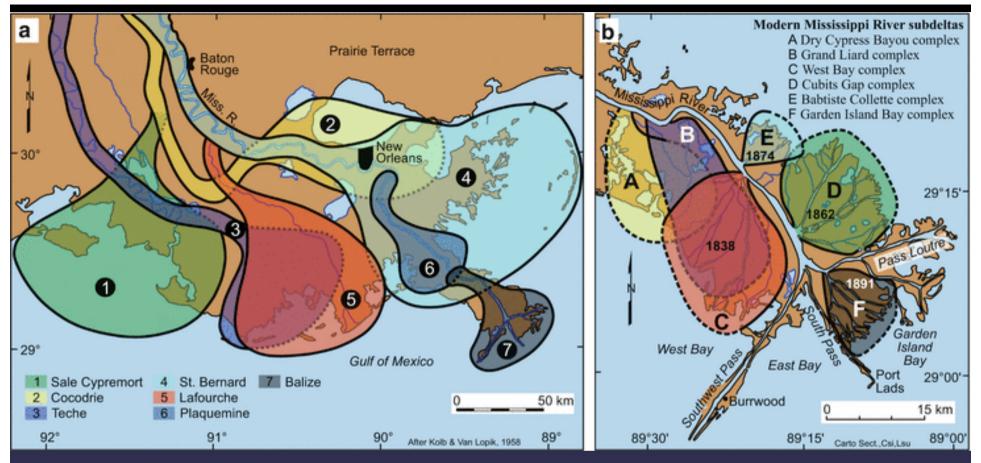
"Lobe switching"



Deltaic sequences cyclic at medium scale (50-150 m thick) due to lobe switching

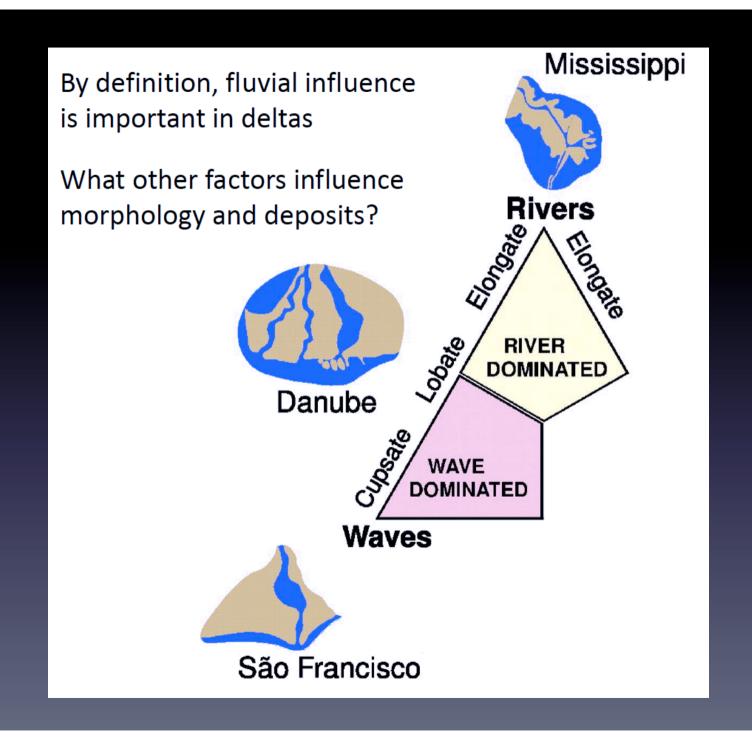


- 1. River builds long distributary channels on delta (progradation)
- 2. Avulsion occurs and river takes new, more favorable path to ocean forms new delta and abandons old lobe
- 3. Abandoned lobe continues to subside (facies retrogradation)

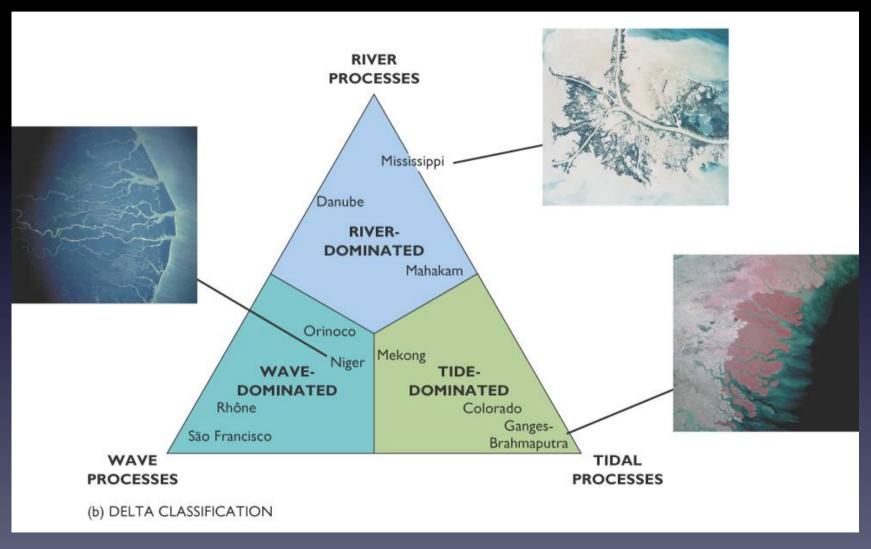


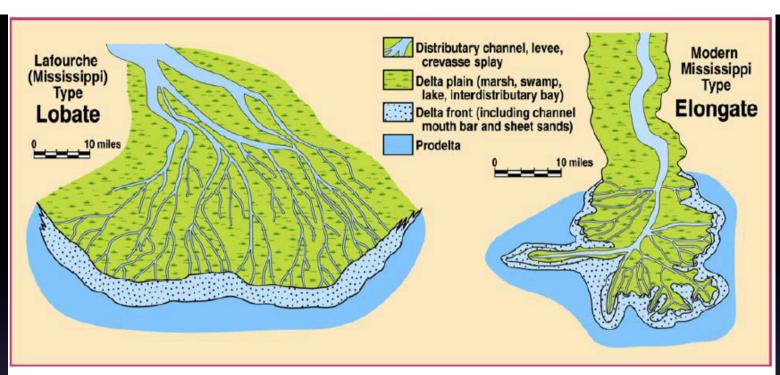
Sedimentation in the modern Mississippi delta. When the river water breaks through the levees, crevasse channels and splays are formed, which help to fill the areas between the channels (Coleman and Prior 1980).

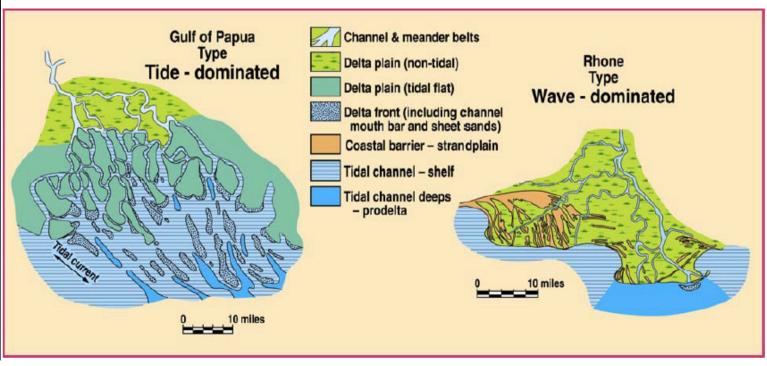
- (a) Delta lobes which show how the sedimentation has changed during the past 7,000 years. Each lobe of the delta appears to be active for 1,000–1,500 years (Coleman and Prior 1980).
- (b) Sedimentation in the last few 100 years.



DELTA Input sedimentari > tasso d'erosione







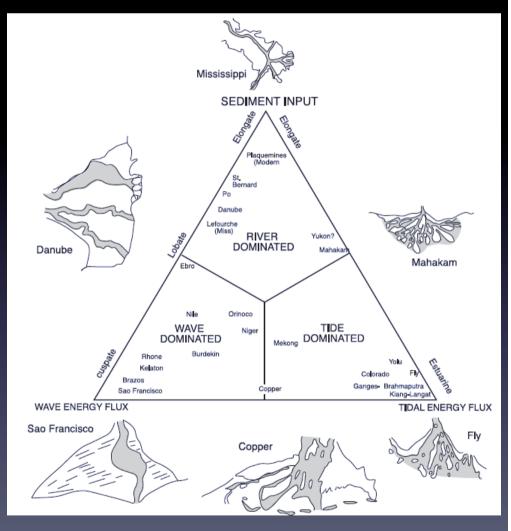
DELTA TYPES

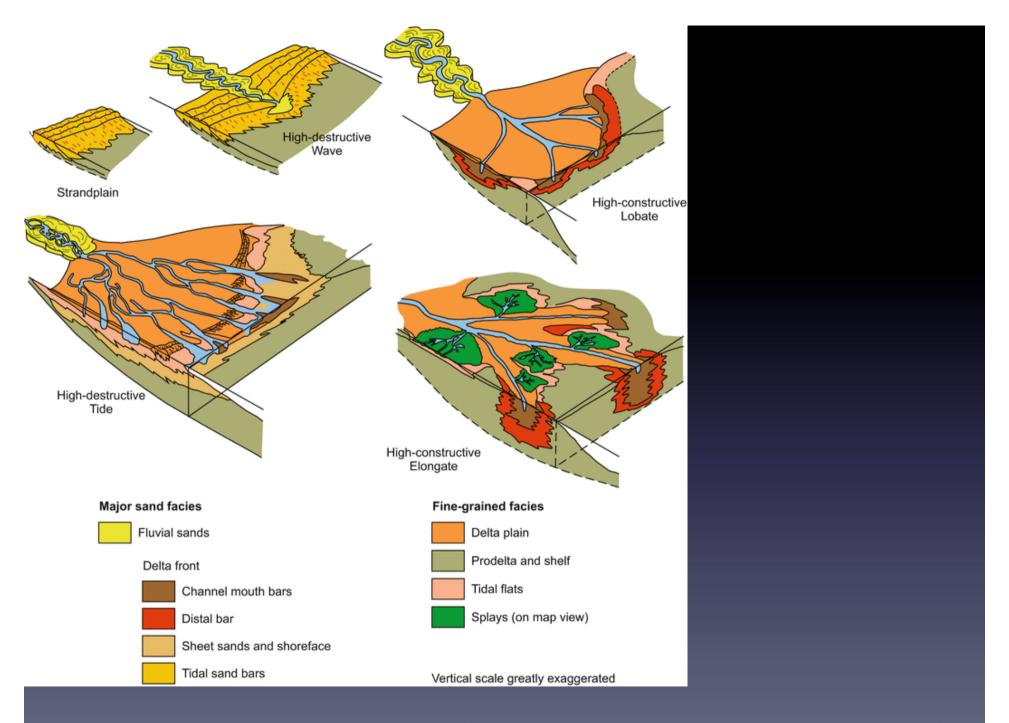
1. Constructional deltas

- Dominated by the fluvial system
- Strongly progradational/regressive
- Lobate Elongate

2. Destructional deltas

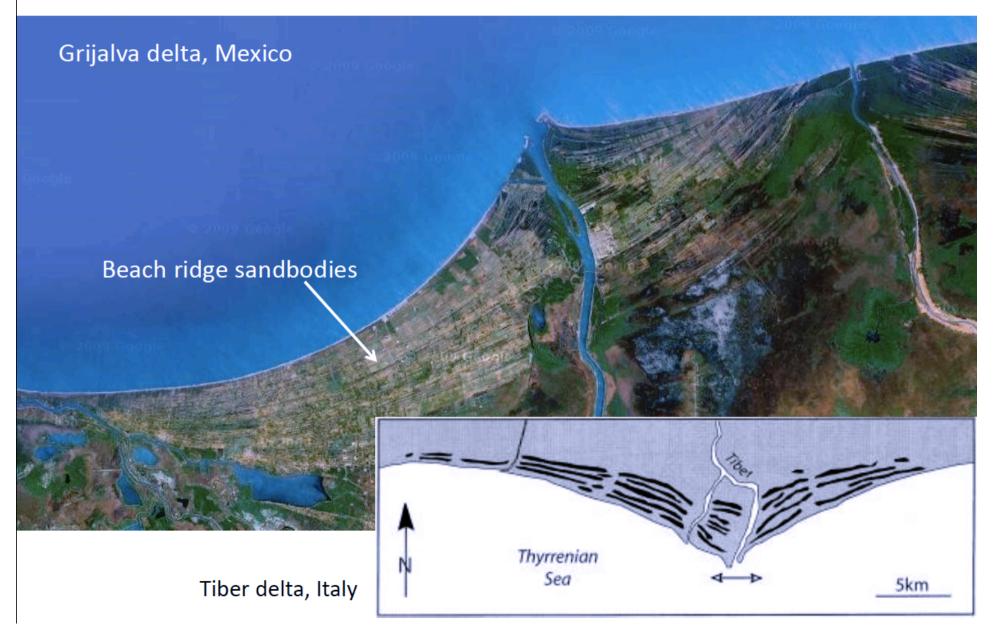
- Dominated by marine processes
- Common marine reworking with transgressive intervals
- Cuspate

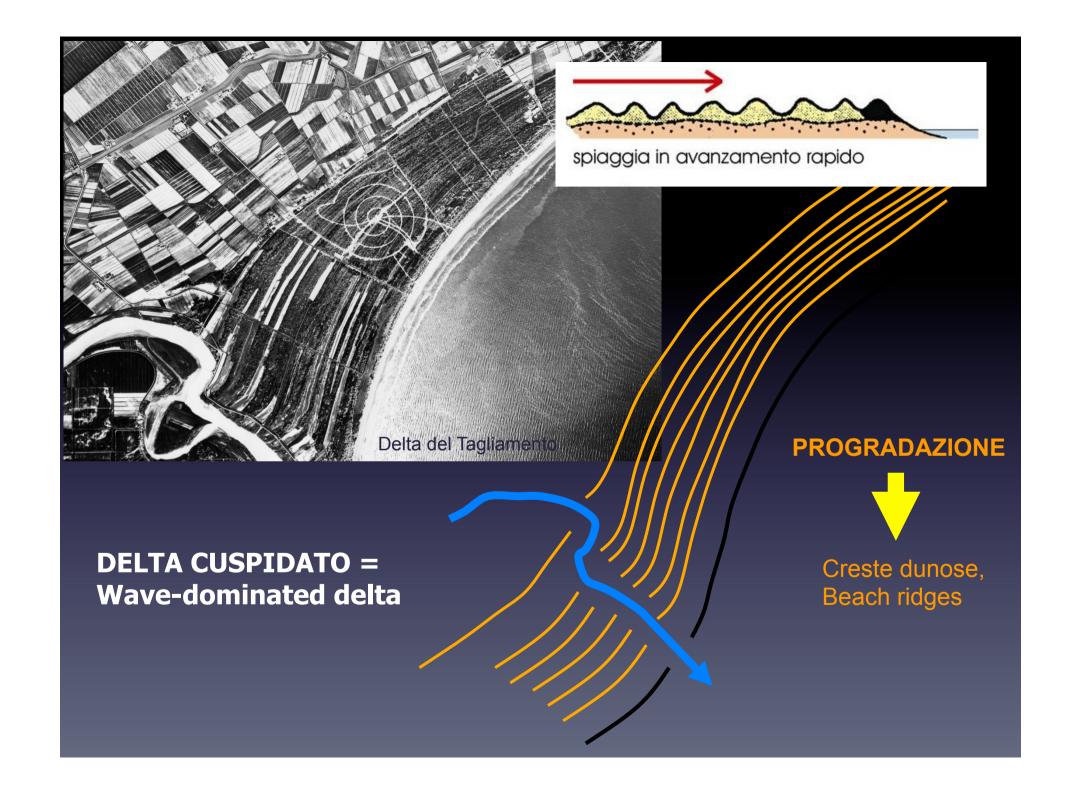




Bjørlykke K. (2015) Introduction to Sedimentology. In: Bjørlykke K. (eds) Petroleum Geoscience. Springer, Berlin, Heidelberg

Mouth bars modified by wave action into linear beach ridge sandbody complexes.





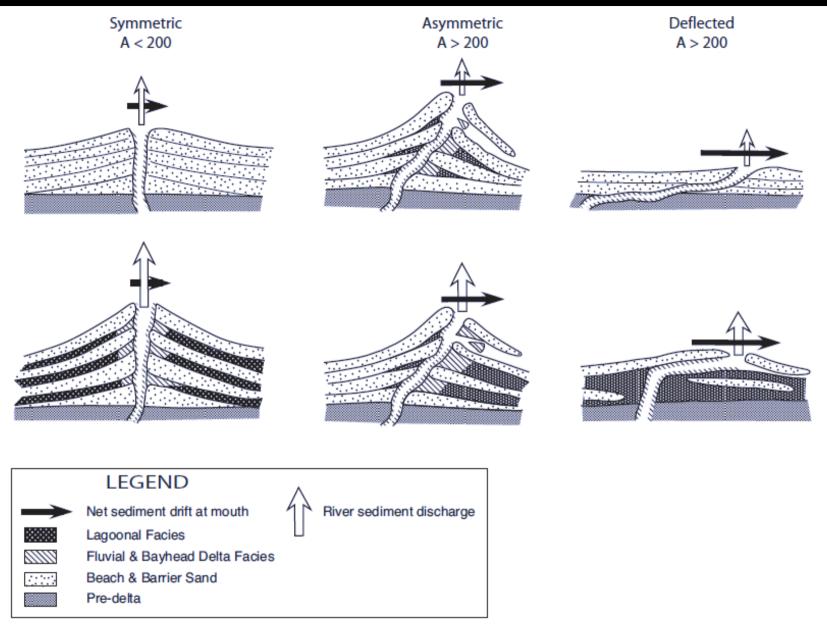
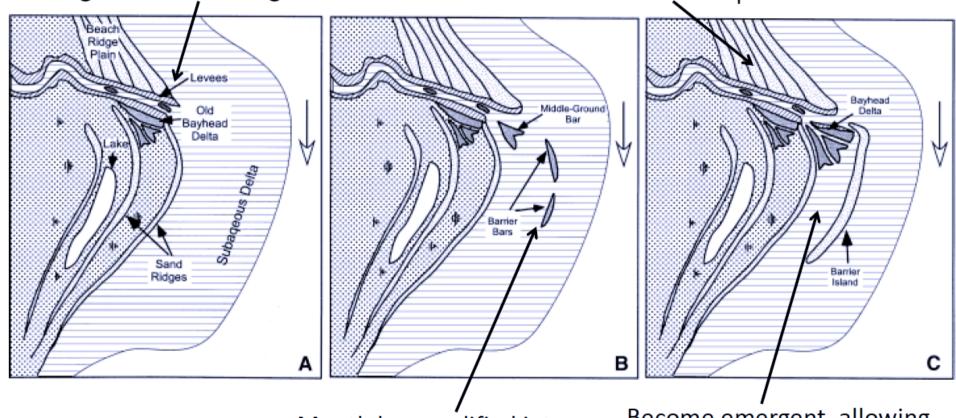


Fig. 6.—Morphology of wave-influenced deltas. Top row represents lower fluvial discharge compared to bottom row. River plume acts as a groyne that traps sediment updrift (after Bhattacharya and Giosan, 2003). Asymmetry index represents the ratio of fluvial sediment discharge to alongshore sediment transport rate.

When significant longshore current present, wave-influenced delta may be substantially asymmetrical

Prograding river levee acts as groin to block longshore drift

Wave-influenced strandplain shoreface forms on upcurrent side



Mouth bar modified into elongate barrier bars

Become emergent, allowing for fine-grained back-barrier deposition